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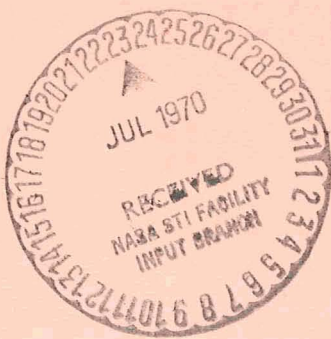
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U.S. DEPARTMENT OF COMMERCE
Environmental Science Services Administration
Weather Bureau

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Weekly Synoptic Analyses, 5-, 2-, and 0.4- Millibar Surfaces for 1967

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ESSA TECHNICAL REPORT WB 12

**Weekly Synoptic Analyses, 5-, 2-,
and 0.4-Millibar Surfaces for 1967**

(Based on meteorological rocketsonde and high-level
rawinsonde observations)

STAFF, UPPER AIR BRANCH,
National Meteorological Center
HILLCREST HEIGHTS, MD.
January 1970

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Weekly Synoptic Analyses, 5-, 2-, and 0.4-Millibar Surfaces for 1967

(Based on meteorological rocketsonde and high-level rawinsonde observations.)

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ABSTRACT

Data from meteorological rocketsonde and rawinsonde observations have been employed to analyze a series of high-altitude synoptic charts. Methods employed for processing the various types of data and the analysis procedure utilized are described.

Broad-scale analyses of the 5-, 2-, and 0.4-mb surfaces, primarily covering North America and adjacent ocean areas, are presented at weekly intervals for 1967. A brief discussion of the height and temperature fields is also given. A major stratospheric warming in December and a less intense event in January are among the discussion items.

INTRODUCTION

This volume represents the fourth in a continuing series of constant-pressure charts for the upper stratosphere, based on rocketsonde and very high level rawinsonde data obtained over North America and adjacent ocean areas. These charts have been analyzed by the Upper Air Branch of the Weather Bureau's National Meteorological Center (Staff 1967a, 1967b, 1969).

Computer-analyzed 10-mb charts (Environmental Science Services Administration 1967), based exclusively on rawinsonde data, are used as a foundation for hydrostatic computation of the height of the 5-mb surface. At 5 mb, both rawinsonde and rocketsonde data are used to derive the height and temperature fields. Above 5 mb, rawinsonde data are seldom available, and rocket observations constitute the basic data source. The Station Model and Reporting Rocket Stations chart (p. —) shows the locations of 19 meteorological rocket launch sites within the map analysis area for which data were available in 1967.

The present series of analyses at weekly intervals includes charts for the 5-, 2-, and 0.4-mb surfaces

(approximately 36, 42, and 55 km, respectively), nominally representing broad-scale synoptic conditions on each Wednesday of the analysis period. In addition to a description of the techniques employed in obtaining these charts, a brief discussion of large-scale features is also included.

PROCESSING OF RAWINSONDE DATA

The preparation of high-level data for analysis presents many problems. Some of the difficulties encountered in the use of rawinsonde information, such as achieving compatibility between daytime and nighttime observations, extrapolation and interpolation of data, and identification of erroneous reports, have been summarized previously (Finger et al. 1965, 1966).

Temperature and height adjustments designed to compensate for instrumental radiation errors were applied to all 5-mb rawinsonde reports. Initially, an adjustment was made which reduced values from daytime observations to the level of those values reported from nighttime observations. The magnitudes of these day-night adjustments, which are intended to account

for the effects of solar heating on the rod thermistor, were determined with the aid of a computer program which calculated monthly mean differences between reported daylight and darkness values. Input to this program consisted of all available 5-mb data for 1964 and 1967 from stations in North America and adjacent ocean areas that employ U.S. outrigger-type radiosonde instruments (for details of the adjustment and further references, see McInturff and Finger 1968).

Theoretical and laboratory studies of the rod thermistor used in present-day radiosondes indicate that a significant error may be induced by infrared cooling at stratospheric levels above about 10 mb. Therefore, a second adjustment (Barr 1966) was added to all reported 5-mb data, including observations made actually in darkness and daytime data that had been adjusted for solar-radiation error. The magnitude of the temperature correction is a function of reported temperature, while the height adjustment varies linearly with the temperature correction.

Rawinsonde data utilized for the 5-mb charts were processed by computer methods. Inputs to the computer program consisted of observations from North America and adjacent ocean areas for the 7-, 5-, 4-, and 3-mb levels in the form of punched cards obtained from the National Weather Records Center. Output listings included all observations for 1 week at each station. Data for levels other than 5 mb were also listed to supply the analyst with as much supplementary information as possible. Thermal winds and corresponding horizontal temperature gradients were computed for the layers from 7 to 5, 7 to 4, and 5 to 4 mb.

After processing, the rawinsonde-derived temperature, height, wind, and thermal wind were plotted on working charts. For charts presented in this publication, the rawinsonde data have been omitted for the sake of legibility.

PROCESSING OF ROCKETSONDE DATA

Winds and temperatures measured by rocketsonde systems comprised the basic data sample for analyses at the 2- and 0.4-mb levels, and were also used in combination with rawinsonde data for the 5-mb charts. The rocketsonde information was obtained for this project in the form of magnetic tape, identical in format to published information contained in the *Meteorological Rocket Network (MRN) Firings* (World Data Center A 1968).

The procedure for extraction of rocketsonde information and for computation of heights of the 5-, 2-,

and 0.4-mb surfaces was as follows:

a. Wind and temperature profiles were processed by a computer program, designed to reduce the amplitude of rapid oscillations in the vertical by using a simple (1-2-1) smoothing technique. This processing was done because analyses are intended to portray only the broad-scale stratospheric features.

b. Temperatures were modified above 40 km in accordance with Drews' (1966) publication, one of several recent theoretical studies (see the review article by Ballard 1967). For the Arcasonde 1A, the instrument most widely employed during 1967, the correction applied to the temperature profiles assumed a representative fall rate and increased from 1°C at 42 km to about 9°C at 58 km.

c. An initial estimate of the geopotential height for each station at the respective pressure levels (5, 2, and 0.4 mb) had to be made before the temperature for that particular level could be selected. This estimate was obtained by deriving the temporal trend of the geopotential height for the station from the lower level analyzed charts and by applying this trend to the geopotential height for the station as shown by the previously analyzed chart. The temperature at this estimated height was then extracted from the temperature profile.

d. The true geopotential height of the pressure surface at each station was then hydrostatically computed for each available temperature profile. This height was obtained by deriving a layer-mean temperature between the extracted temperature of the upper level and the temperature obtained from the previously completed lower level analysis. For the 5- and 2-mb levels, a linear lapse rate was assumed to be representative. Because a temperature maximum (associated with the stratopause) usually exists between 2 to 0.4 mb, mean temperatures were derived by using values extracted at each whole kilometer between the two pressure levels.

e. Wind components were extracted from the smoothed profiles at the height of the constant-pressure surface, and a resultant wind was computed.

f. Thermal winds were determined for 6-km layers surrounding each pressure surface at the reporting station. An experimental method of smoothing the wind profiles was used to derive thermal winds for the January and February analyses. This method of fitting a first-, second-, or third-order polynomial to the data proved quite useful for eliminating scales of motion smaller than those being analyzed and led to generally consistent computations of the horizontal temperature gradients.

The rocketsonde data derived (temperature, height, wind direction and speed, and thermal wind) were plotted on working charts. On the charts presented for publication, three observations are shown when available for each station. Heights and thermal winds have been omitted for the sake of legibility. The Station Model and Reporting Rocket Stations chart illustrates the symbols used to distinguish the data obtained on Wednesdays from off-time data.

ANALYSIS PROCEDURE

Conventional techniques, including differential analysis methods, were utilized to construct the 5-, 2-, and 0.4-mb charts. However, because of the sparsity of observations, data acquired for the entire week were examined to determine the synoptic changes that took place during the week, and thus conditions pertaining to Wednesday.

The analysis system consisted of the following steps:

a. Isotherms were derived with the aid of processed rocketsonde and rawinsonde temperatures. In spite of the corrections applied to the 5-mb rawinsonde temperature data, large differences could sometimes be seen when comparing rocketsonde data that were close in time and distance. Generally, the problem could be resolved by giving greater weight to rawinsonde stations employing hypsometers. In a few cases, the difficulty even occurred with hypsometer radiosondes; therefore, some compromise value had to be chosen, usually closer to the rocketsonde-reported temperature. Computed thermal winds were especially useful in determining horizontal temperature gradients and the relative location of warm and cold areas. Where possible, time-height sections of temperature were plotted as a further aid in deriving the isotherms.

b. A mean temperature field representative of the layer between the previously analyzed lower surface, and the selected surface was derived graphically. This mean field represents a geopotential thickness which, when added to the lower level height field, yields a smooth, conservative first approximation to the contour pattern at the upper surface.

c. Reported winds and computed heights for individual stations were then used to adjust the first approximation of the contour field, using the assumption of geostrophic flow. Winds were accorded the highest priority for this adjustment.

d. The analysis was checked for vertical and temporal consistency. For example, centers of systems, ridges, and troughs were examined with the aid of all

available data to verify vertical slope and movement with time. Time-height sections and height-change charts were especially useful for these purposes.

The above procedures appear to produce excellent results at 5 mb, and were successfully applied to obtain the 2- and 0.4-mb charts. Generally, only slight adjustments of the first-approximation height fields were necessary at the 5- and 2-mb levels. However, rather formidable analysis problems were evident at the 0.4-mb level. Obviously, the most severe problems arise from the sparsity of observations. The geographical area of analysis for the charts has thus been reduced in accordance with the available data. Another difficulty was the apparent occurrence of large day-to-day temperature changes at times exceeding 10°C (Miller 1969), and persistent oscillations in many wind profiles. In some cases, deviations of reported temperatures and winds from the fields derived by differential analysis techniques could be accounted for by identifiable synoptic changes. Occasionally, it was impossible to make a reasonable reconciliation of station values with the values determined by differential analysis.

A further analysis problem arises from the apparent intersection of the stratopause with the 0.4-mb level. Because the normal stratospheric temperature inversion ceases at the stratopause level, the graphical method for obtaining mean temperature, which depends on the existence of a linear profile, is no longer valid. Large adjustments must be made, especially at lower latitudes, in the graphically derived height field to conform with the computed height at each station (as discussed in the previous section).

An additional problem is encountered in the summer charts. During that season, the circulation at the upper surfaces (2 and 0.4 mb) would be expected to follow the pattern established for lower levels, that is, a rather uniform easterly flow about an anticyclone centered at or very near the North Pole. However, on a typical summertime 0.4-mb chart, most reported rocketsonde winds exhibit significant southerly components. If these winds were to be given full weight in the analysis, the resulting pattern would consist of contours, oriented from southeast to northwest, spiraling toward a high center located apparently over northern Europe.

The prevalence of positive meridional components in summertime rocketsonde winds has been noted previously (Miers and Beyers 1964). Recent studies (Reed et al. 1966), based on MRN data for several summers have demonstrated that the meridional wind component resulting from the diurnal tide reaches

maximum southerly strength at about noon, local time. Since most MRN firings occur near local noon, the measured winds naturally contain this component. Although no adjustments were made in individual reports, the analysis procedure includes suitable compensation. Differences in the reported winds and orientation of the contours may be quite pronounced, especially on the summertime charts. Strong westerly winds of winter usually mask the weak diurnal component, but there is a possibility that this component may contribute to some bias in analysis during periods of weak circulation.

Although careful consideration of high-level data allows a broad-scale depiction of circulation patterns up to 0.4 mb, the sparsity of reports requires an increasing amount of subjectivity as the analysis proceeds to this high level. As yet, the analyst has little in the way of synoptic models for guidance with respect to the probable contour and isotherm patterns and the phase relationship between these patterns in areas of sparse data. The justification for some analyses depends on the interpretation of the limited amount of data in such a way as to portray a coherent sequence of synoptic events. In spite of these factors, surprisingly little alteration in the principal features of the circulation and temperature distribution shown in the final analysis can be made without inordinately violating some of the data. In general, the contours and isotherms depicted are felt to be good approximations to the flow patterns at these levels. Even so, the same degree of accuracy that is found customarily in the analysis of charts at lower levels should not be expected.

A contour interval of 320-geopotential meter (gpm) was used throughout the year. In addition, intermediate dashed contours were used to outline areas of relatively weak gradient, especially during the spring and fall changeover periods. Isotherms were drawn and labeled at 5°C intervals.

DISCUSSION OF THE 1967 CHARTS

The principal features of the circulation during 1967 may be seen in figures 1 through 5 which illustrate the analyzed height and temperature values extracted from the weekly charts at the location of five representative rocket launch sites. From these figures, we can see the large changes that occurred and also infer the cycle of events during the course of the year. The annual trend toward maximum height and temperature during the summer and toward minimum during the winter is clearly visible in figures 1 through

4. Largest ranges in values are seen at the more northerly locations. The pronounced sinusoidal temperature curves, most noticeable at 5 mb, are considerably flattened at 0.4 mb (Johnson and Gelman 1968). Although not clearly shown by this graph, Antigua, B. W. I. (fig. 5), exhibits the semiannual variation associated with a tropical station. The major anomalies of the winter period are seen in the curves as these disturbances affect the various locations. Outstanding departures from the seasonal trend occurred in January when a minor warming affected the upper stratosphere, and in December when the first indications of a major stratospheric warming could be seen. Perturbations seen in February, March, and November were associated with the intensification and movements of the Aleutian anticyclone, with events in April leading to the final spring change to summer conditions.

The following discussion will concentrate on a brief description of the major synoptic events of 1967 and will show how the charts depict them.

In January, the circulation was dominated by a cold cyclone near the North Pole with anticyclonic activity apparent in the Aleutian region and the western Atlantic Ocean. The isotherm patterns, especially at 2 and 0.4 mb, were not oriented parallel to the contours, suggesting that strong dynamic changes were taking place.

As an aid to the readers, it may be instructive to review the various segments of information used by the analyst to determine the circulation features on typical charts; for example, the January 4 charts. The basic isotherm pattern at 5 mb was determined by several rocketsonde reports at each station and by the approximately three times as many radiosonde reports received from stations located primarily on the North American continent. Off-time reports were used as an aid in determining conditions shown in the analysis.

In addition to the cold trough extending from Greenland over western Canada to the mid-Pacific, two warm areas were also apparent. A report from Bermuda, together with thermal wind reports from Wallops Island, Va., and Cape Kennedy, Fla., indicated the location of warm air off the U.S. east coast. Another warm area in eastern Siberia was determined largely on the basis of the computed thermal winds at Fort Greely, Alaska. This latter location, however, agrees well with analyses at lower levels where more data were available. The contours were then determined, with very little adjustment needed, in accordance with the method previously discussed. It is apparent that all available data fit into a coherent pattern

for the determination of the temperature and height field at 5 mb.

At the 2- and 0.4-mb levels, only rocketsonde data were available, so that added emphasis had to be placed on the thermal winds to outline the temperature fields. The thermal winds indicated a sharp constriction of the cold trough between the warm centers, with a system of cold air over the Gulf of Alaska and another in the polar regions. The very strong thermal gradients shown on these charts may have been difficult to justify solely on the basis of the single cold report at Fort Greely on January 3 and from the computed thermal winds. However, substantiating evidence appeared the following week when rapid temperature changes were reported by all western stations. This reinforced the pattern of intense cold and warm centers which moved rapidly during the period.

The situation up to this point strongly suggested that a major stratospheric warming would take place. Warming was noted over the North Atlantic Ocean and the Alaskan region; the contour and isotherm patterns appeared to be oriented in a hemisphere pattern of wave two (two warm anticyclones diametrically opposed and an elongated polar low with isolated cold air in the middle latitudes). In spite of the intense temperature changes at the upper stratospheric levels, the temperature fields below 5 mb were relatively unaffected. Because the height fields were only slightly changed and westerly circulation remained strong, this event in January was termed a minor stratospheric warming.

During the last half of February, a significant circulation change occurred as the Aleutian anticyclone moved eastward over western North America. Warm air associated with this system was detected northeast of the center. By March 22, the anticyclone had receded to lower latitudes while warm air had moved steadily toward the polar basin. This latter movement resulted in a reversal of thermal gradient by the end of March—warm air dominated the North Pole and relatively cold air was located in the midlatitudes. The height field slowly responded to the thermal changes, but westerly circulation remained evident until the second week in April. At that time, the anticyclone redeveloped in the Aleutian region, moved over Alaska, and gradually displaced the polar cyclone. As the warm polar anticyclone intensified in response to the springtime radiational changes, the westerlies associated with the cold trough moved southward. By the beginning of May, the westerlies had disappeared at 0.4 mb; at the end of the month,

easterlies were predominant over the entire analysis area.

While the higher latitudes of the Northern Hemisphere were under the influence of circumpolar easterlies which increased in intensity with altitude, the light and variable winds at Fort Sherman, C. Z., indicated a more cellular structure in the tropical latitudes. This structure can be explained at least in a qualitative way partly as the coupling mechanism between the dominant easterlies of the Northern Hemisphere and the westerlies of the Southern Hemisphere (Miller and Finger 1969).

The warm summertime polar anticyclone reached maximum temperature and intensity at the beginning of July. Thereafter, this system began a slow decay as solar radiation over the northern latitudes declined. By August 23, a cold core cyclone had formed near the North Pole at 5 mb. This system gradually intensified, displacing the remnants of the polar anticyclone (in the form of cellular ridges) southward. By the second week of September, easterly winds were still apparent at the middle latitudes, but general circulation features were relatively complex. For instance on September 27, the westerlies at 2 mb, as reported by Fort Sherman, indicated a trough was north of the station, as were the dominant cyclone and subtropical ridge. At 5 mb and 0.4 mb, however, this tropical trough was south of the station.

The polar region cooled and the vortex intensified in October, producing a stable-appearing circumpolar westerly circulation. At 0.4 mb, however, the cooling in polar regions was interrupted by relatively warm air in the region north of Fort Greely. Charts for other periods have shown rather clearly that temperature variations may occur in the upper stratosphere without any significant change in the parameter even at middle stratospheric altitudes. It was not until the middle of November that the entire layer from 5 to 0.4 mb was affected.

On December 6, the very warm report from Fort Greely at 0.4 mb may have been an indication of the major stratospheric warming that occurred during the following weeks (Johnson 1969). Rapid movement in the temperature pattern occurred so that by December 13, warm temperatures were reported at West Geirinish, Scotland, and Cape Kennedy, and the calculated thermal winds at all eastern stations indicated that at 0.4 mb there was a very large area of the Northern Hemisphere under the influence of warm air. This rapid change becomes more evident when viewed in the context of the sequence of events at lower levels. The warm air seen in the eastern Atlantic region, for

instance at 5 mb on December 20, was part of a large warm area around Eurasia, extending sometimes to Alaska. Relatively small position changes in this warm area, with the warm air extending into the Aleutian region one time and into the North Atlantic region the next, could have accounted for the rather radical change apparent at 0.4 mb from December 6 to December 13.

Very strong gradients in the height field were also seen during this time period, with a wind of over 350 knots at the 47-km level reported on December 13 at West Geirinish. During succeeding days, the effects of the warming were observed at lower altitudes. By December 20, a very warm region was seen at 5 mb

over Europe, with rapidly intensifying anticyclones in the Atlantic and western Pacific Oceans. Warming in some areas was as much as 70°C during that week. The Atlantic anticyclone moved northward as the warm area expanded and intensified. By the end of December, the temperature gradient between the polar latitudes and midlatitudes had reversed throughout the entire middle and upper stratosphere. This remarkable event constituted the earliest (by about a month) major warming of the stratosphere ever recorded. Consequences of this stratospheric warming carried into the beginning of the following year, with major circulation changes occurring in January 1968. (This discussion prepared by Melvyn E. Gelman.)

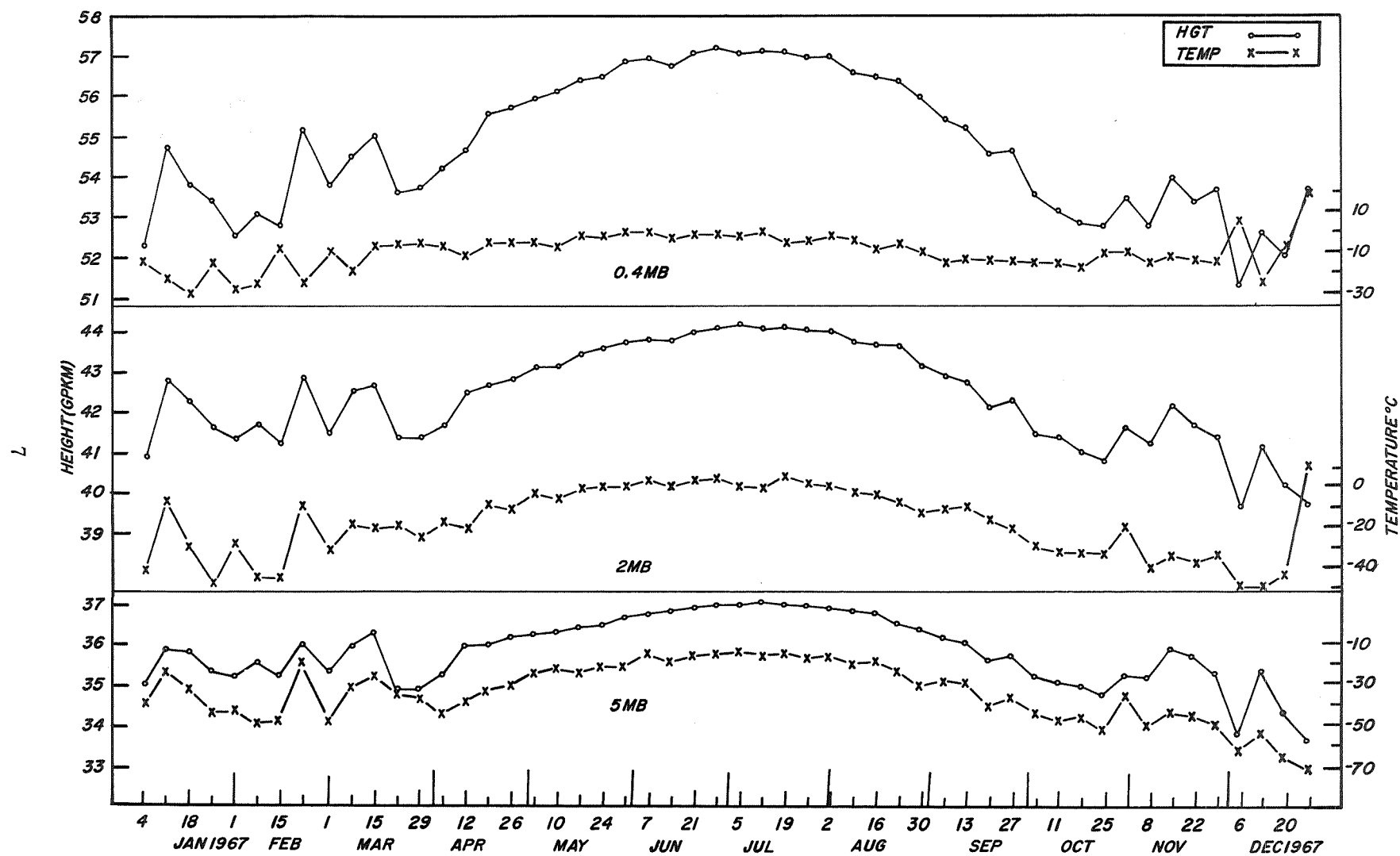


FIGURE 1.—Fort Greely, Alaska ($64^{\circ}00'N, 145^{\circ}44'W.$), analyzed values extracted from weekly 5-, 2-, and 0.4-mb charts.

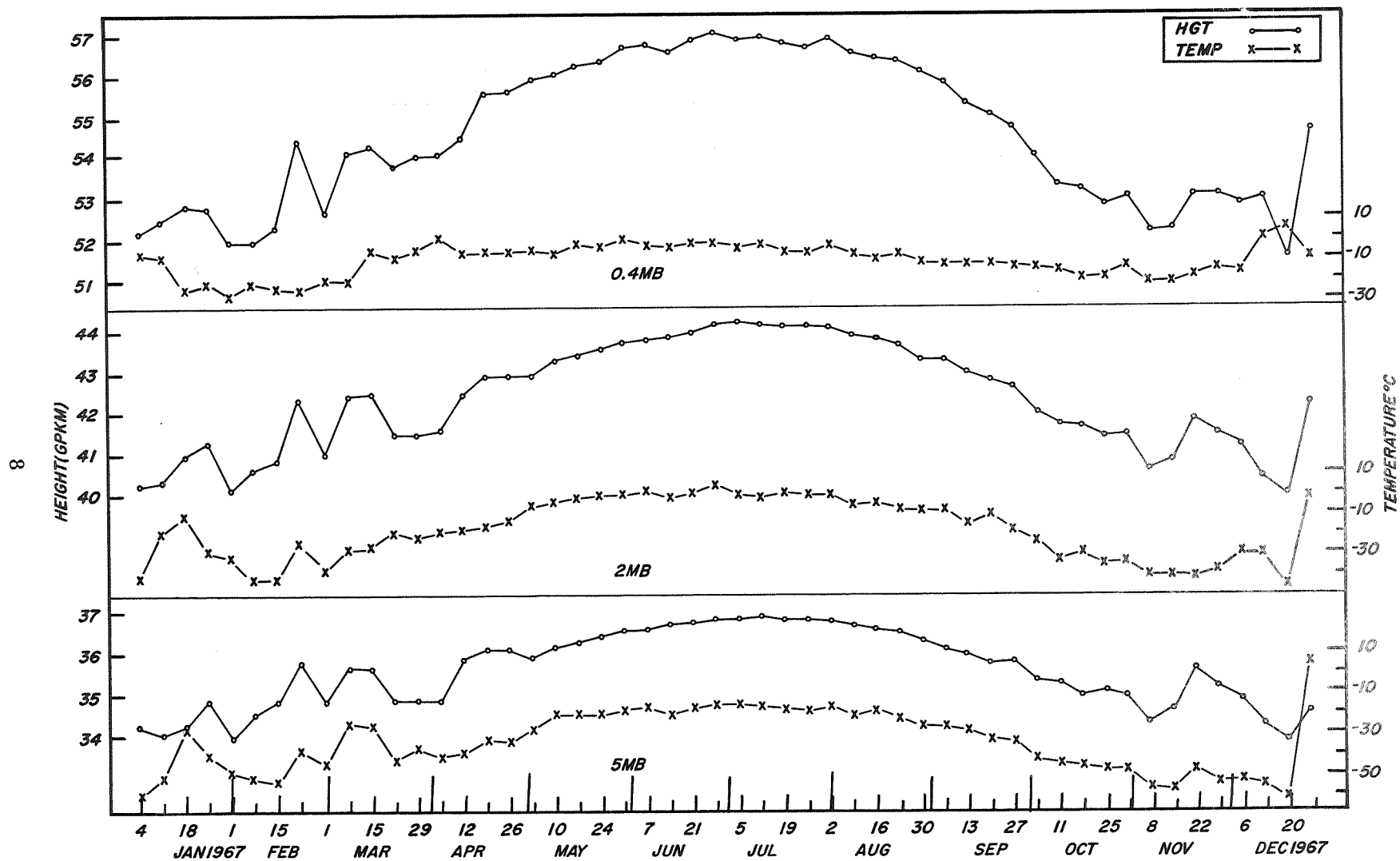


FIGURE 2.—Fort Churchill, Canada ($58^{\circ}44'N$, $93^{\circ}49'W$.), analyzed values extracted from weekly 5-, 2-, and 0.4-mb charts.

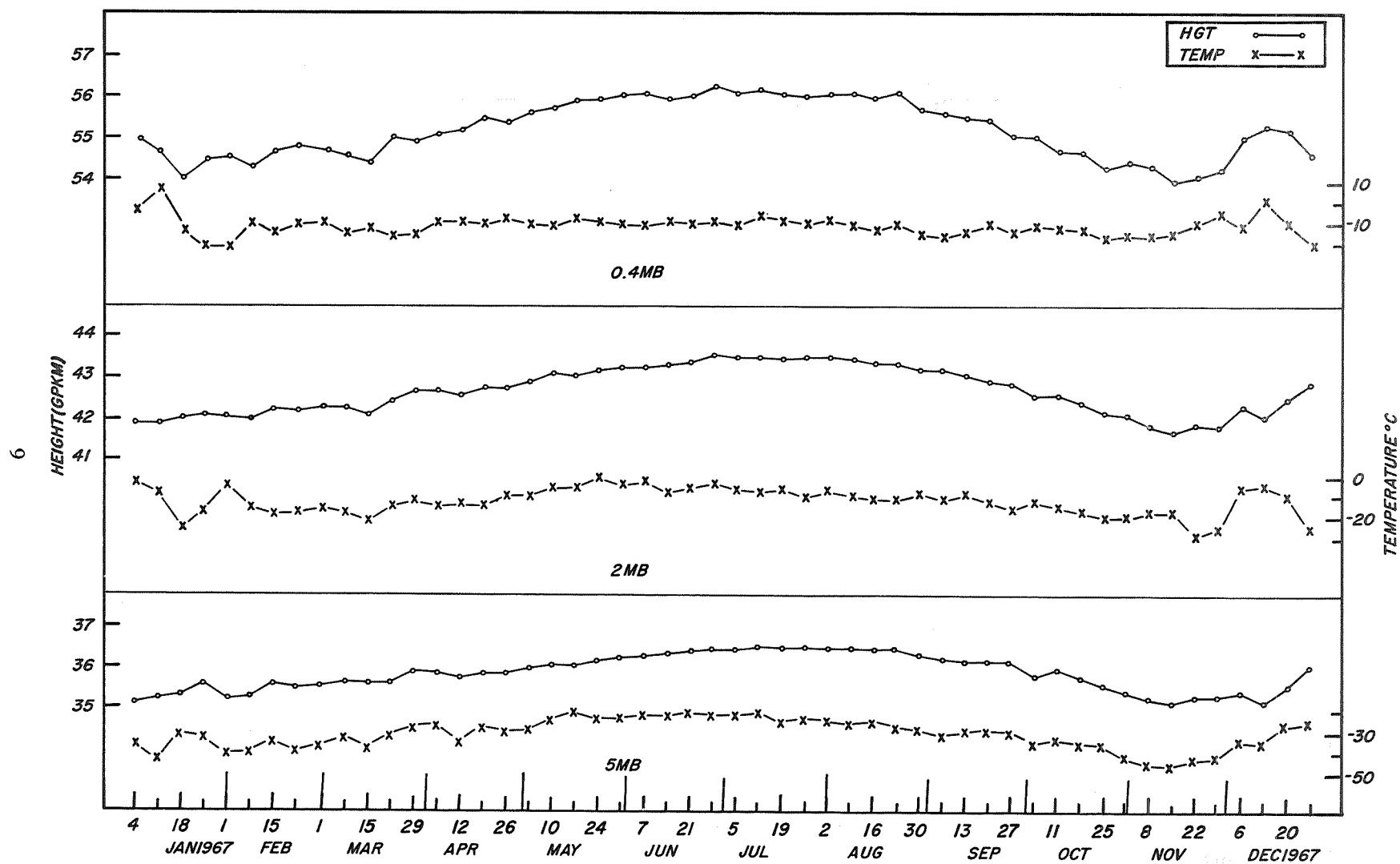


FIGURE 3.—Wallops Island, Va. ($37^{\circ}50'N$, $75^{\circ}29'W$.), analyzed values extracted from weekly 5-, 2-, and 0.4-mb charts.

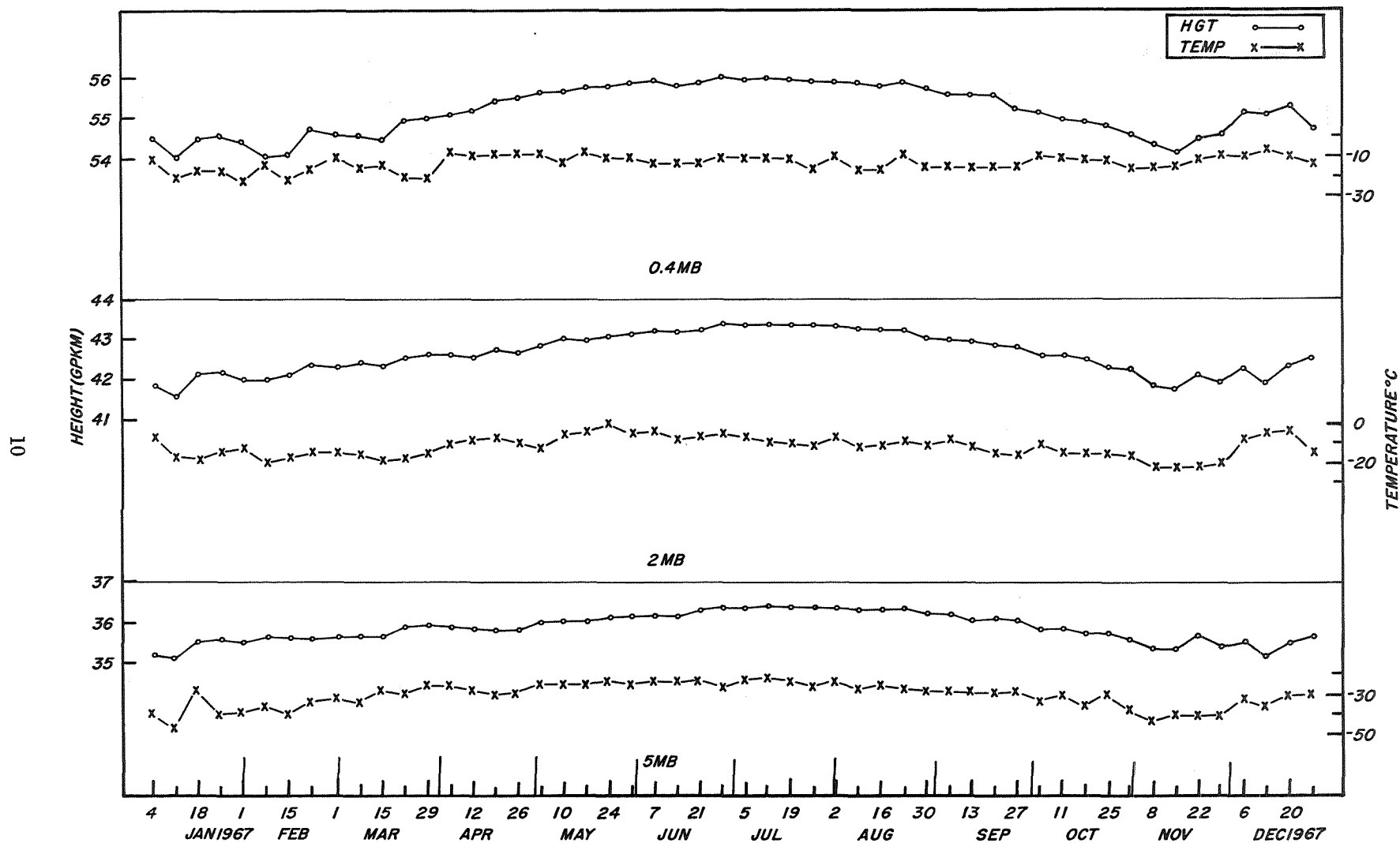


FIGURE 4.—White Sands, N. Mex. (32°23'N., 106°29'W.), analyzed values extracted from weekly 5-, 2-, and 0.4-mb charts.

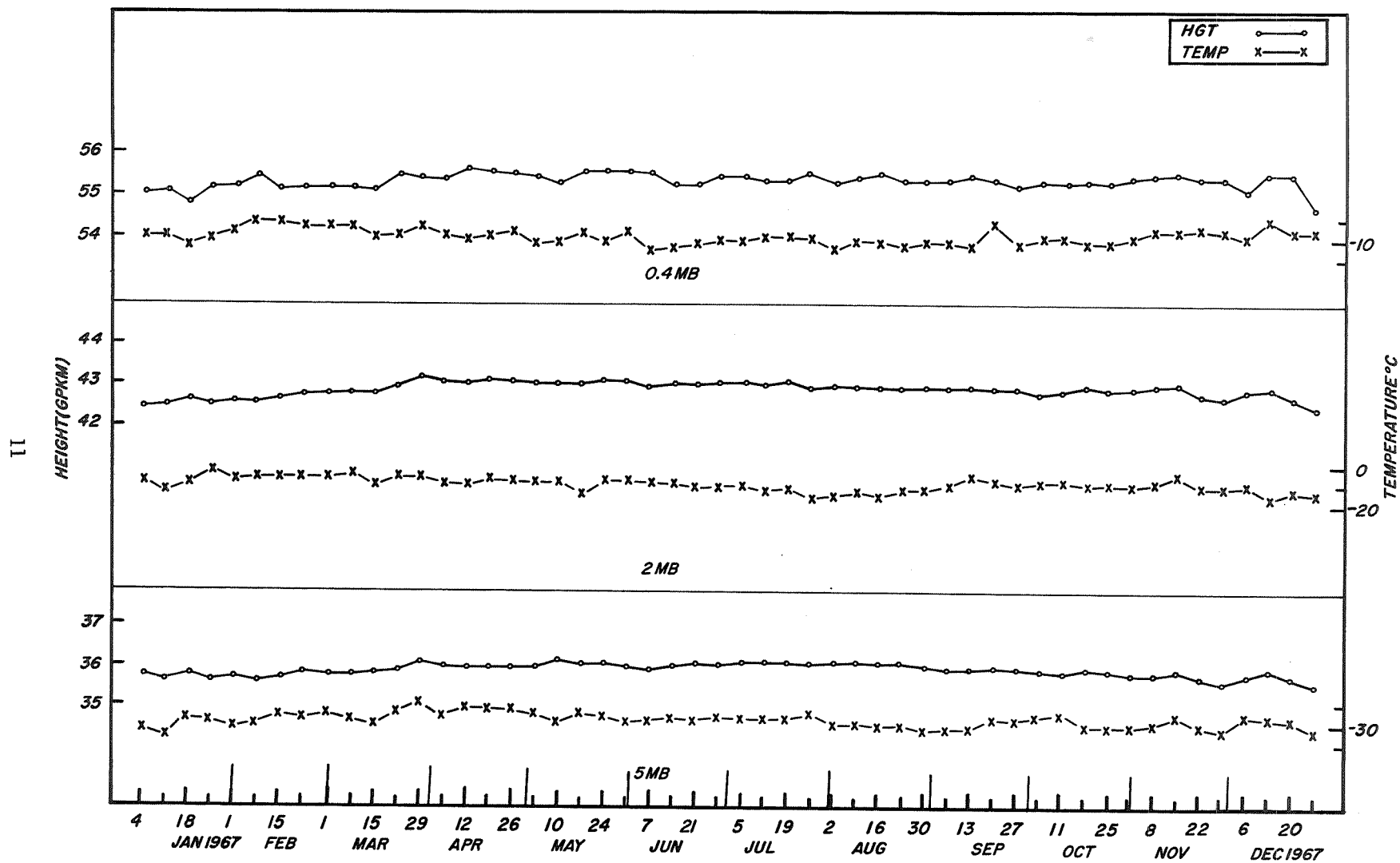


FIGURE 5.—Antigua, B.W.I. (17°09'N., 61°47'W.), analyzed values extracted from weekly 5-, 2-, and 0.4-mb charts.

ACKNOWLEDGMENTS

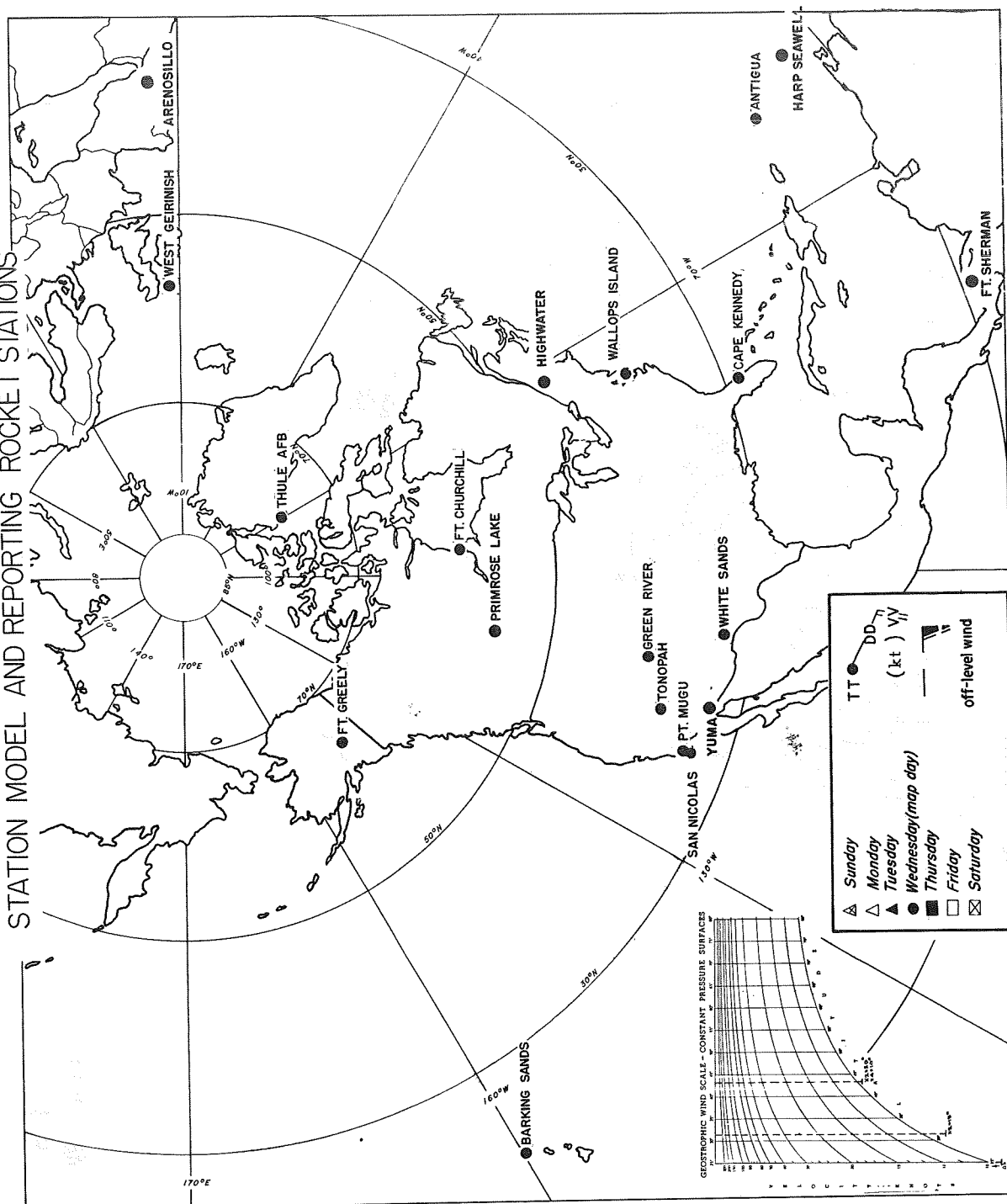
Funds for the analysis of the 1967 chart series were provided by the National Aeronautics and Space Administration under contract P-40346 (G).

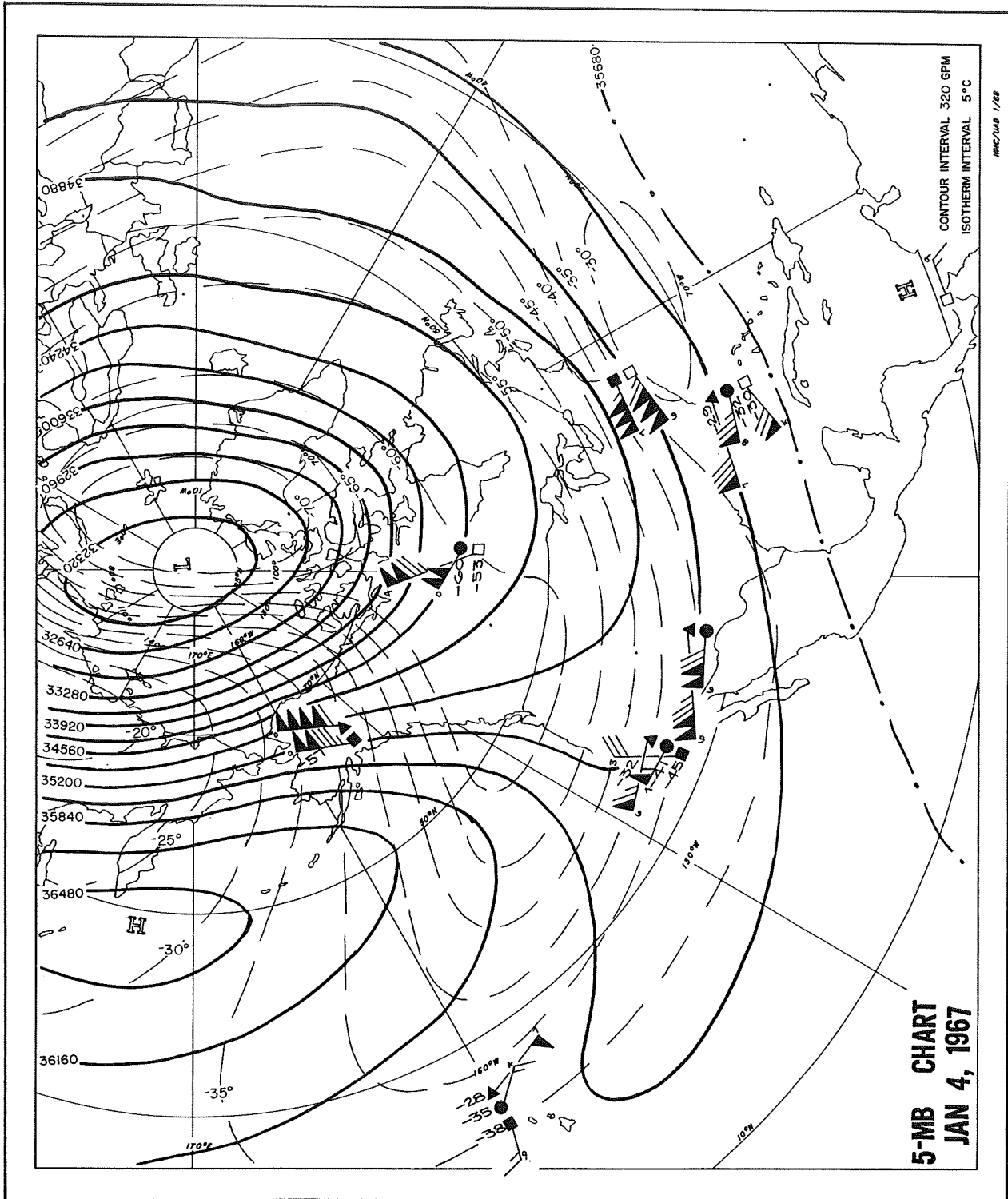
Weather Bureau personnel actively engaged in the analysis project were: Project Manager, Frederick G. Finger; Project Analyst, M. E. Gelman; Research Consultant, A. J. Miller; Data Processing, L. P. Manello, J. A. Booth, and D. L. Griffith; and Computer Operations, K. W. Johnson and W. P. Townshend.

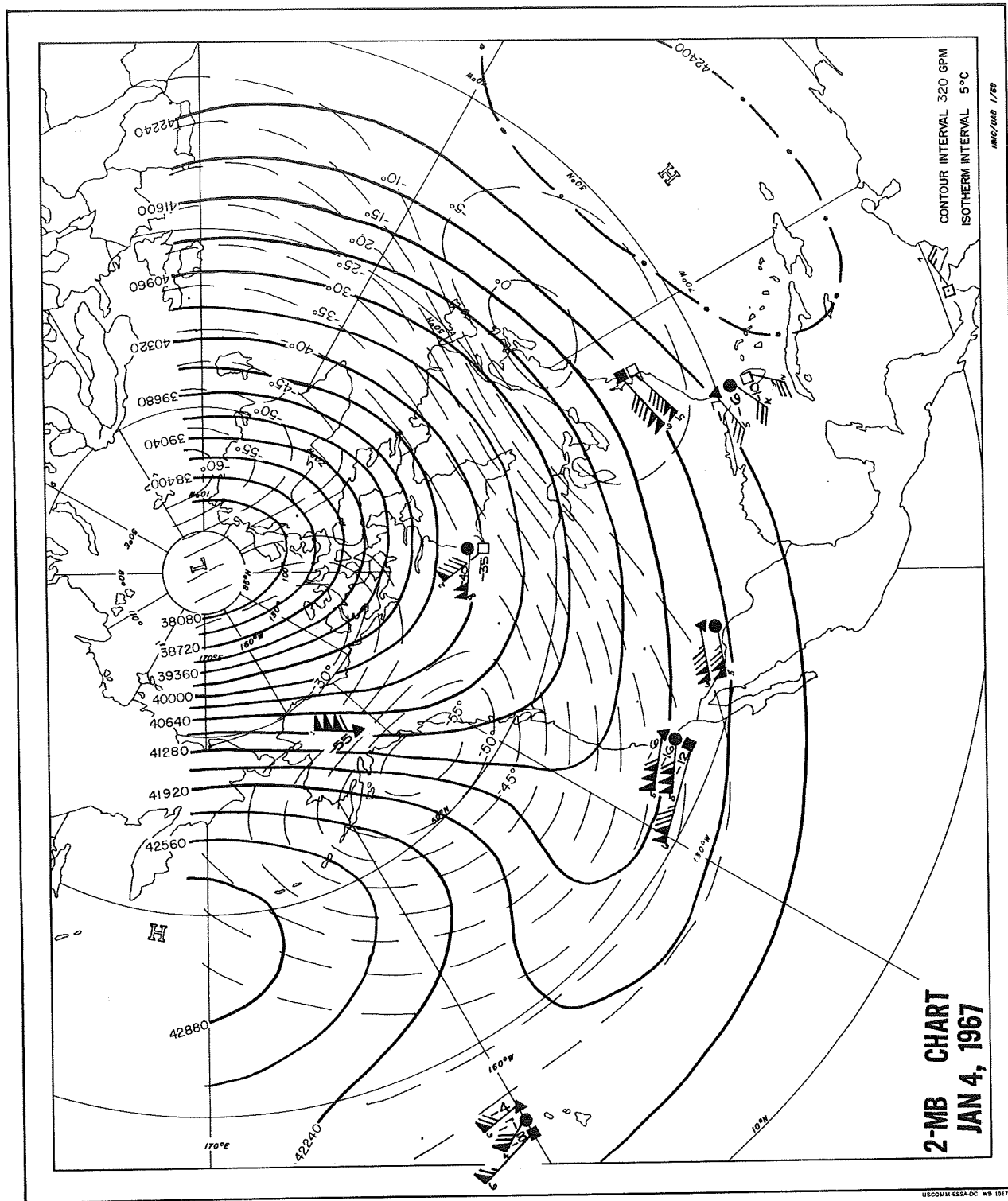
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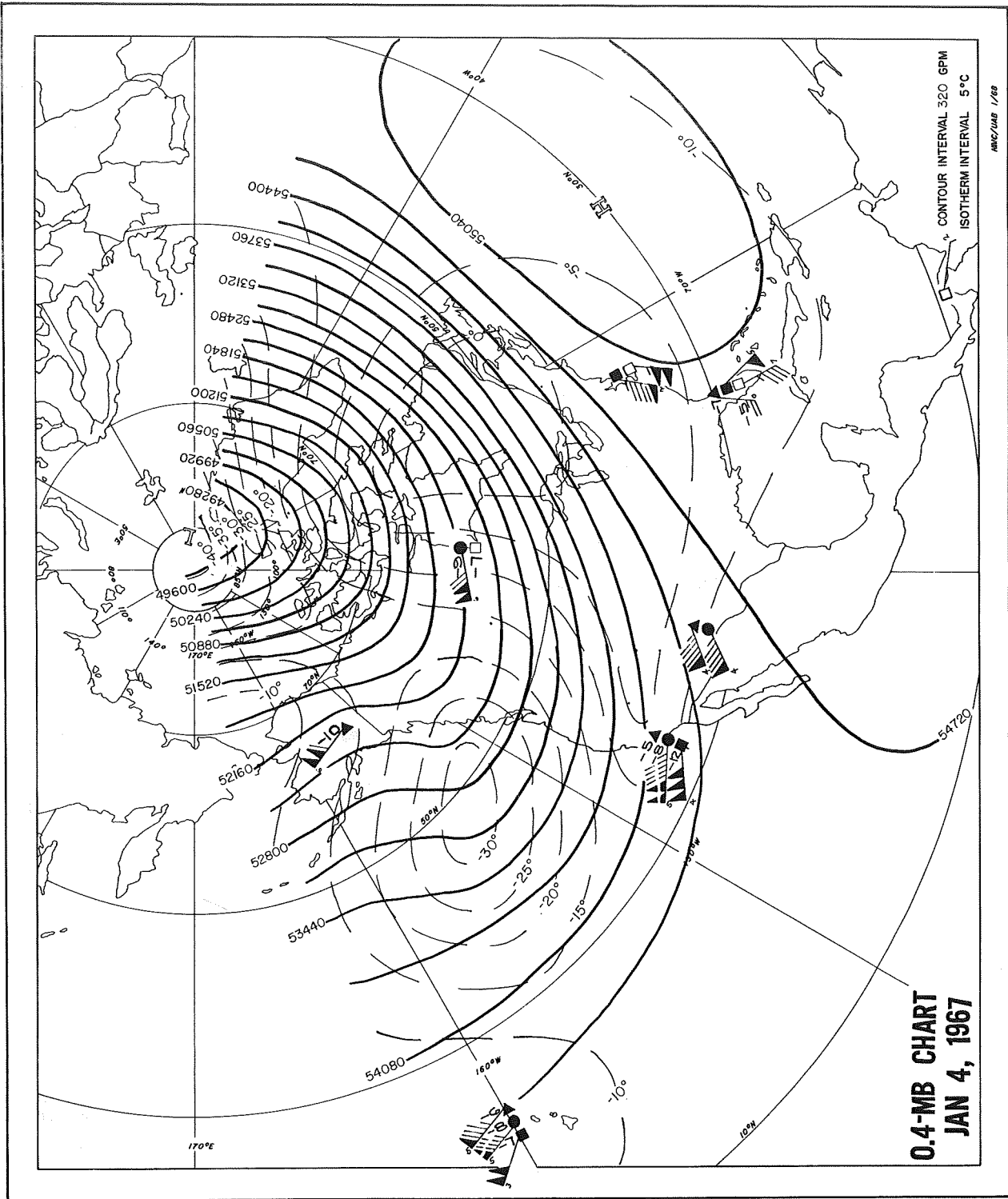
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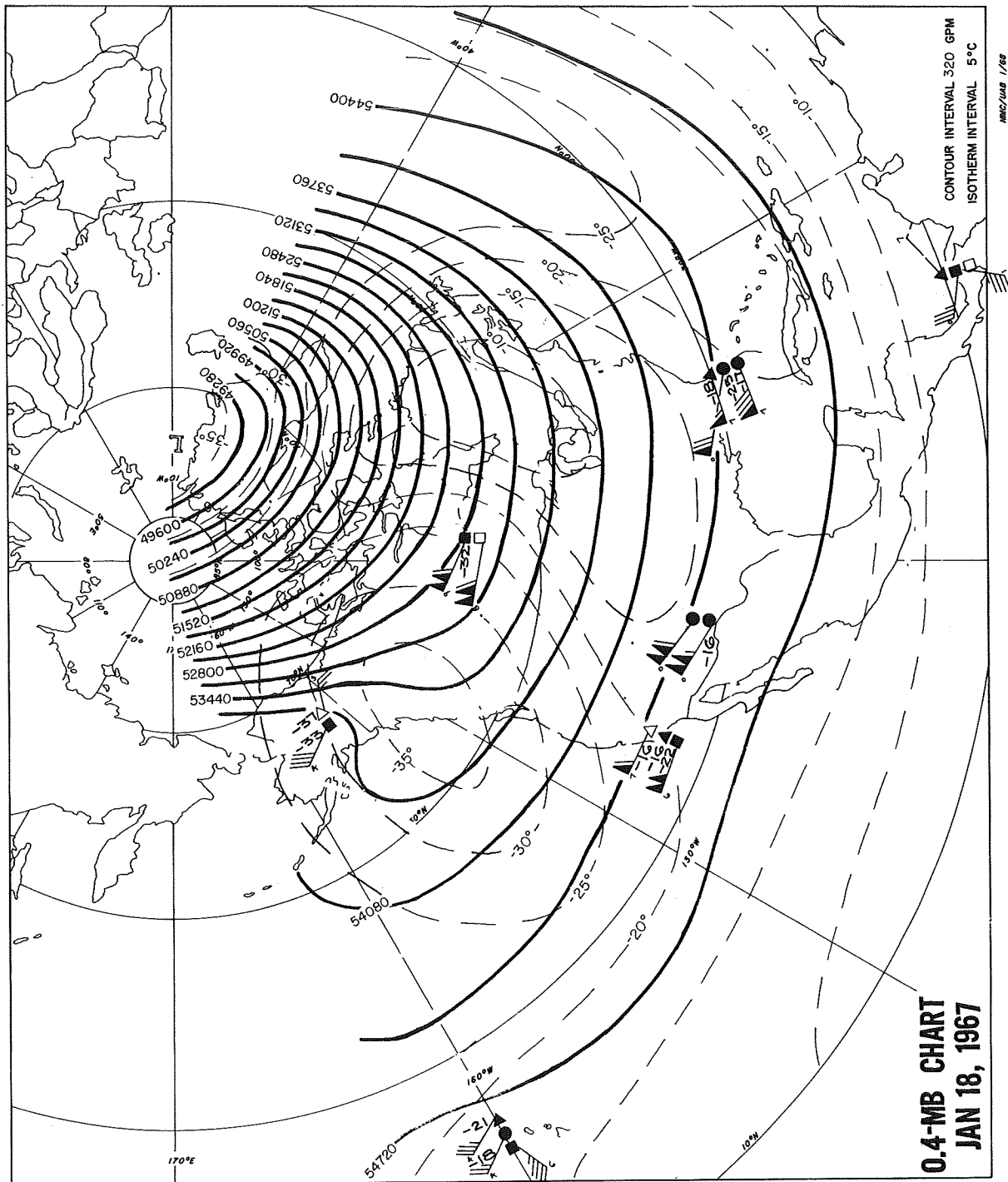
STATION MODEL AND REPORTING ROCKET STATIONS

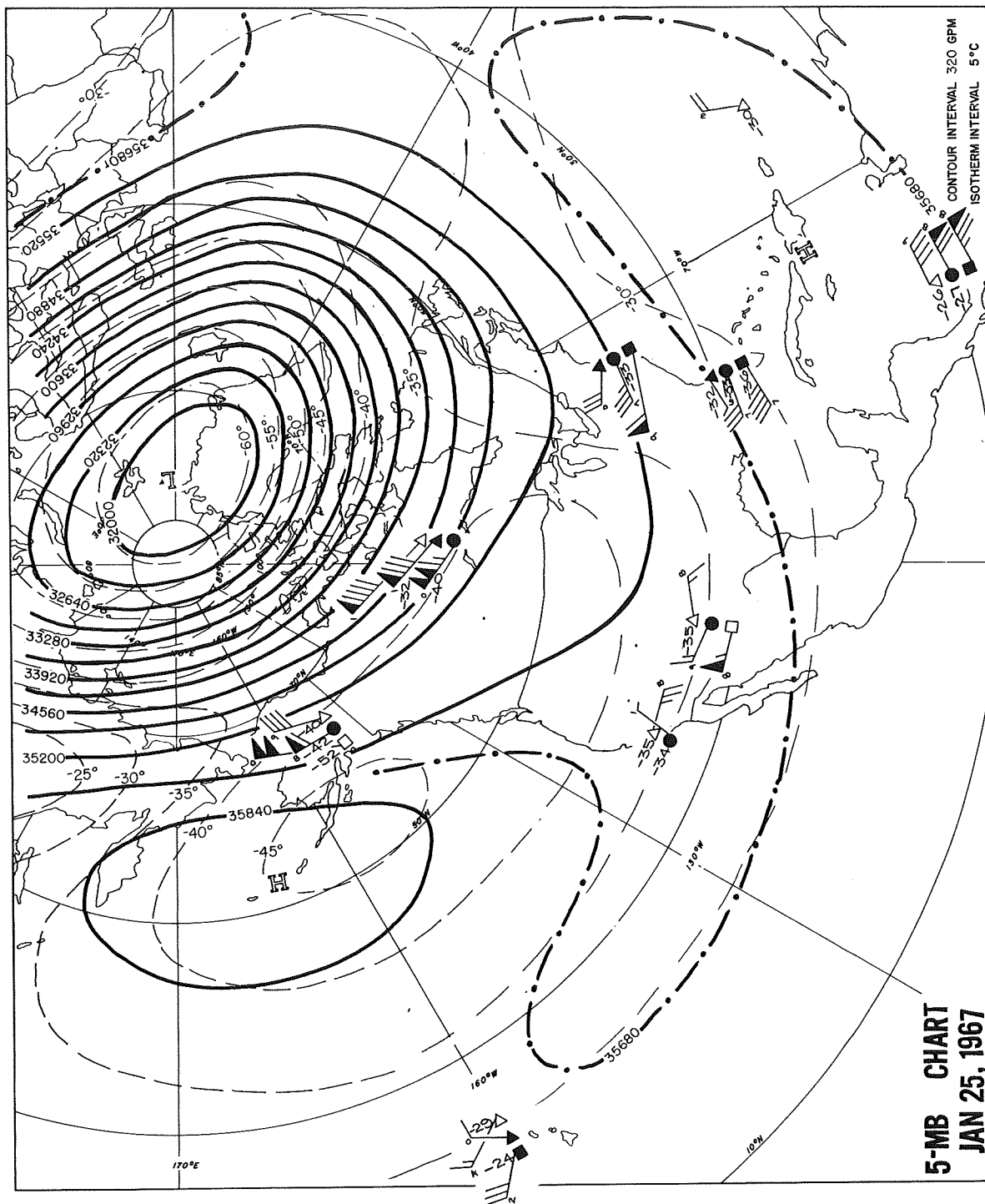




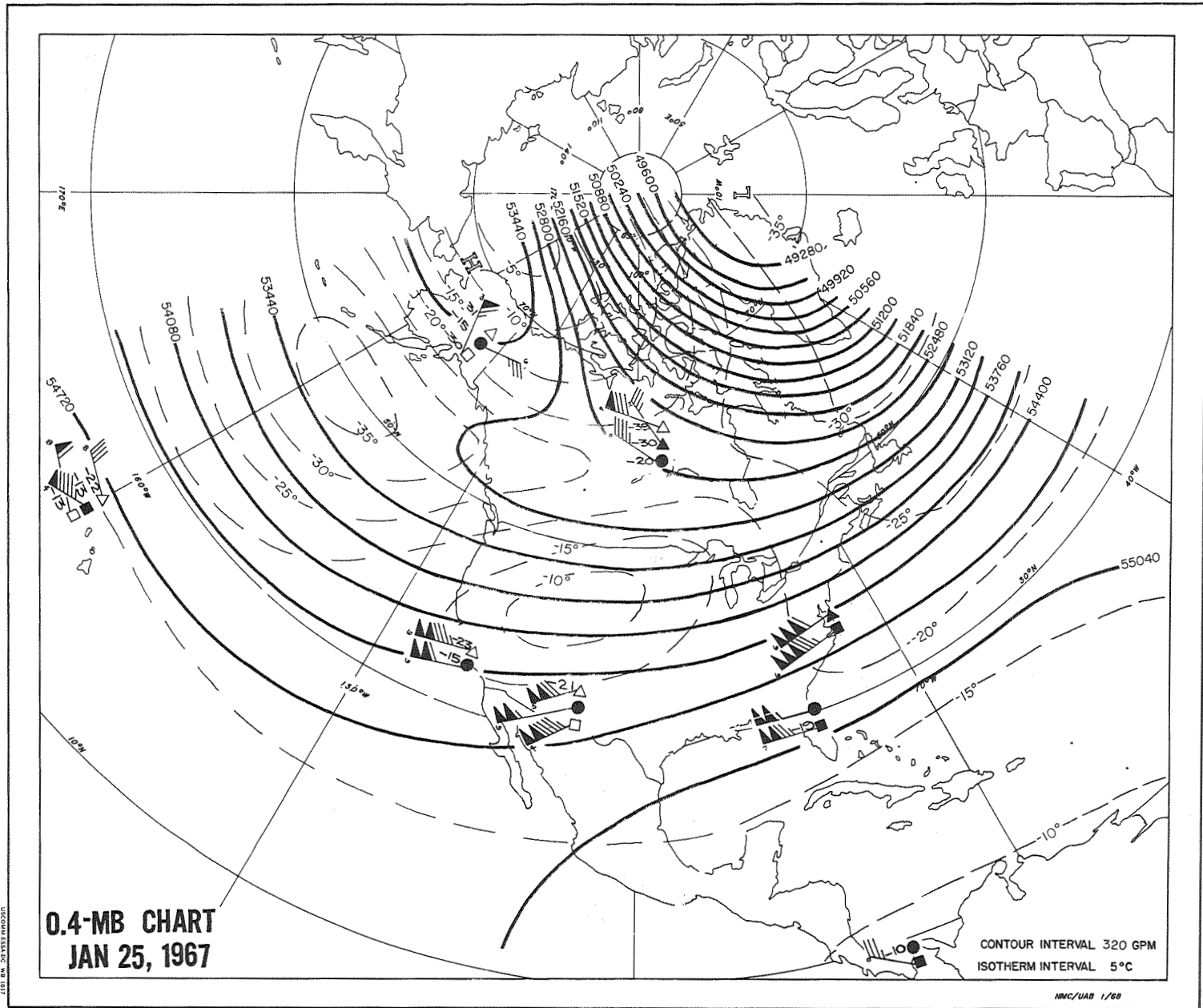


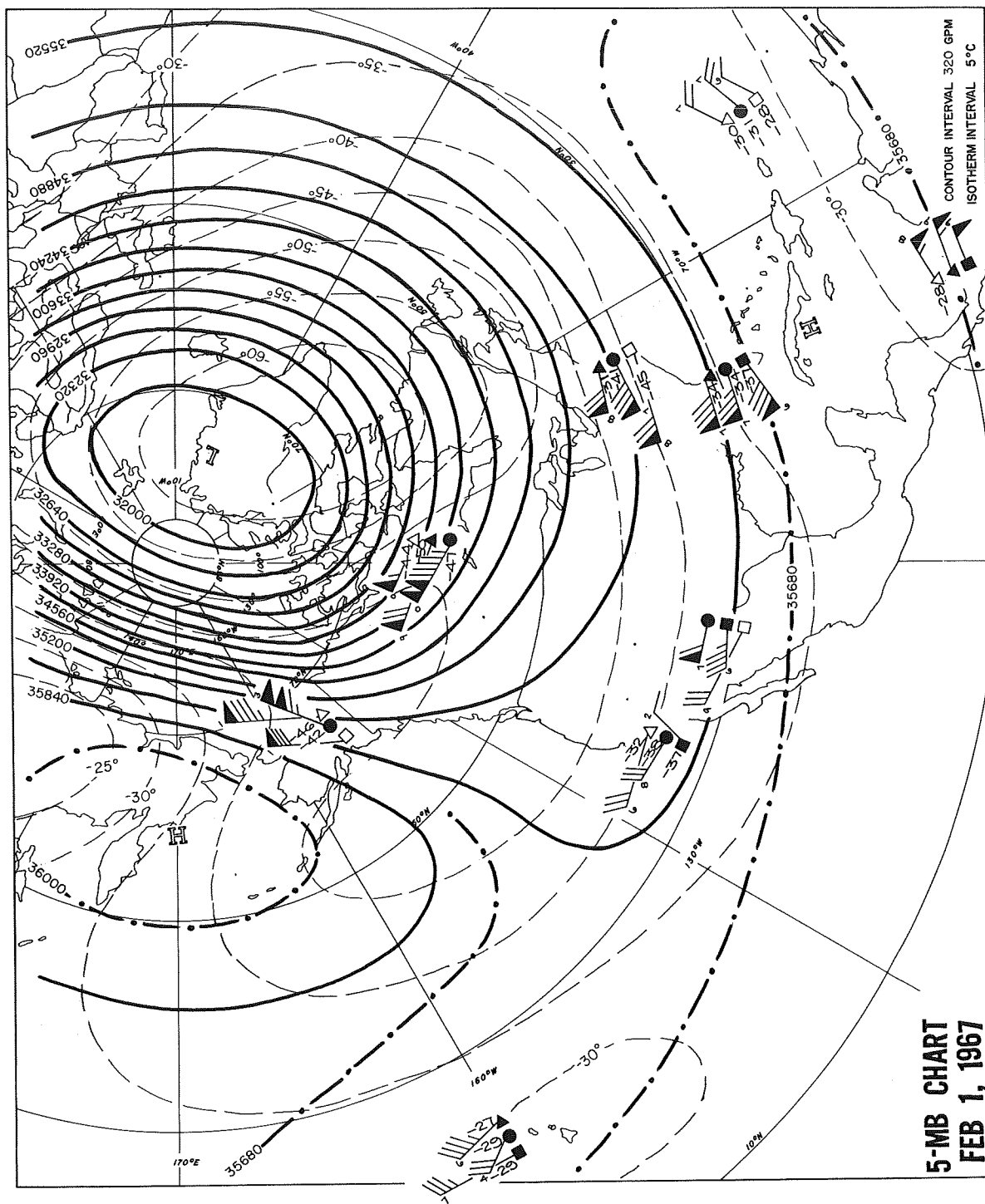




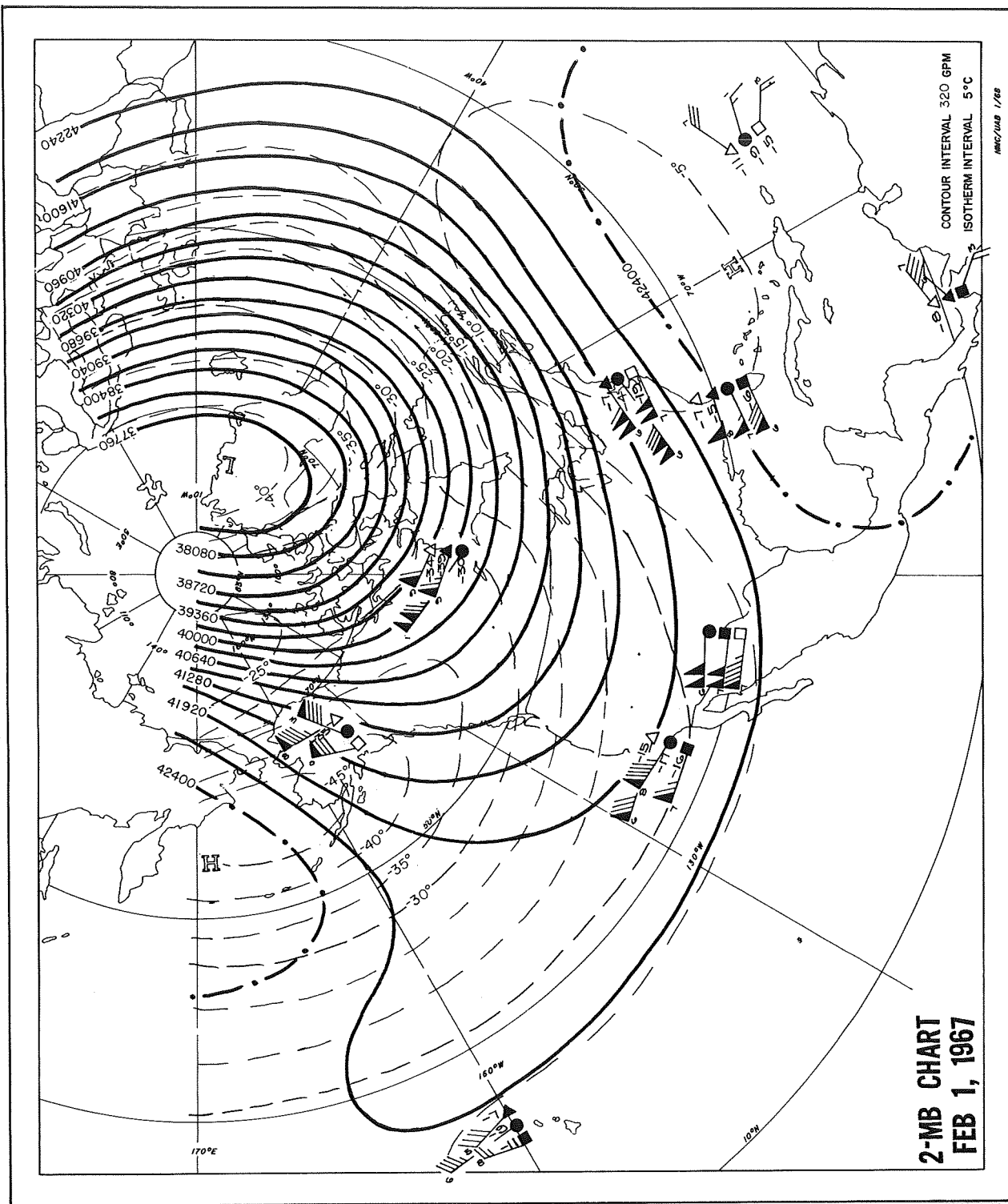


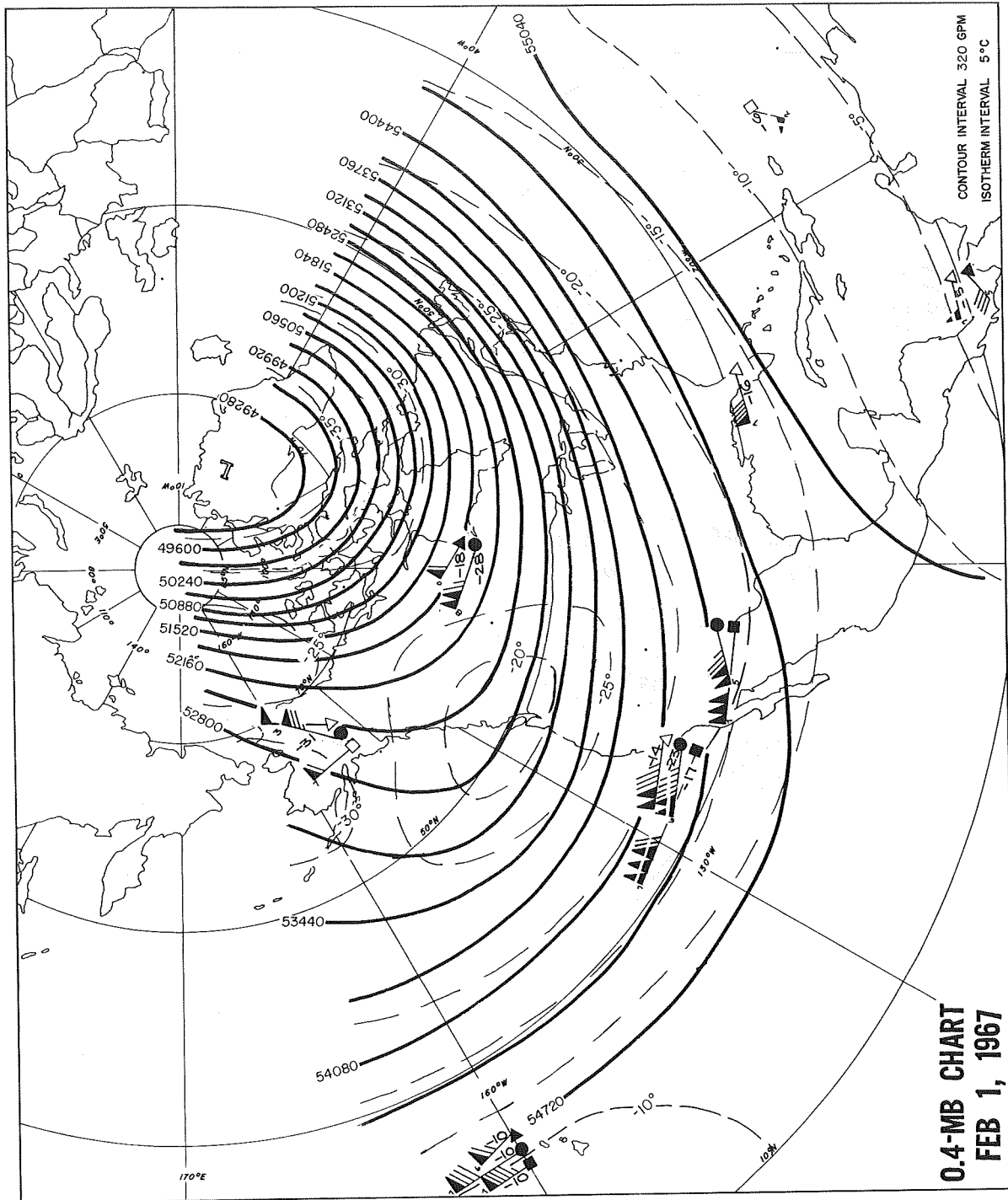
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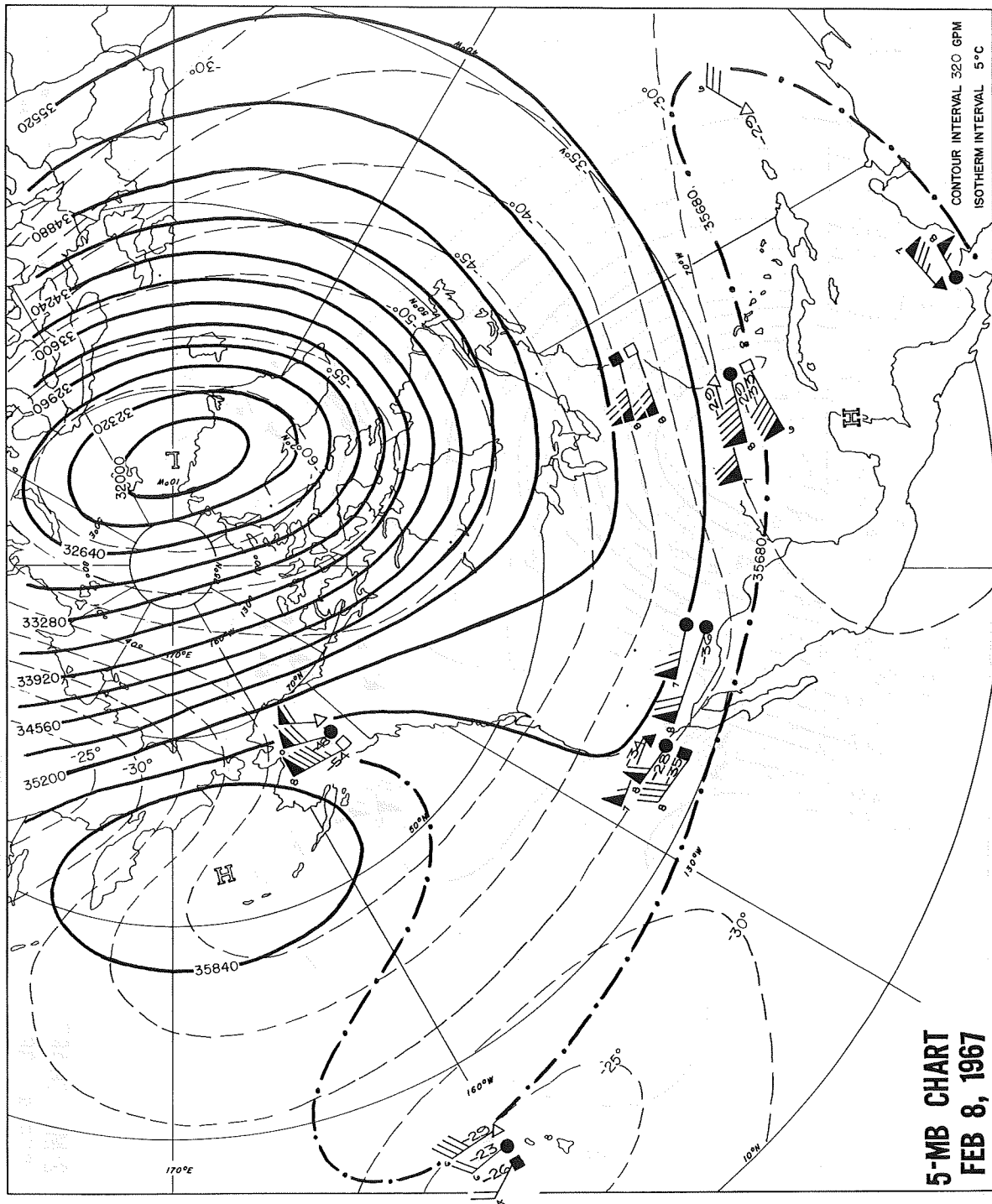


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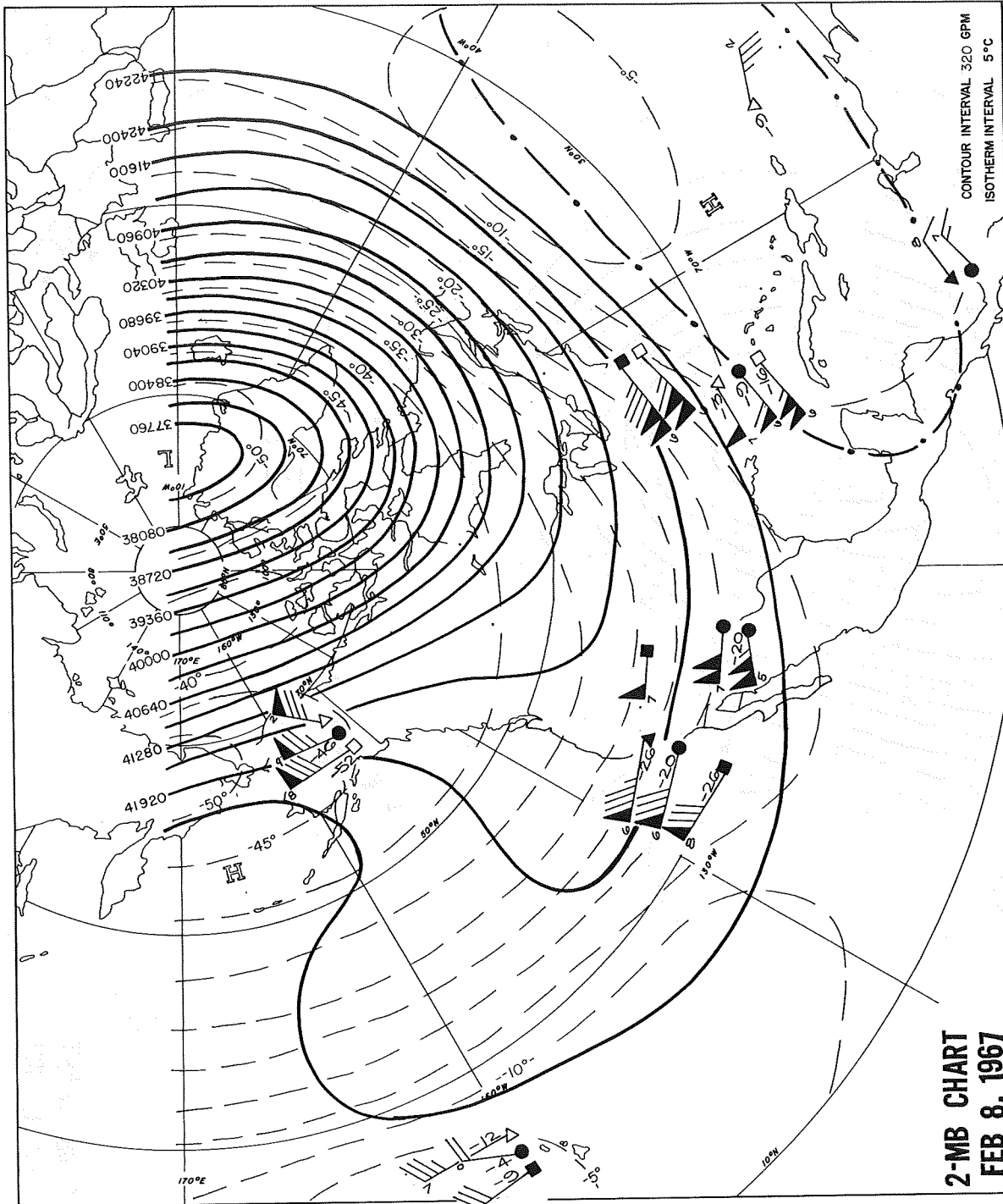




USCOMM-ESSA DC WB 1017



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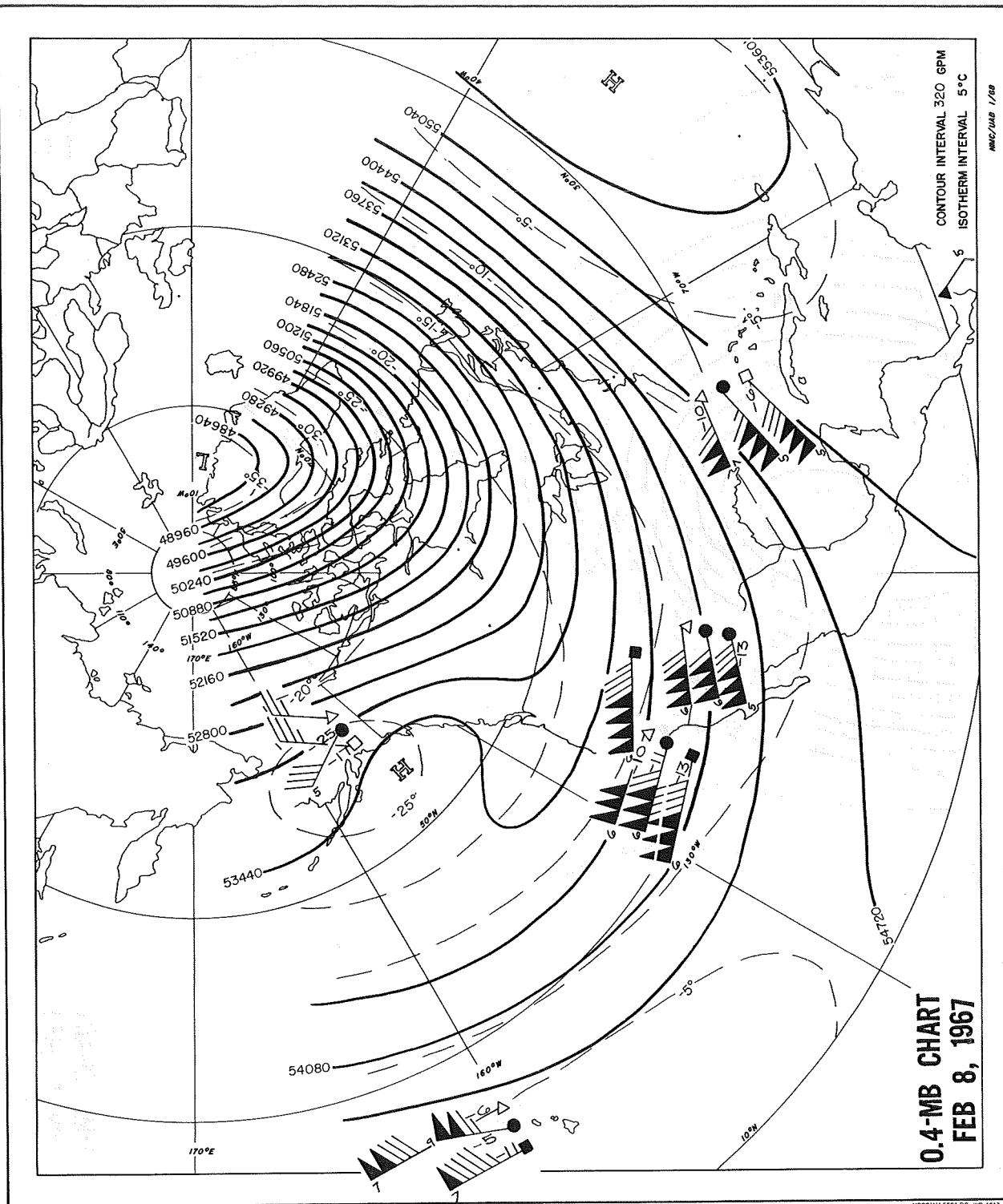


2-MB CHART
FEB 8, 1967

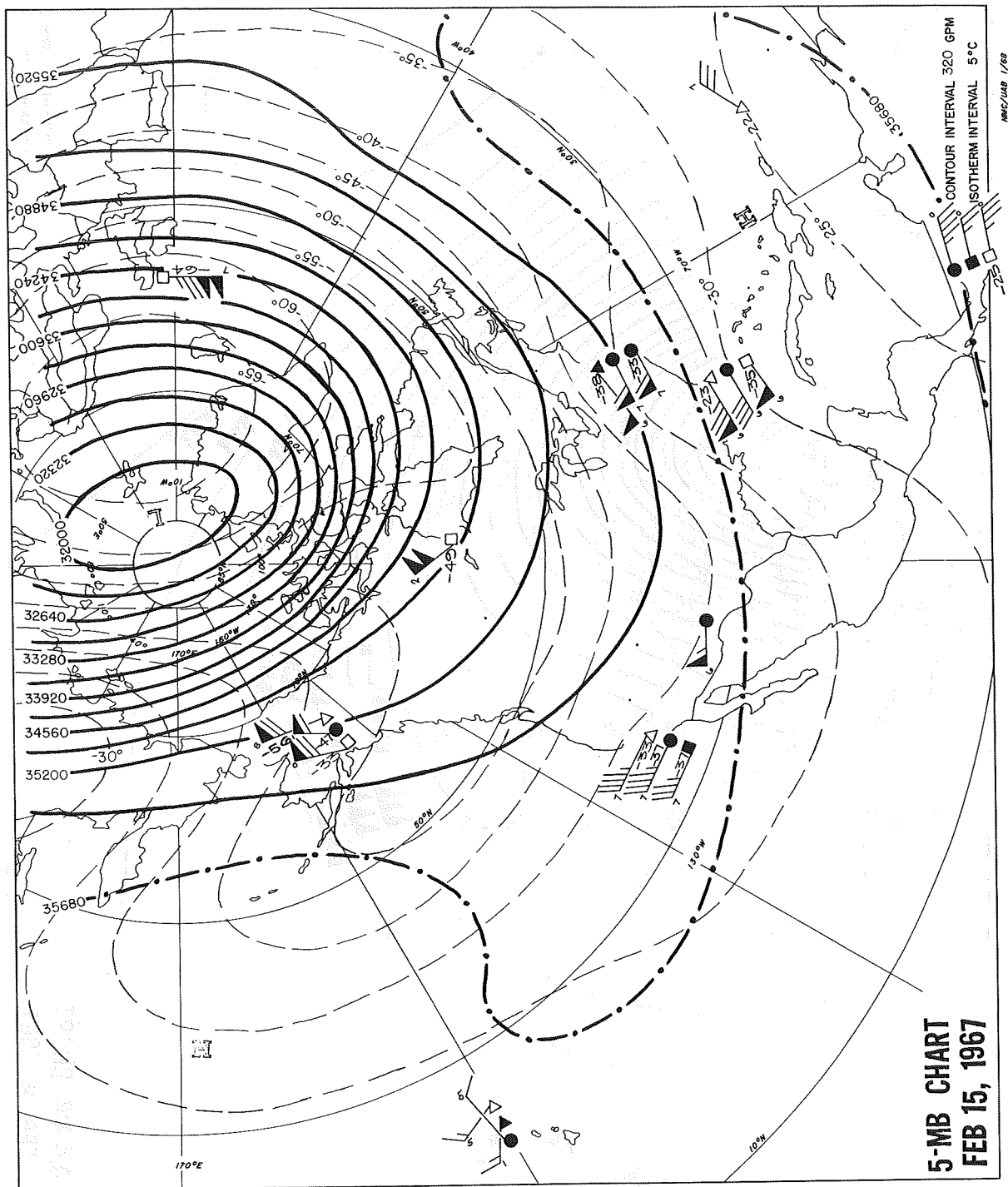
CONTOUR INTERVAL 320 GPM
ISOTHERM INTERVAL 5°C

NRAC/USAF 1/68

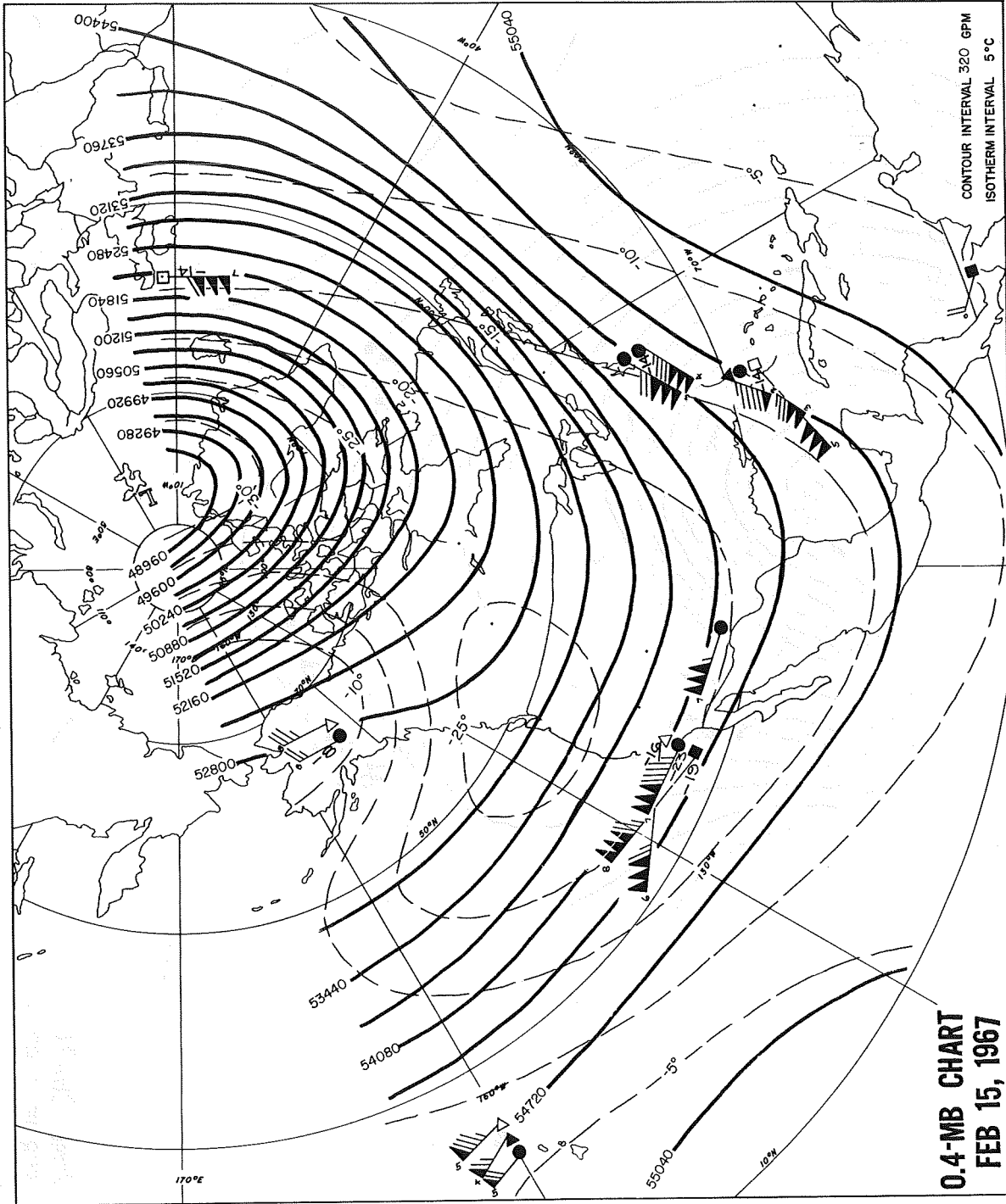
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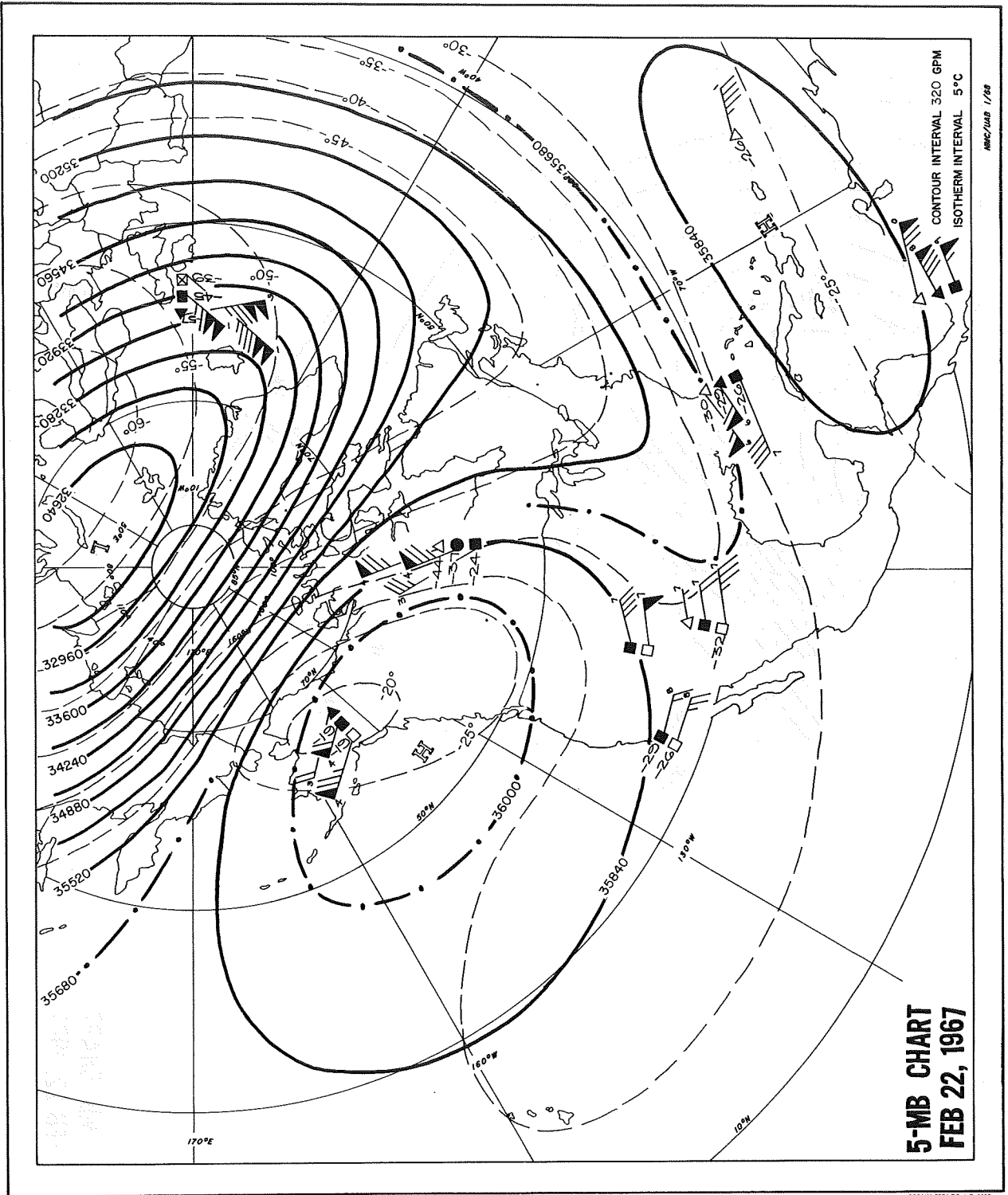
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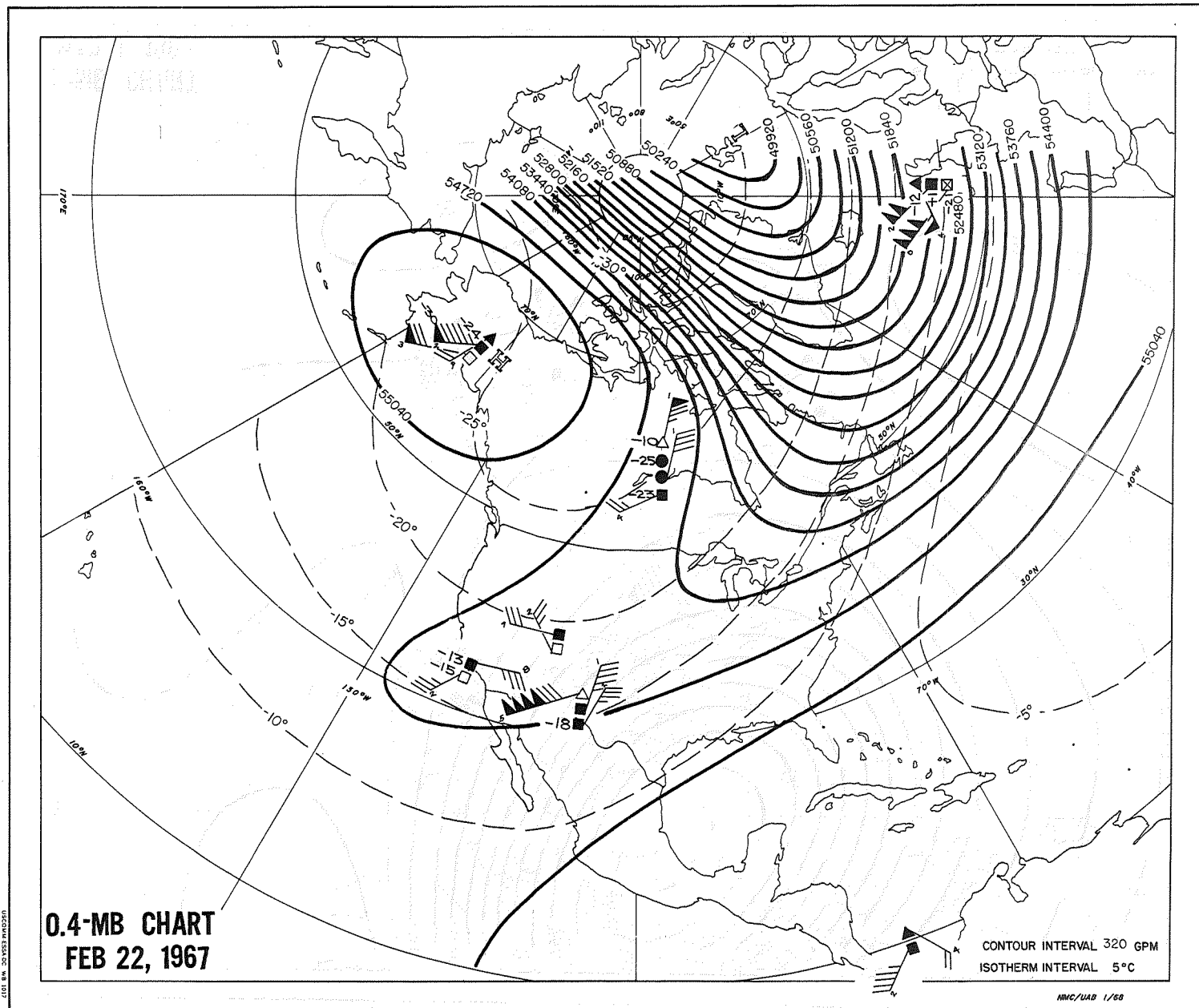


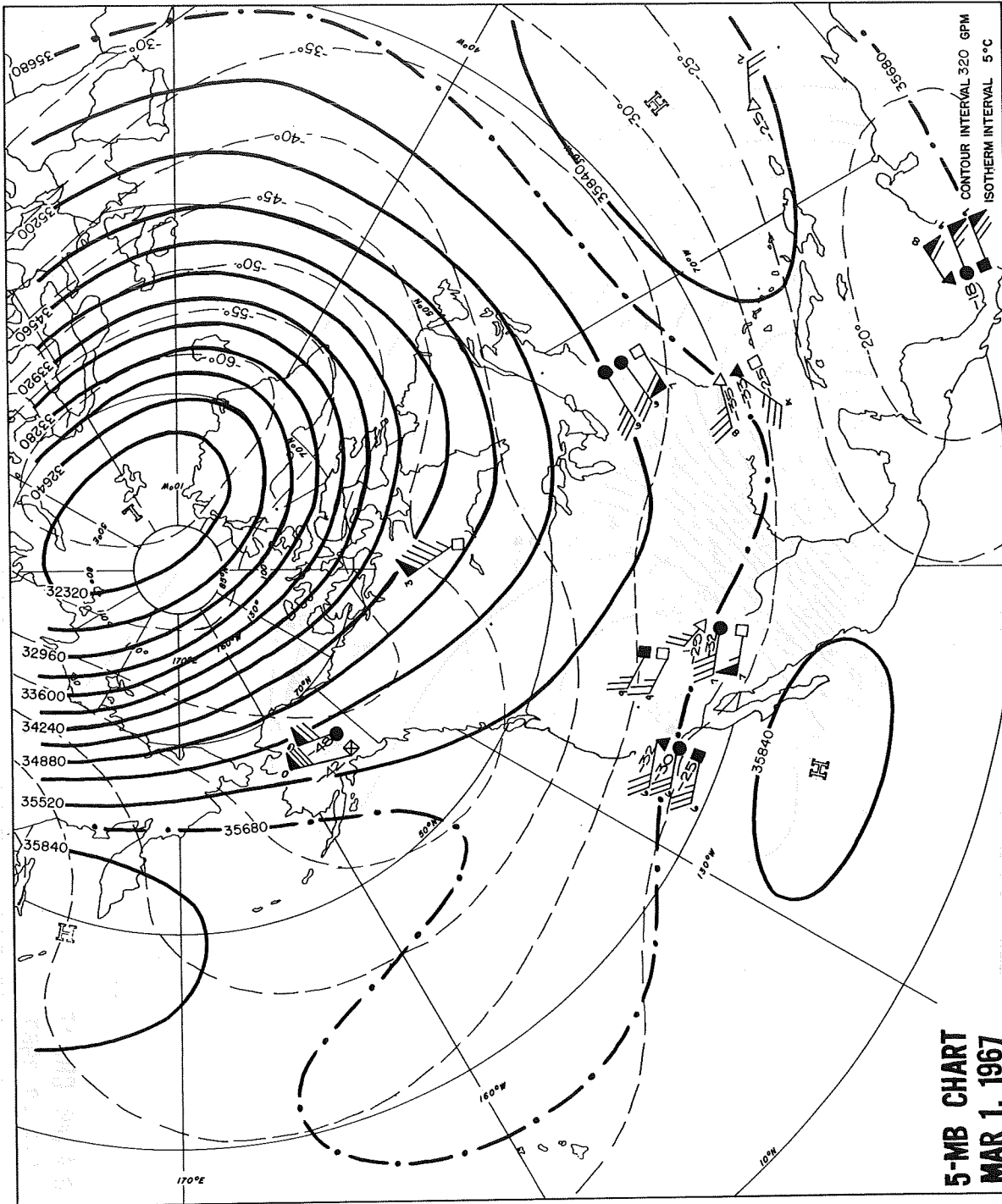
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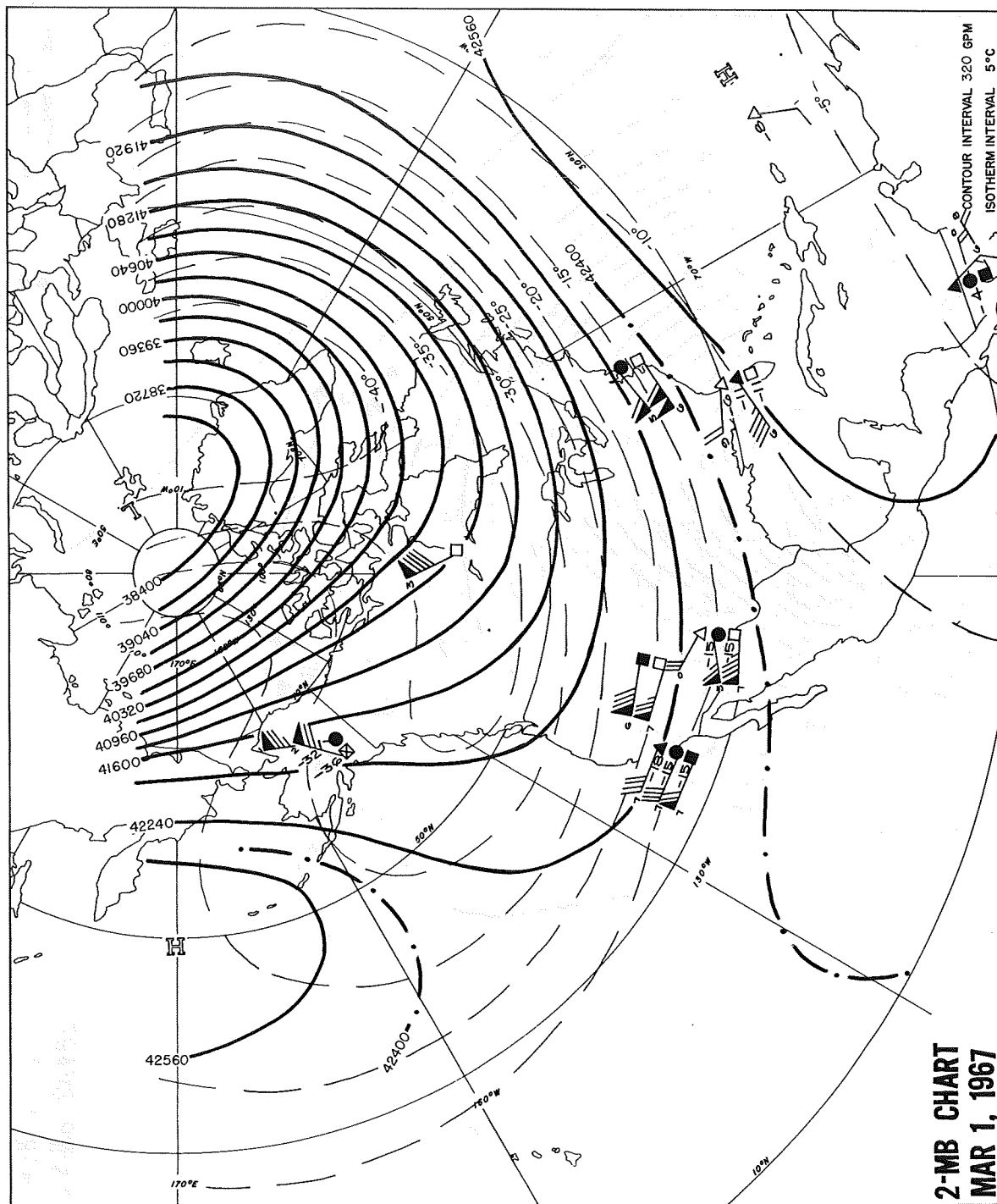


USCNAV ESSA DC WB 1017





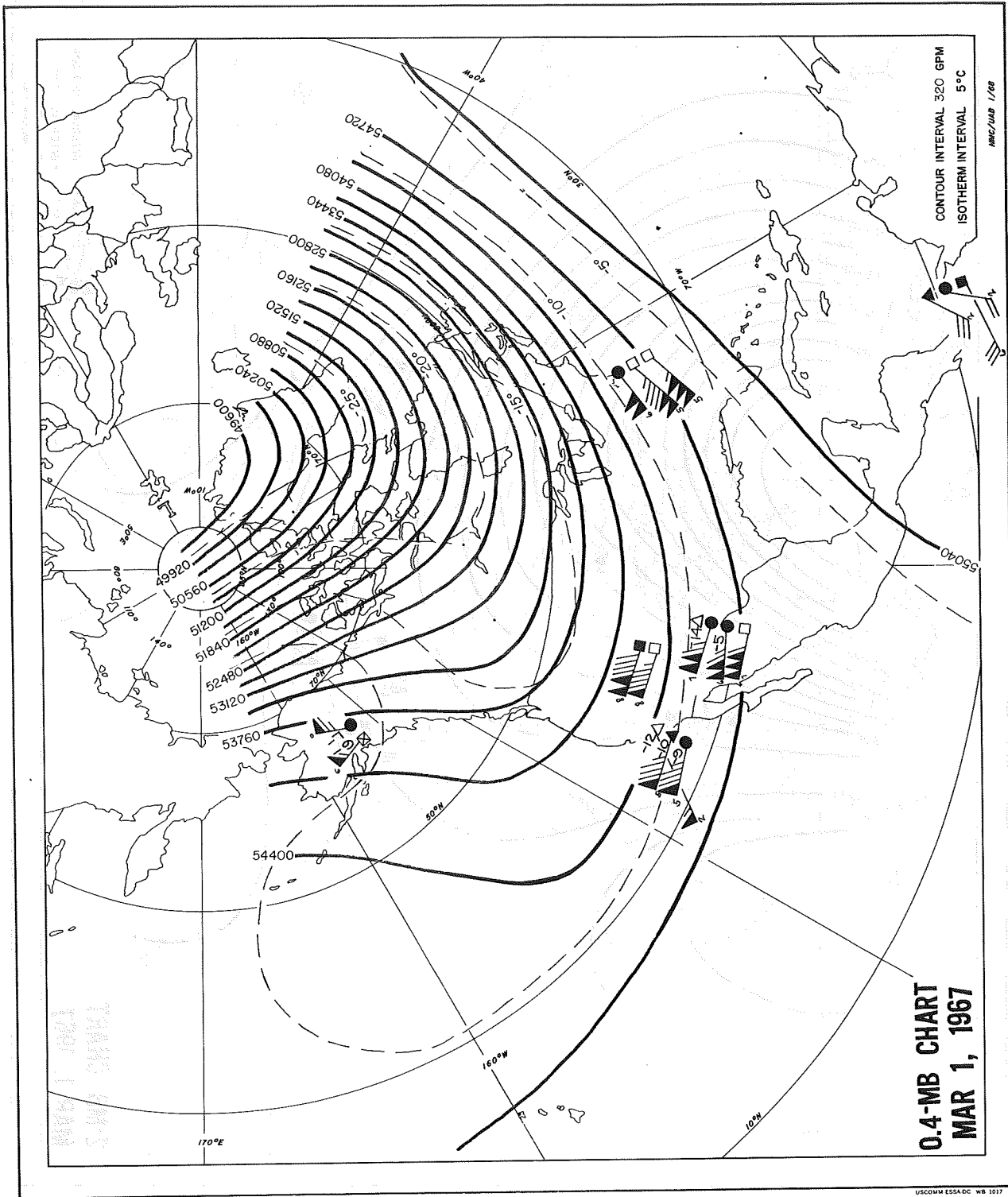




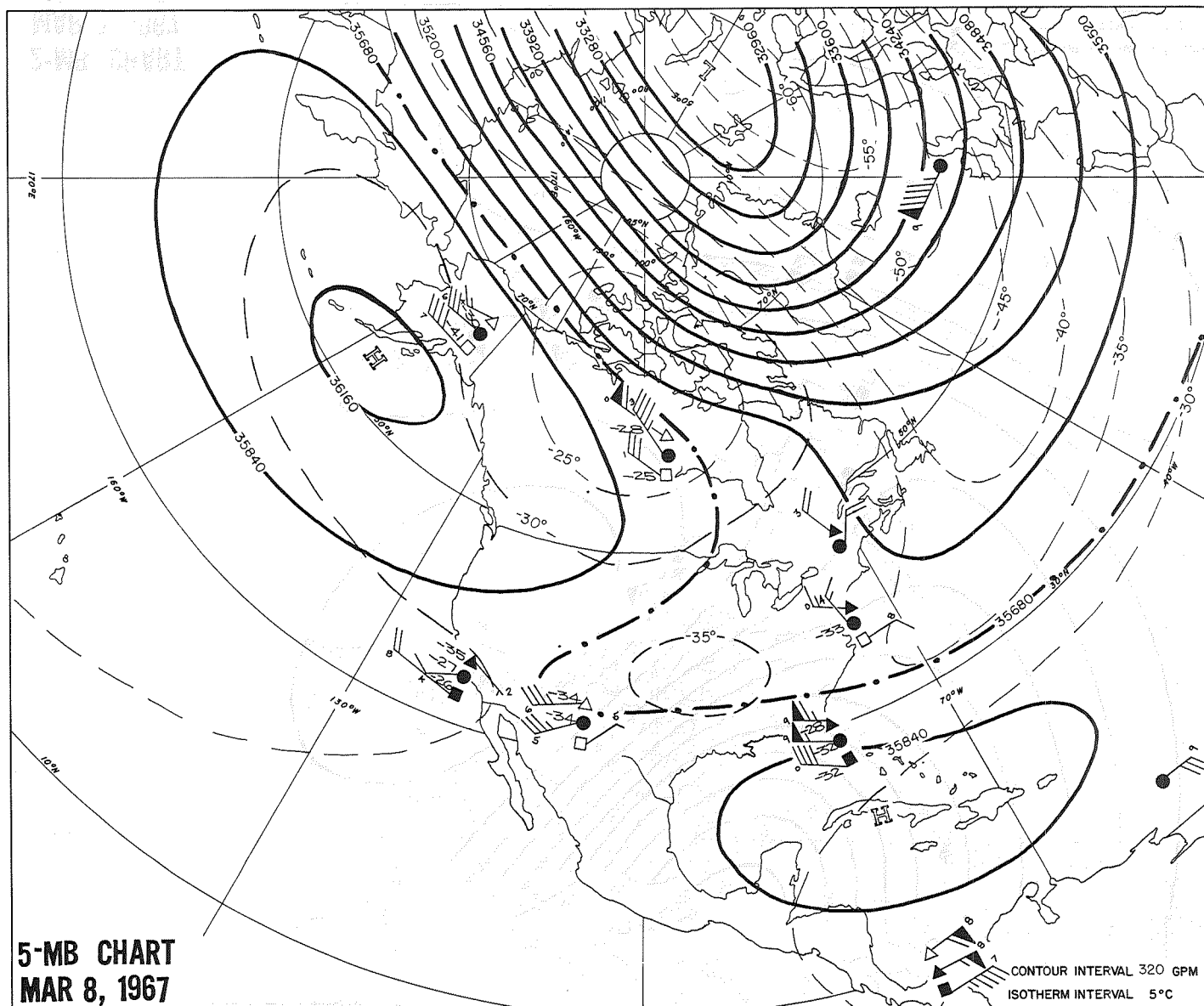
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MAR 1, 1967**

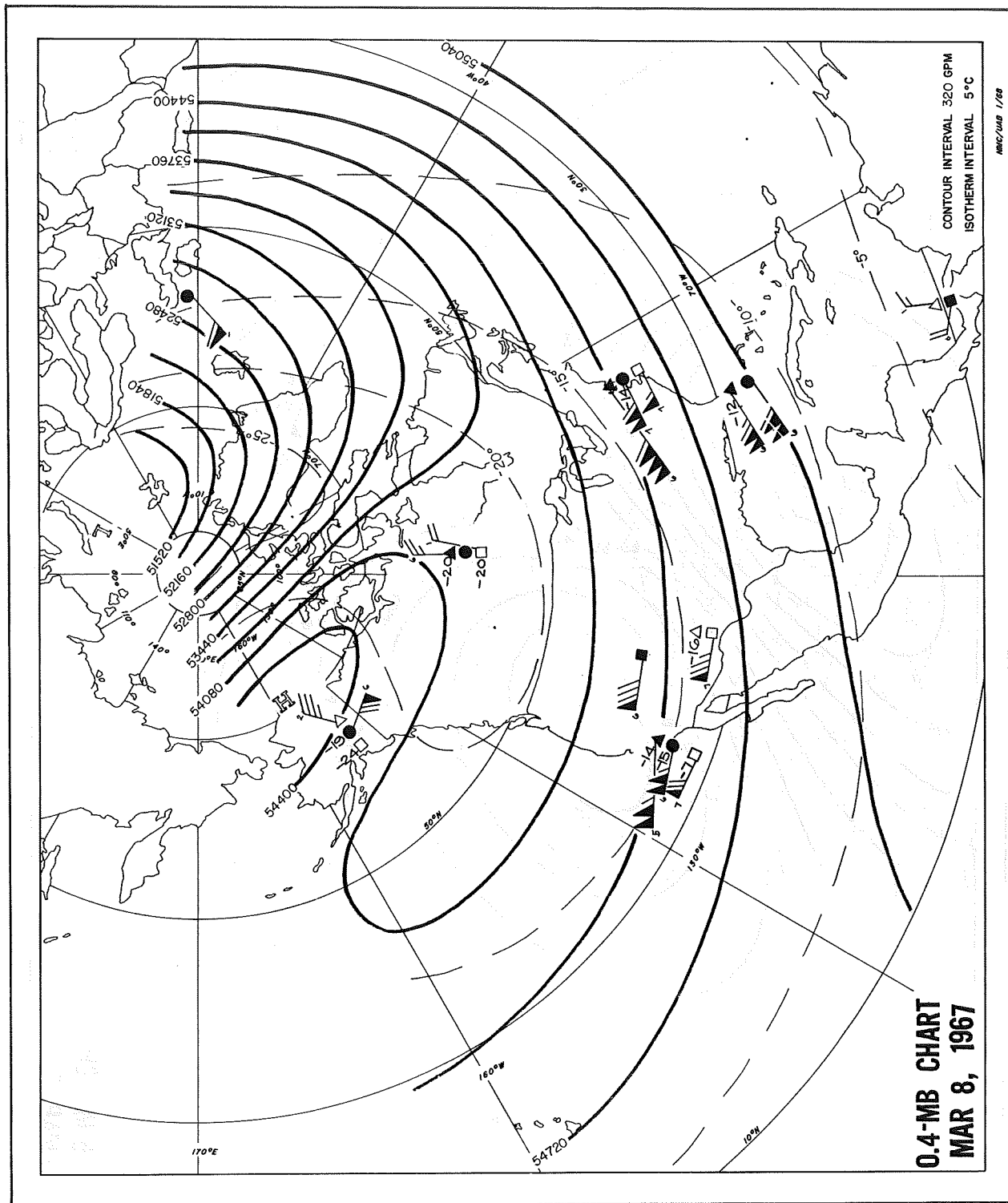
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NAVJAG 1/68

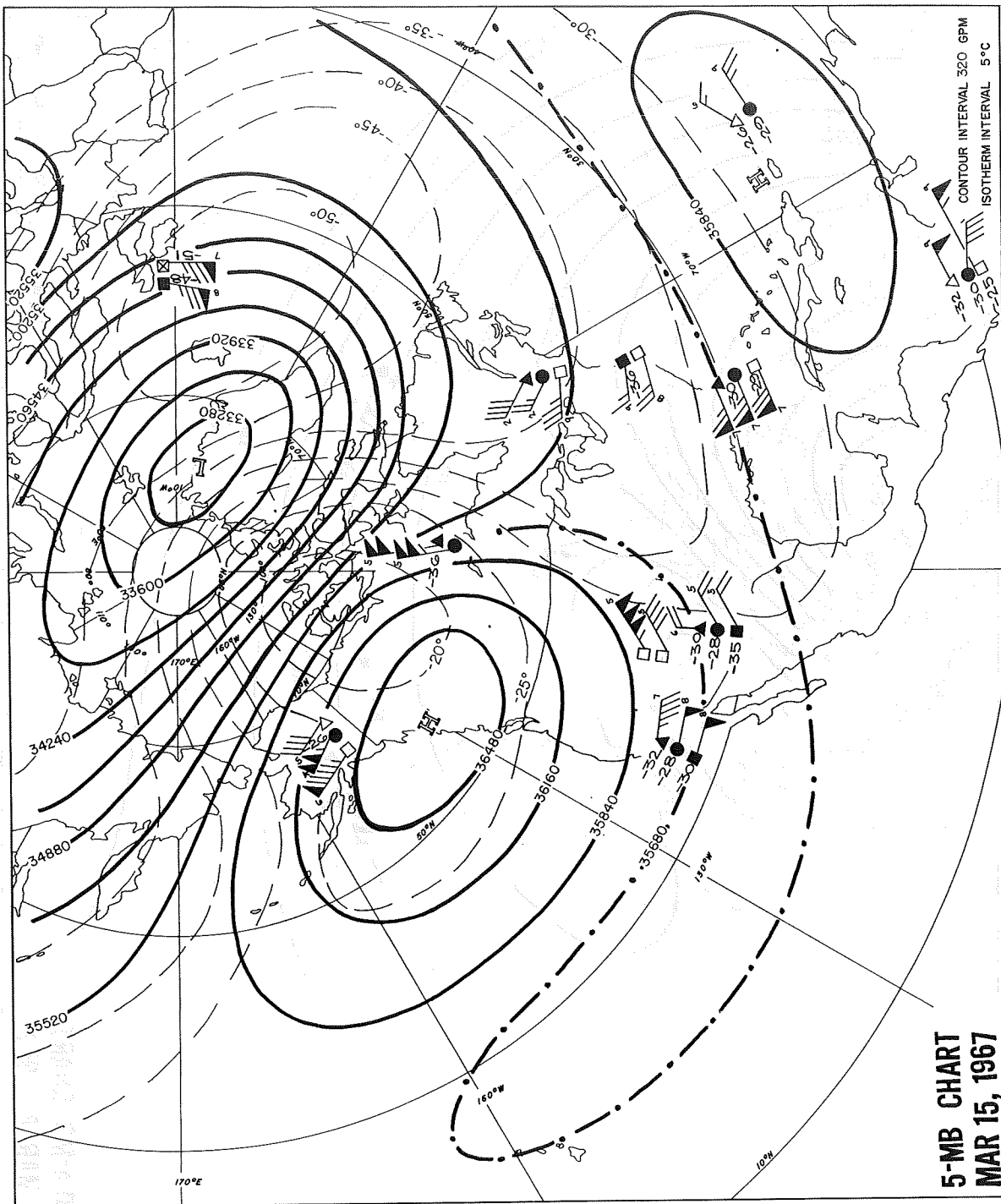


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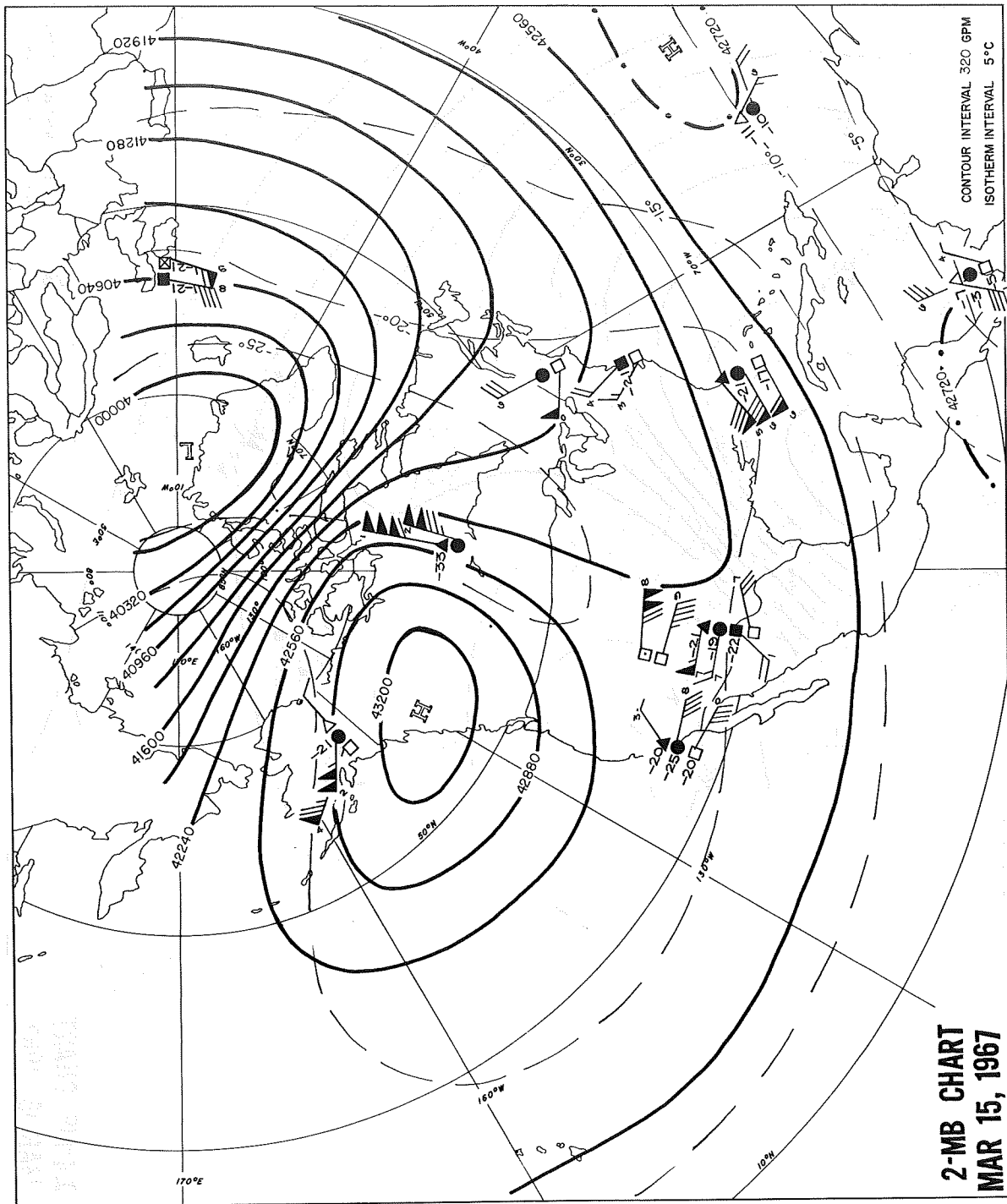




USCOWH ESSA-DC WB 1017



USCOMM ESSADC WB 1017

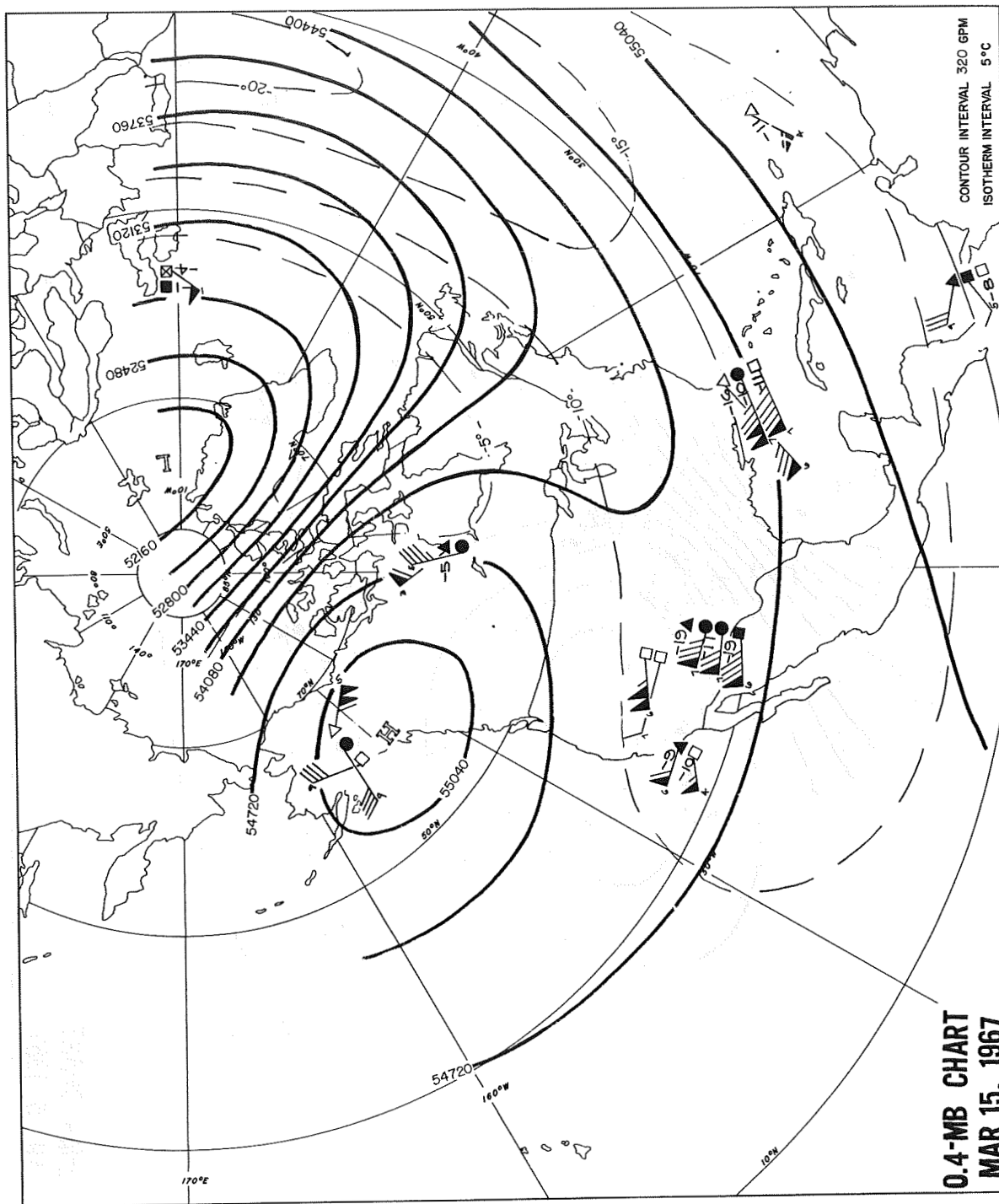


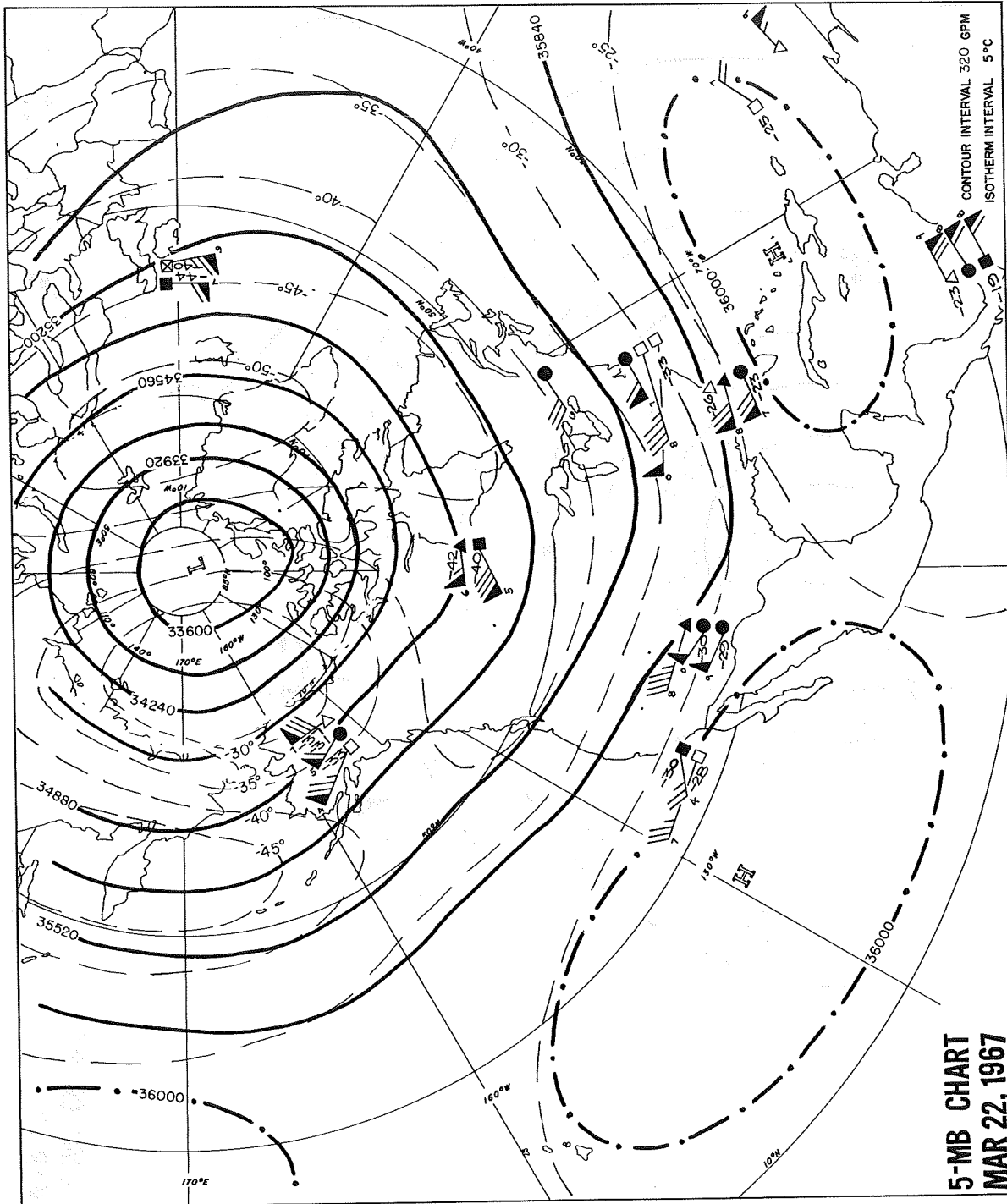
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CONTOUR INTERVAL 320 GPM
ISOTHERM INTERVAL 5°C

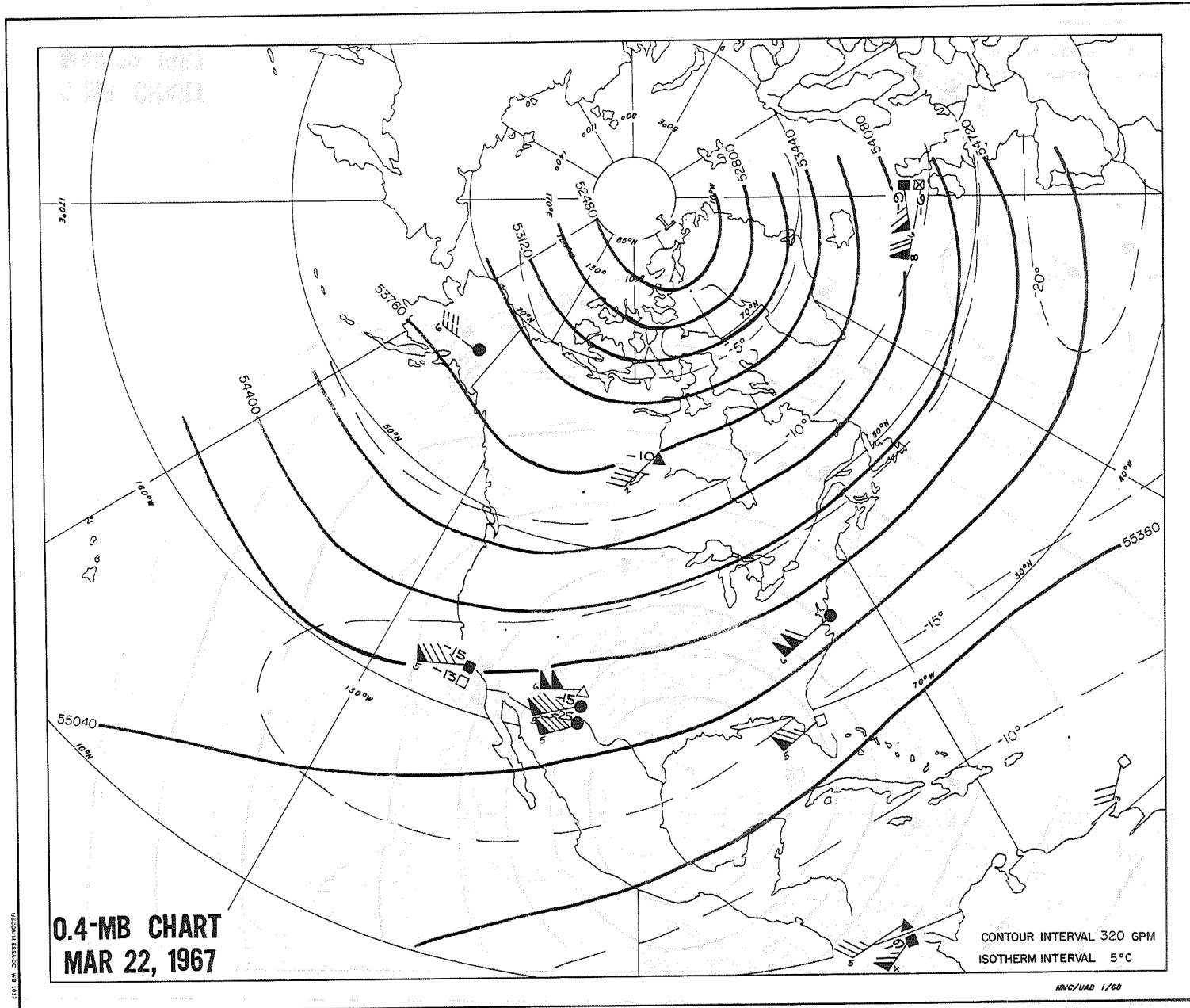
WMO/USAF 1/68

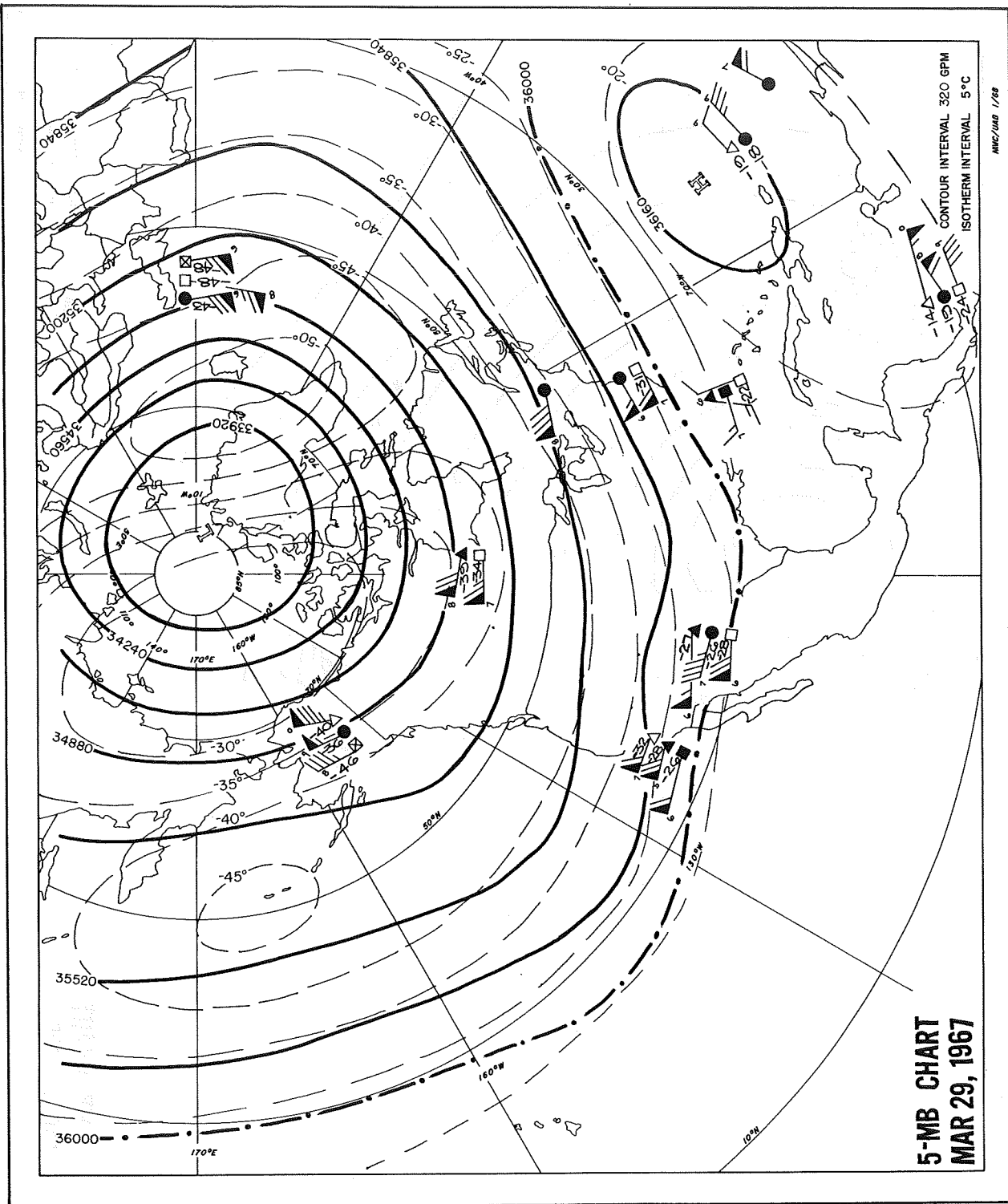
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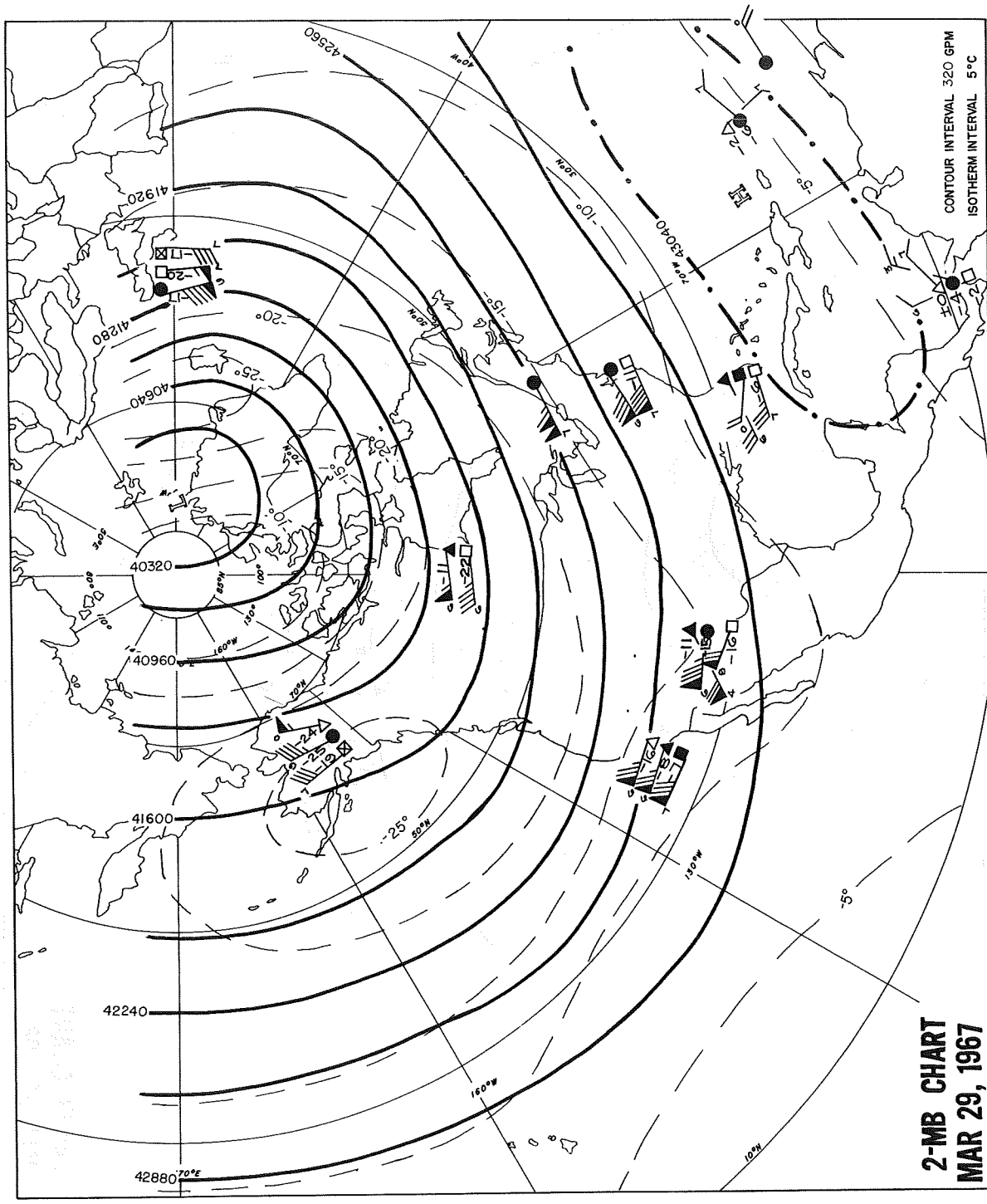




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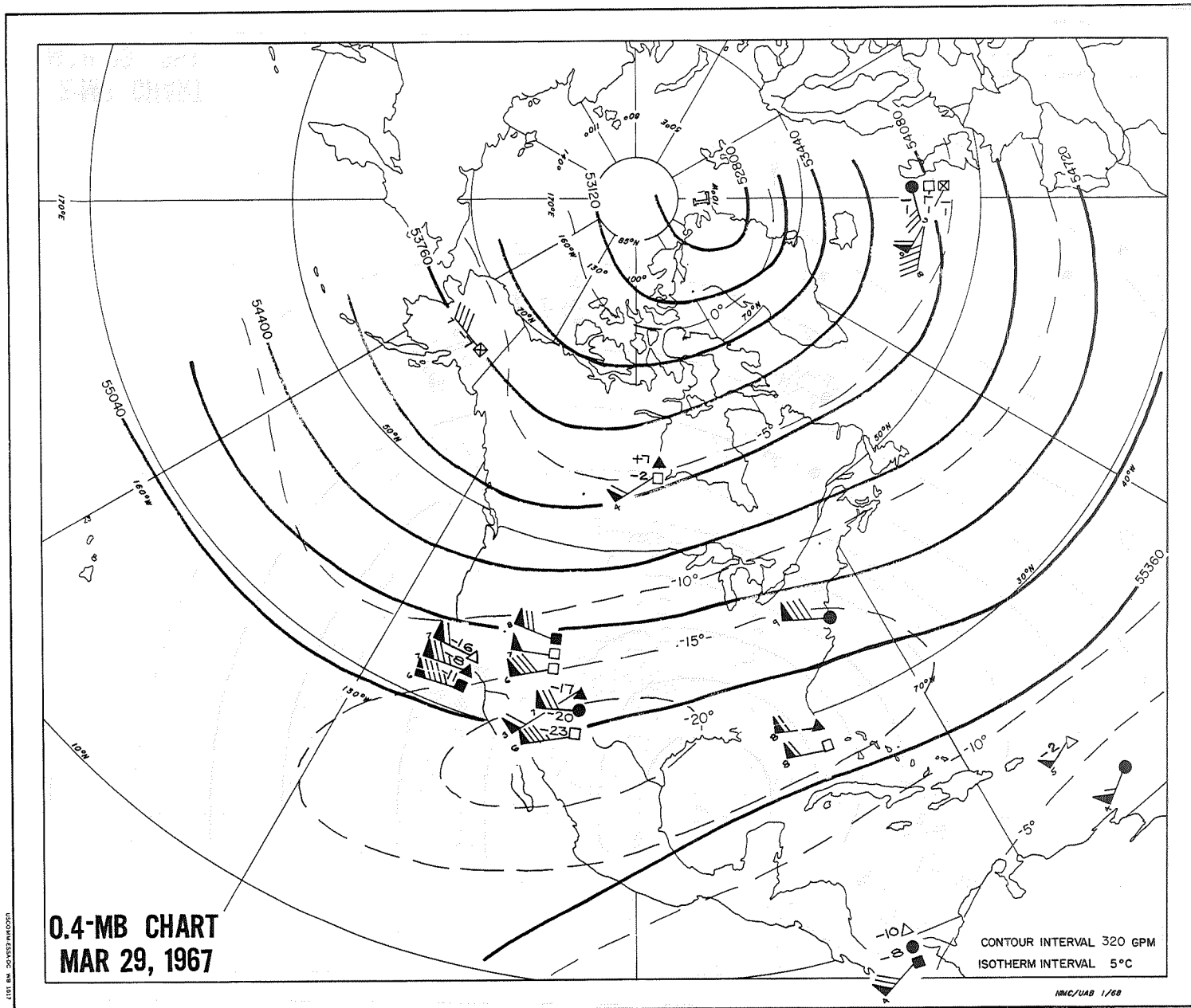


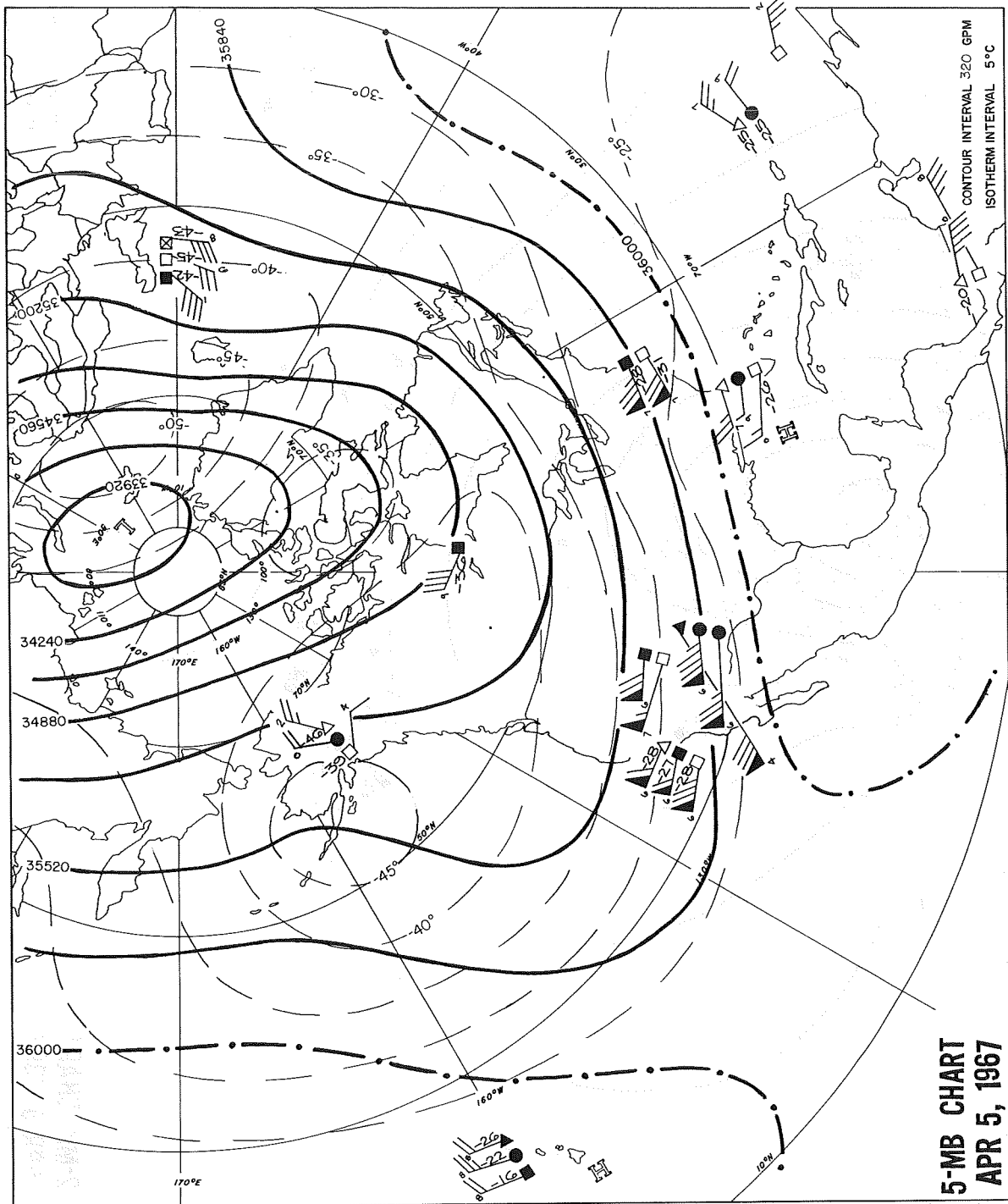




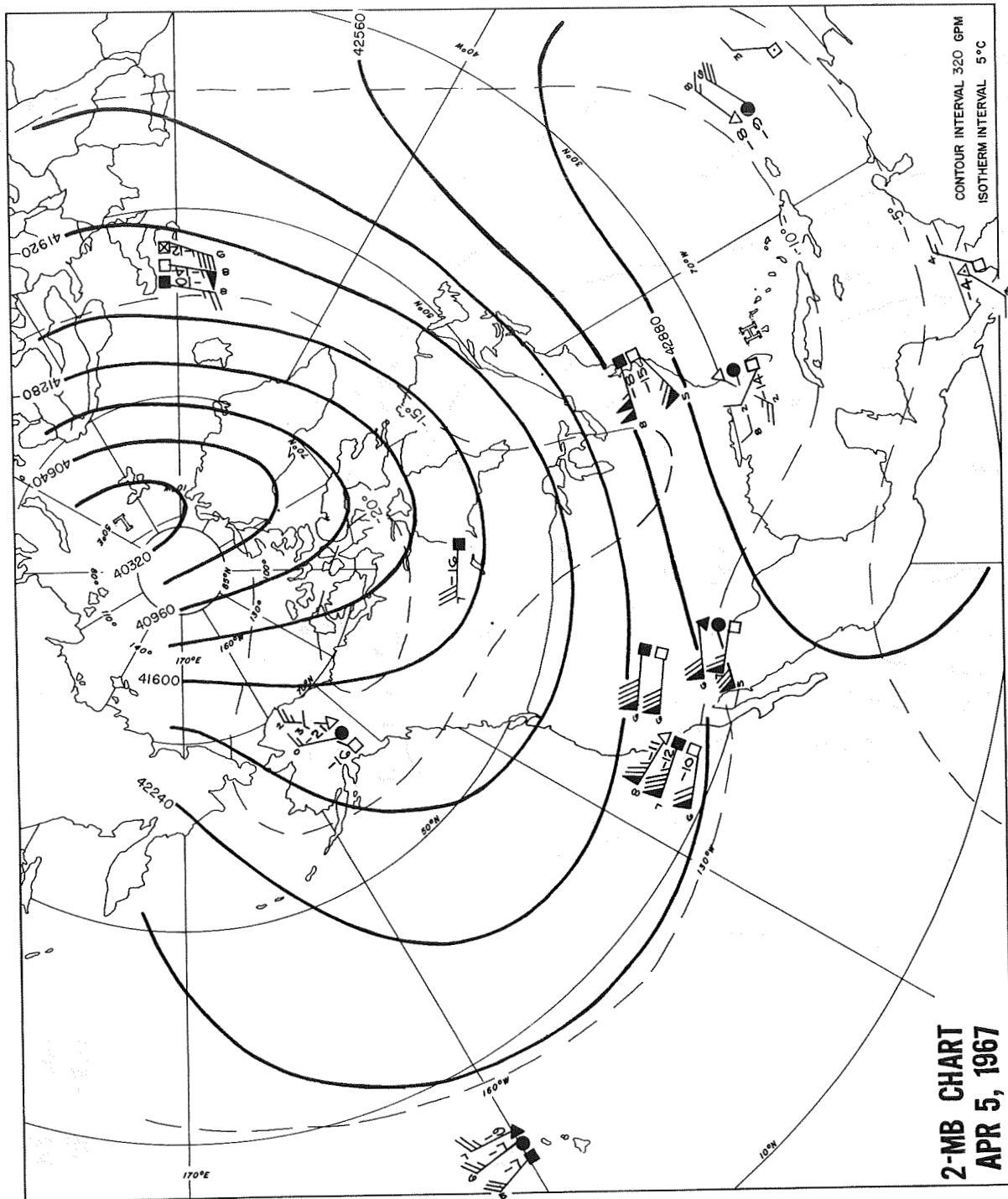
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NMC/US 1/68

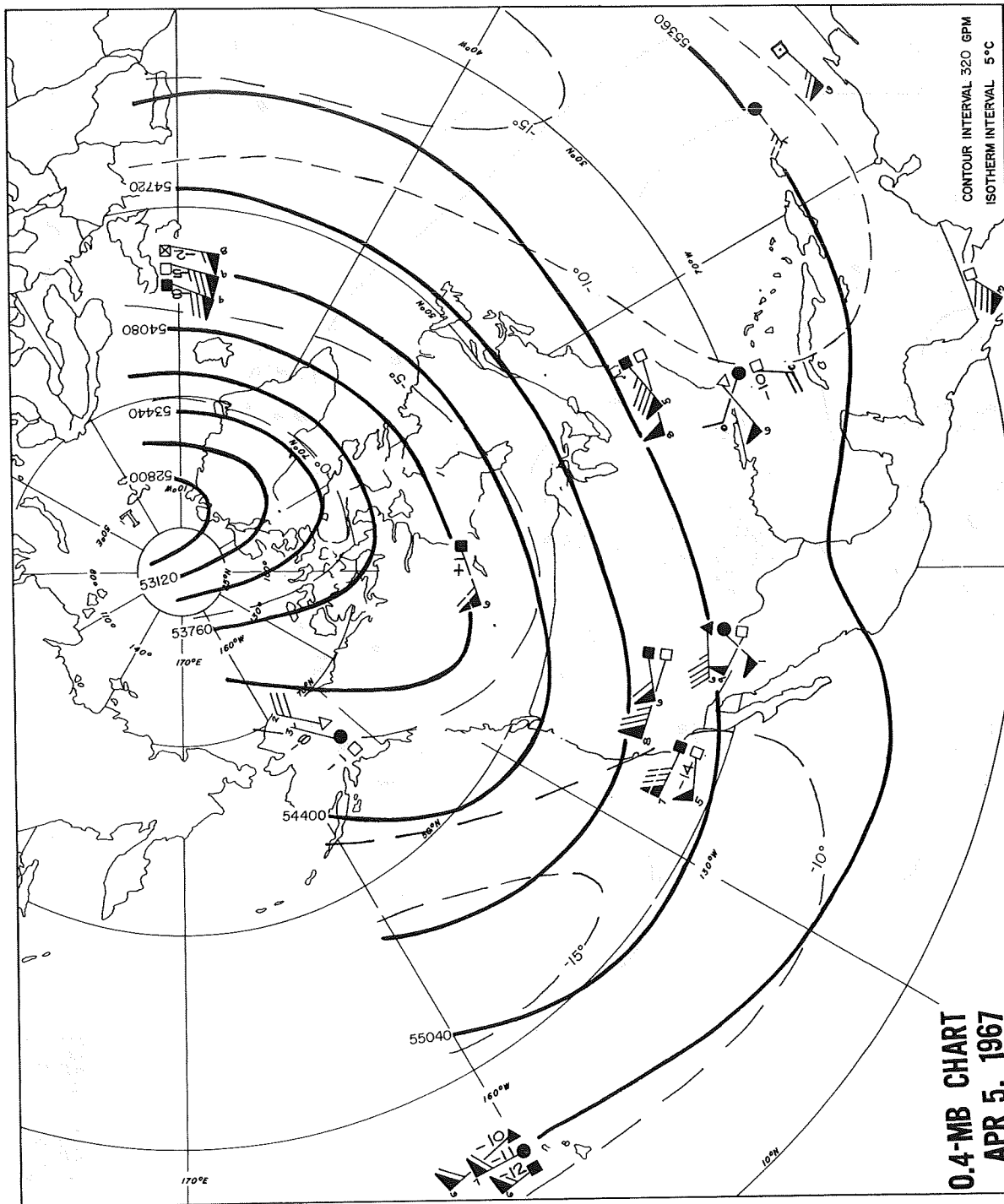




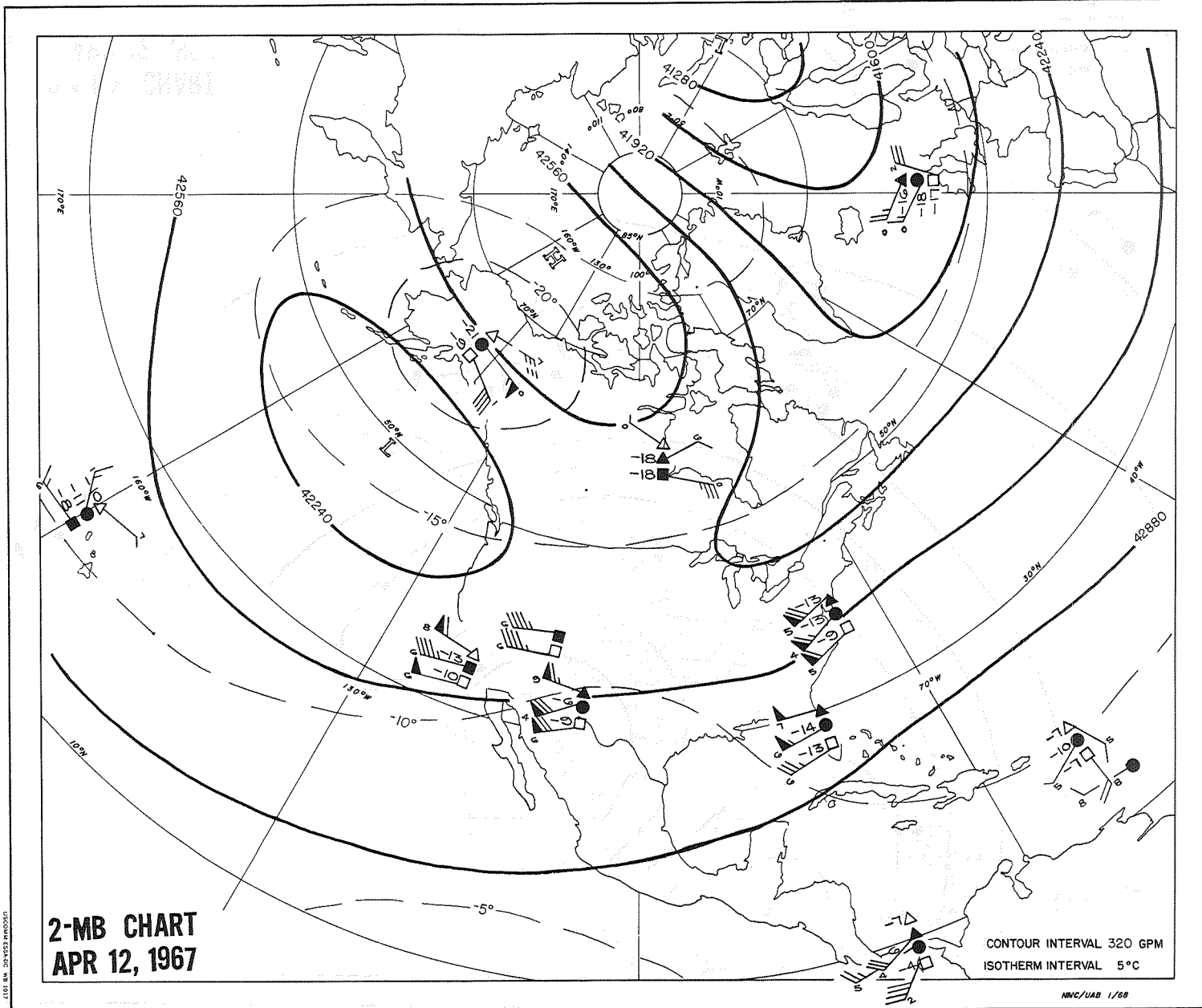
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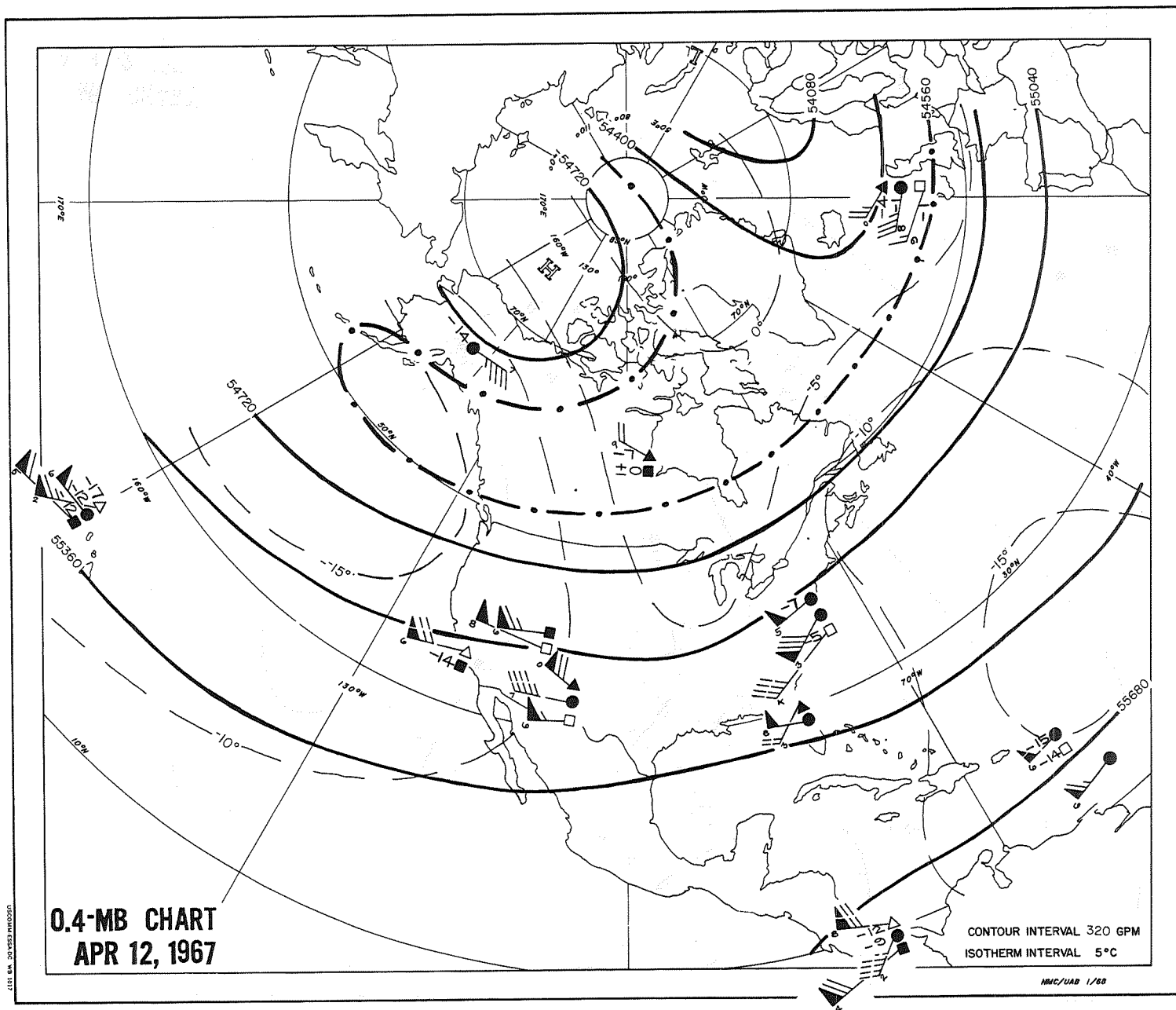


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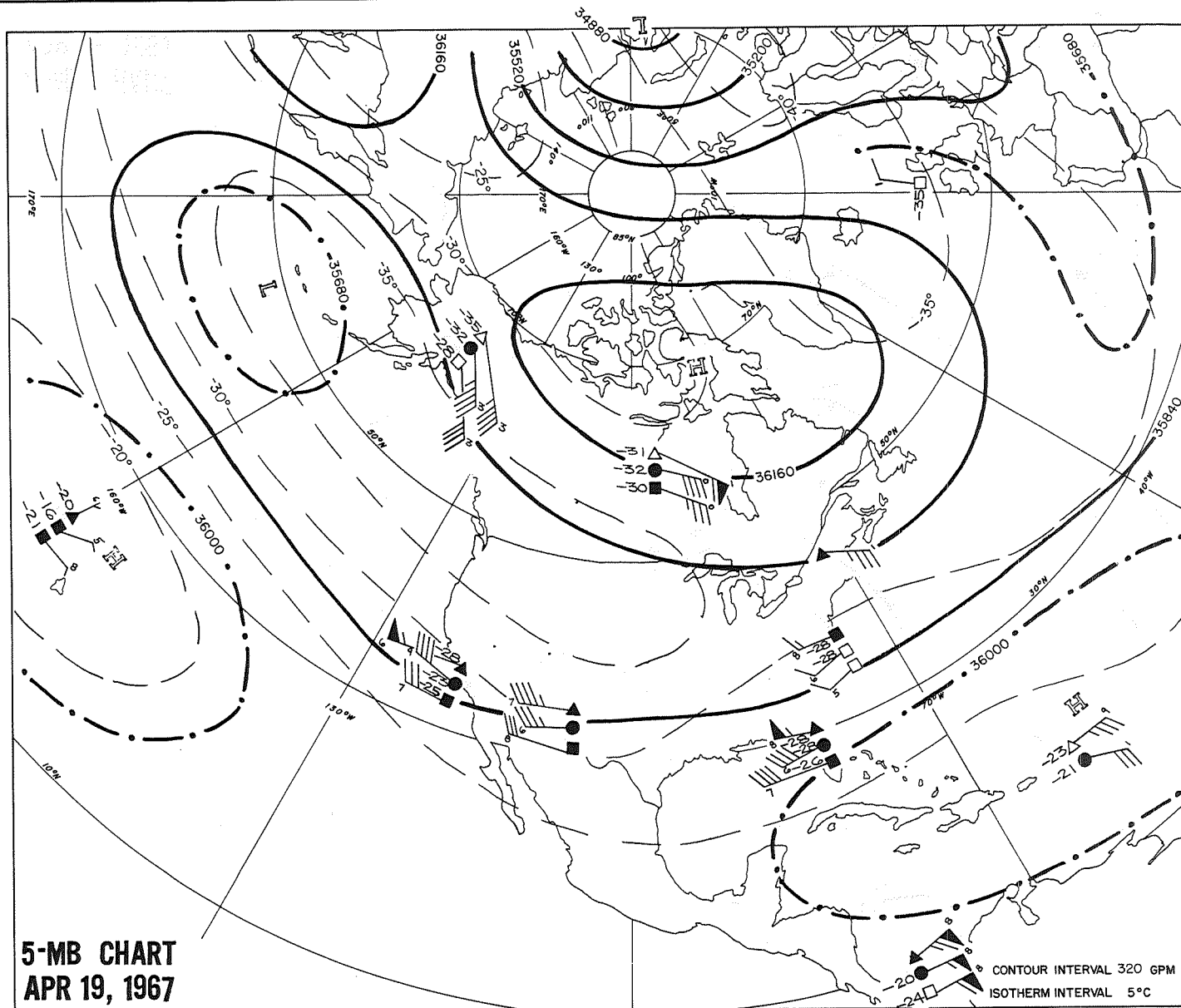
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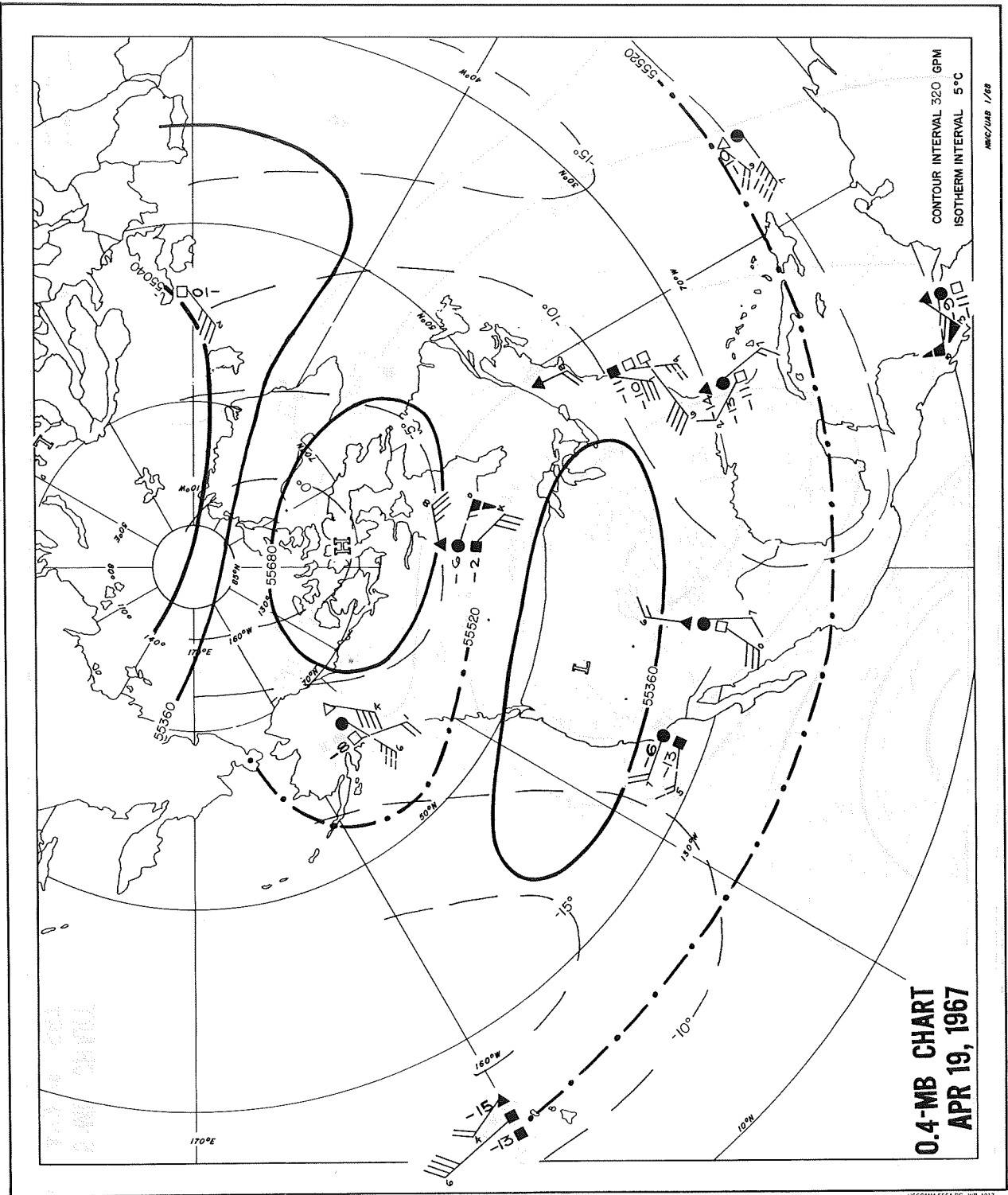


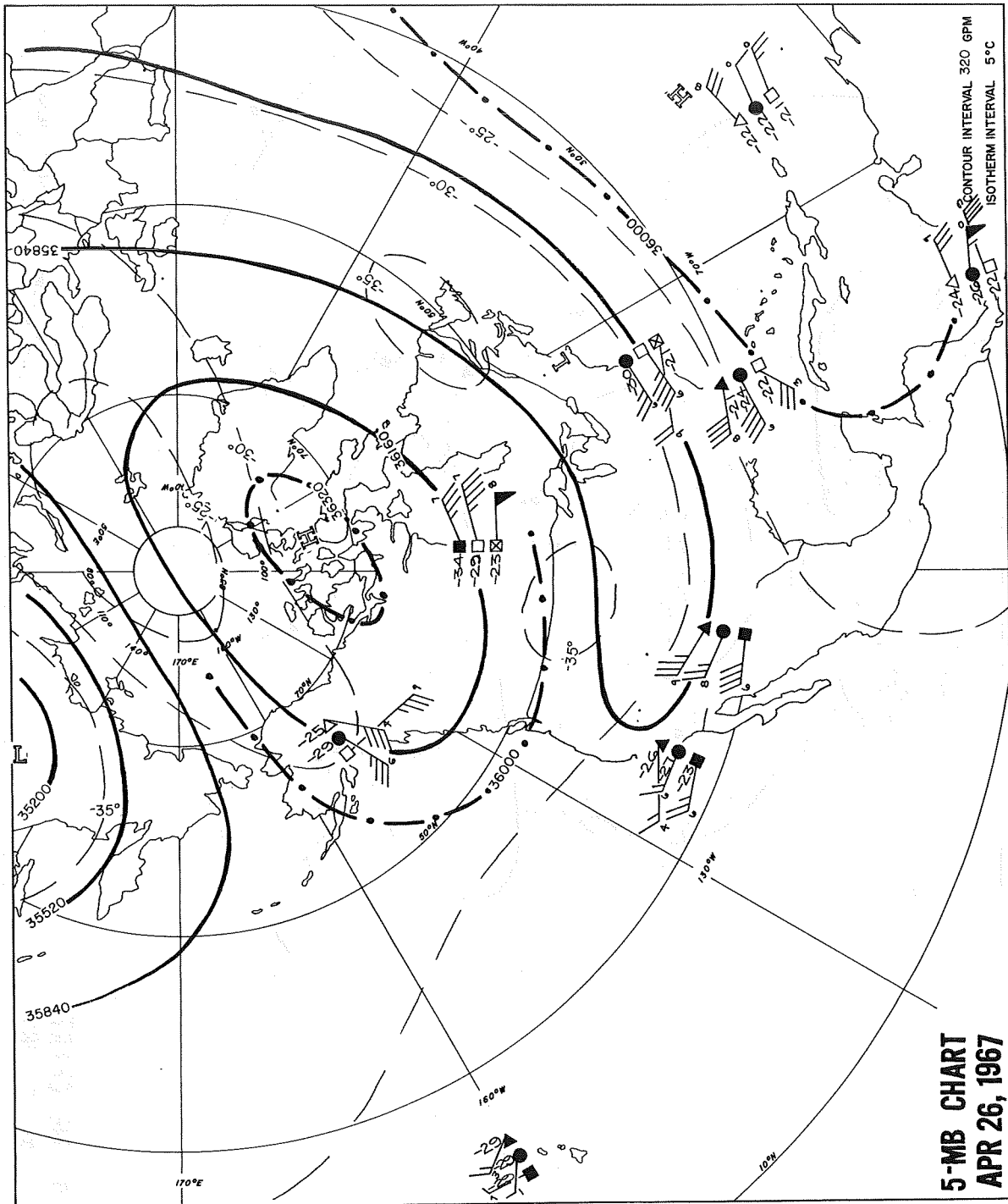
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5-MB CHART
APR 19, 1967



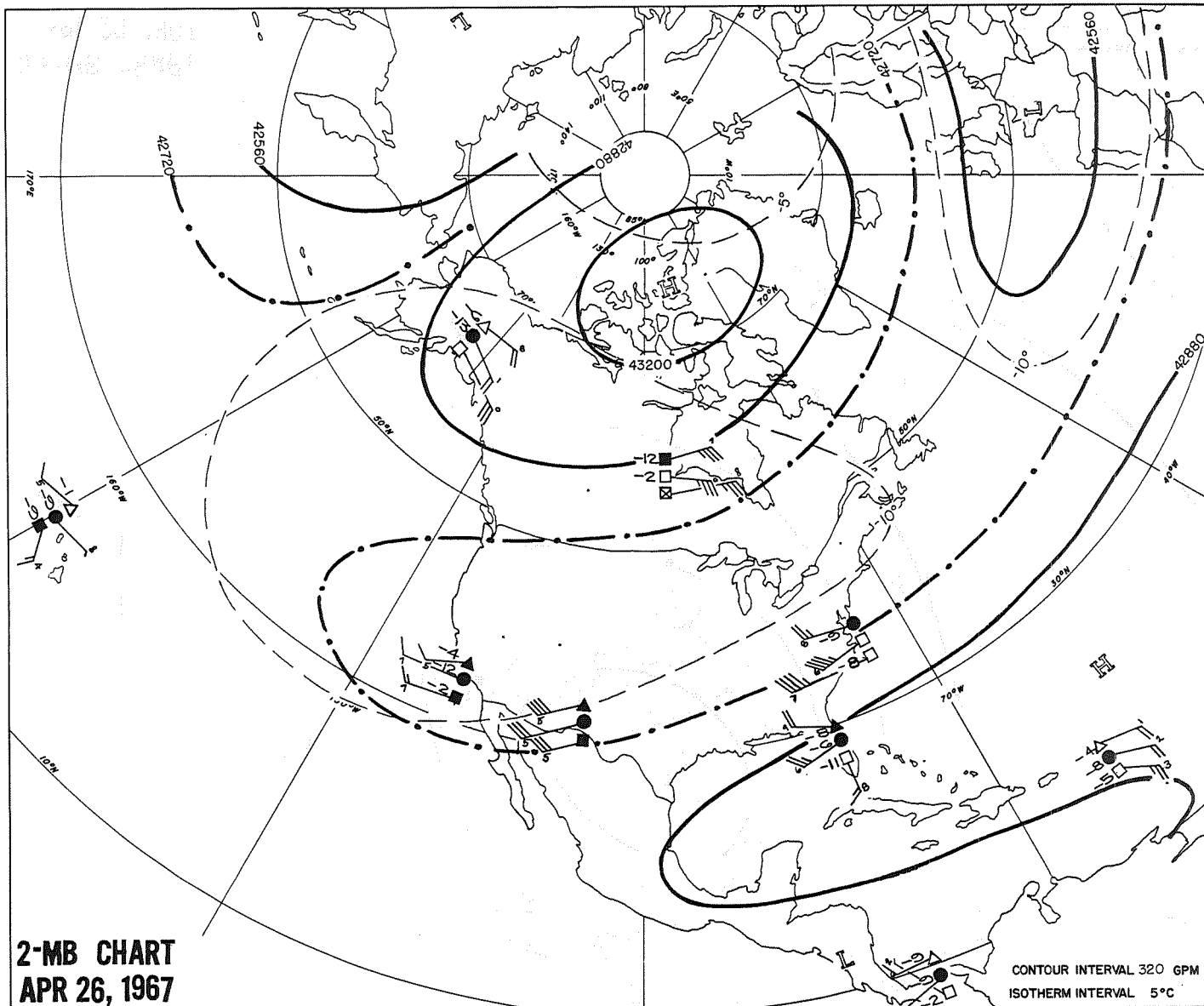
NMIC/UAB 1/68



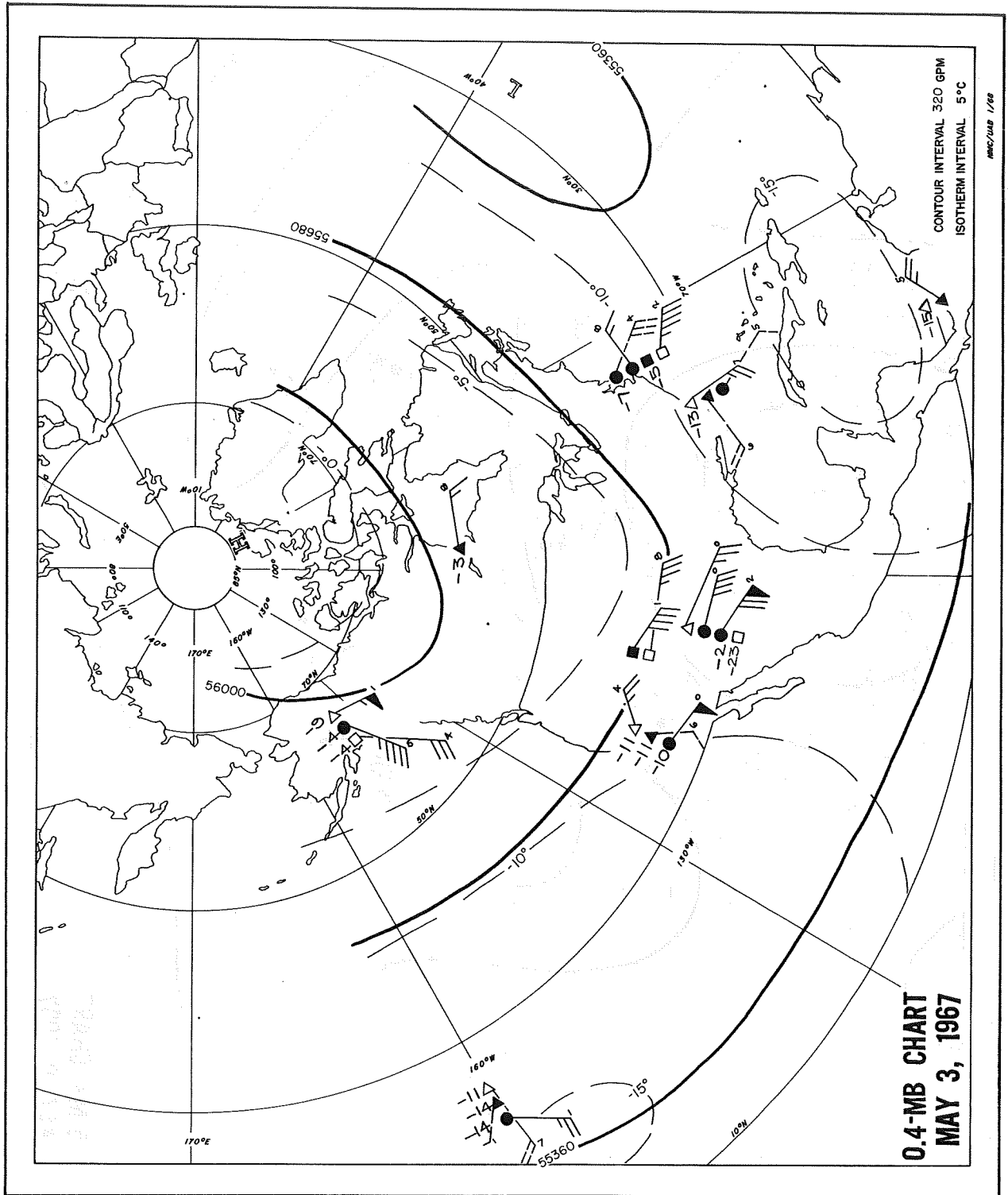


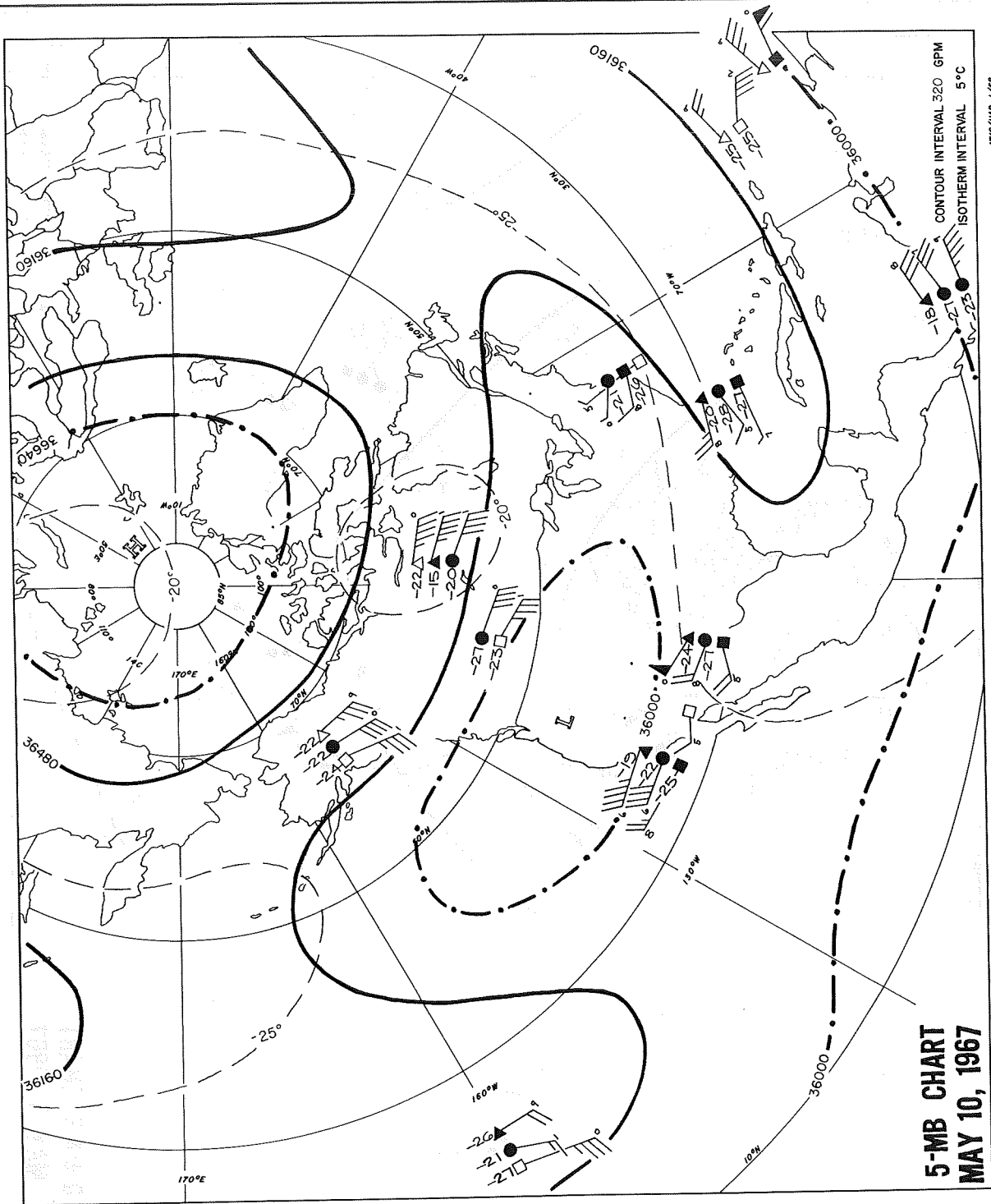
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APR 26, 1967

USCOM-ESSA DC WB 1017

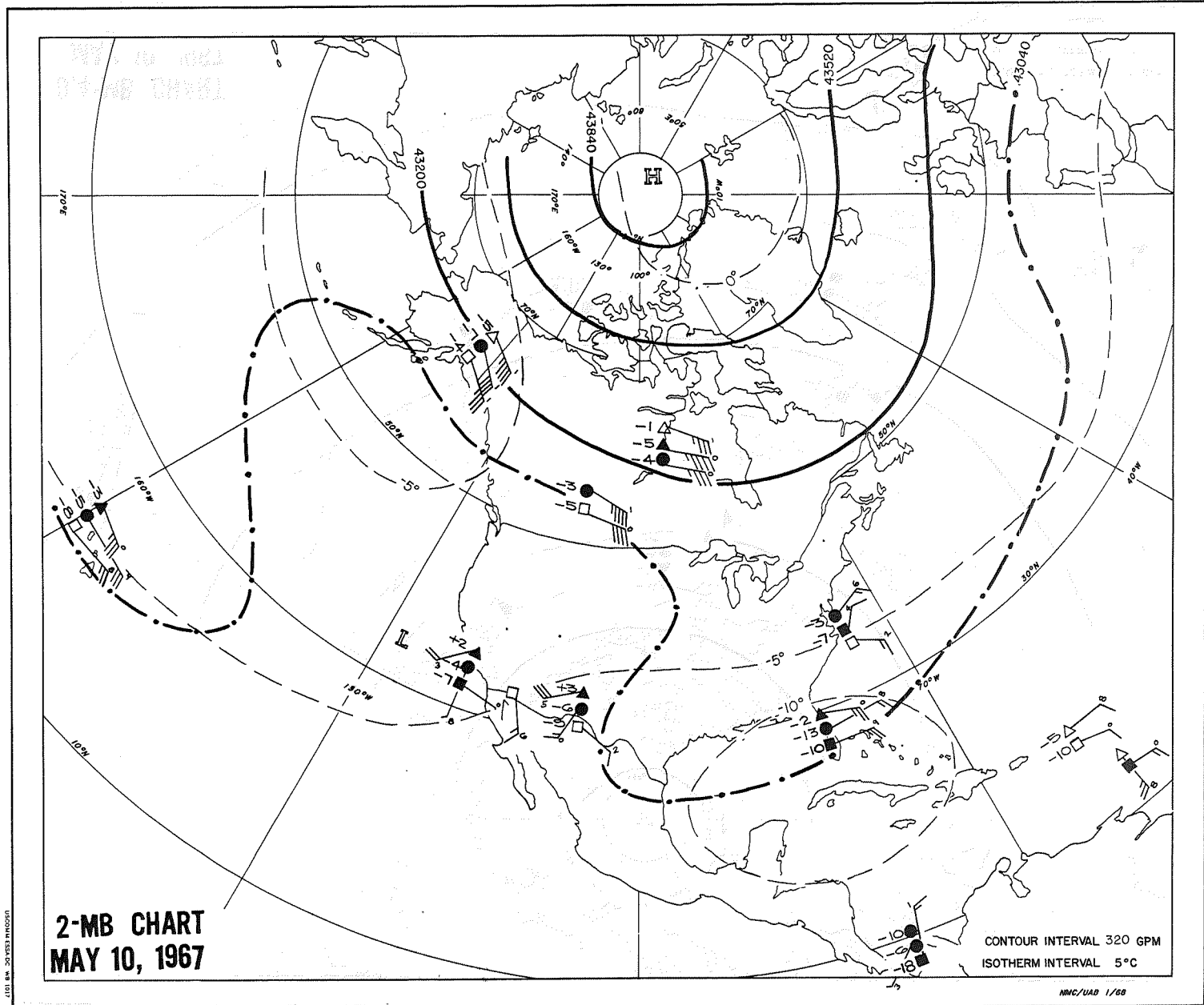


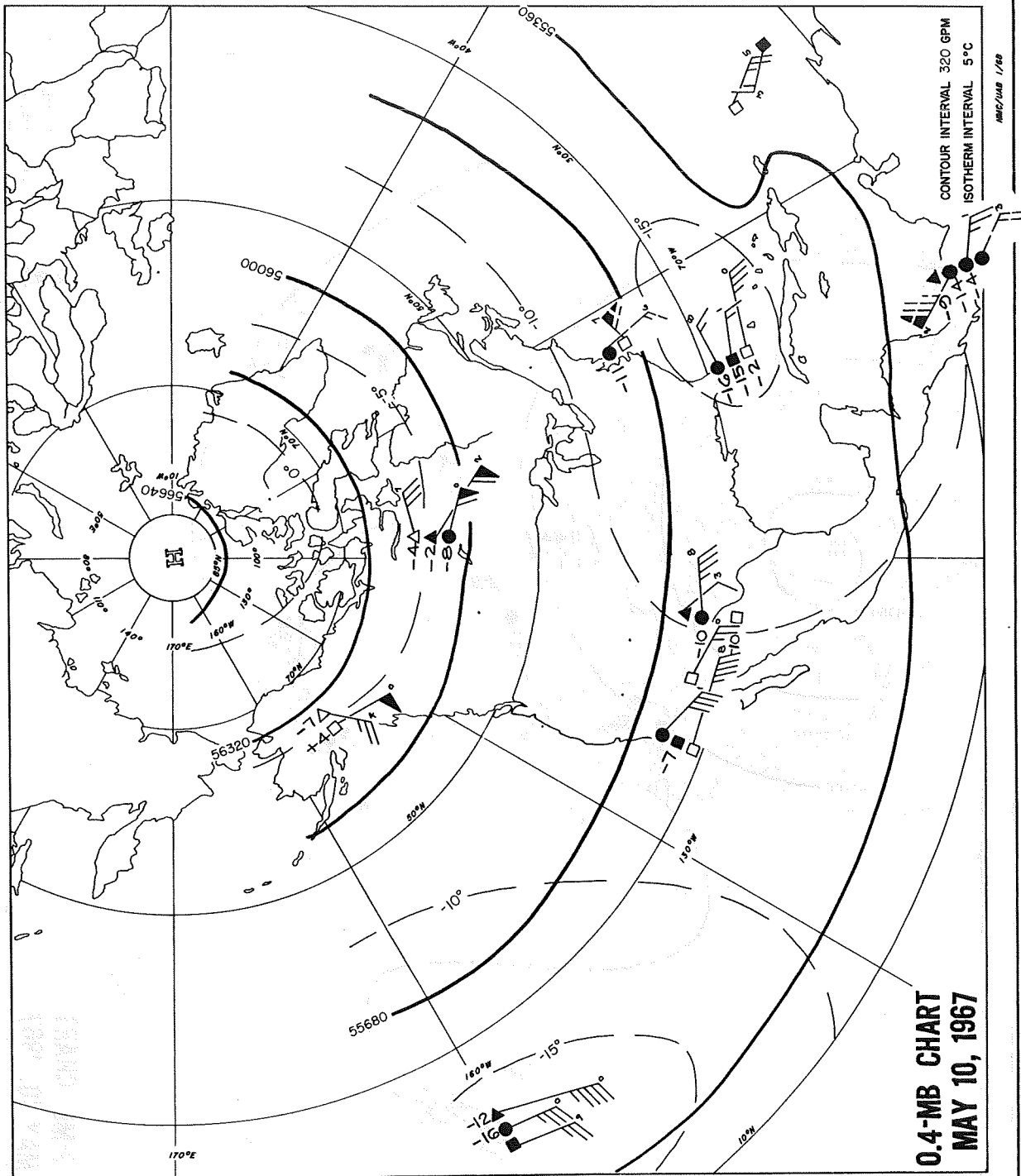




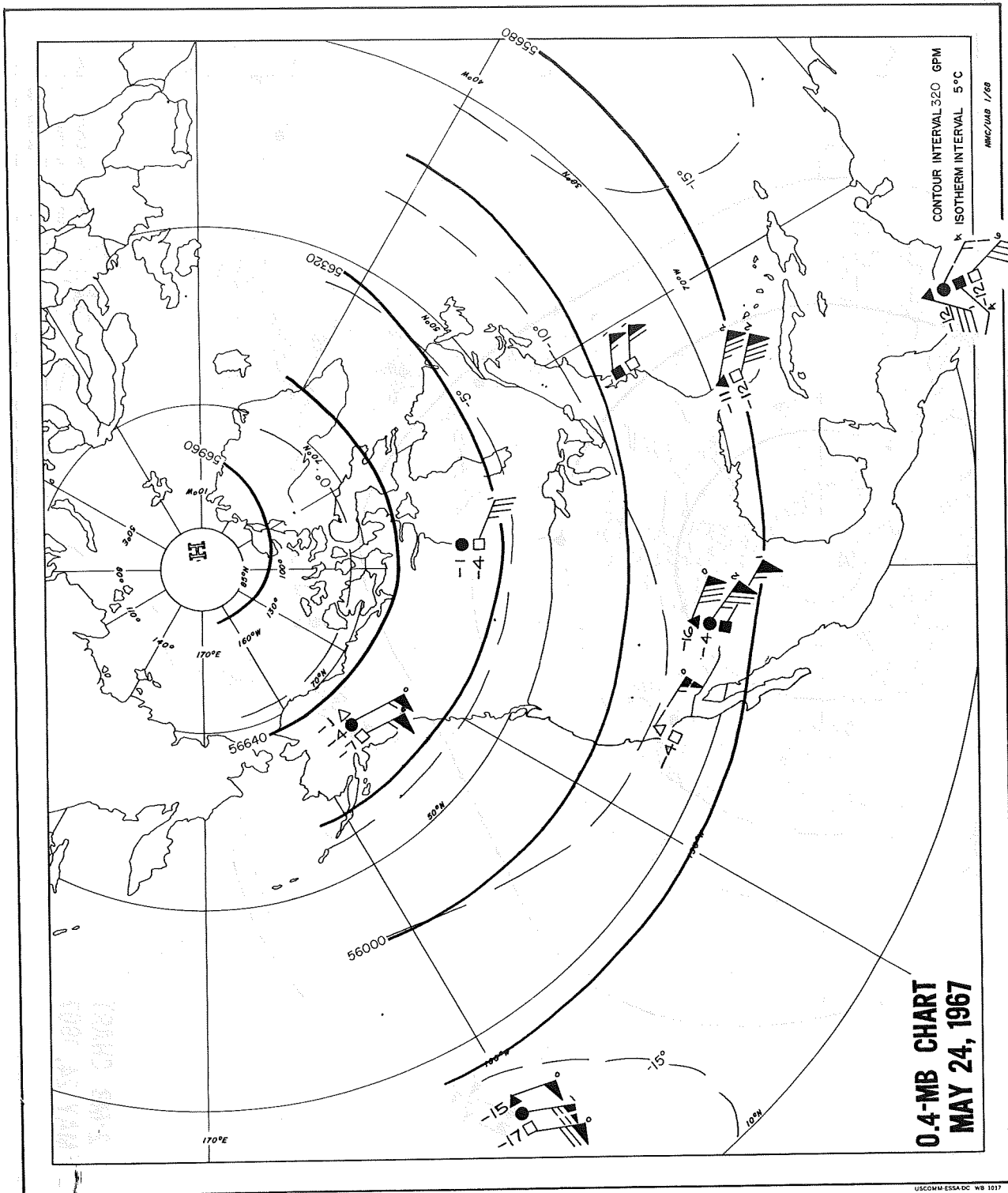


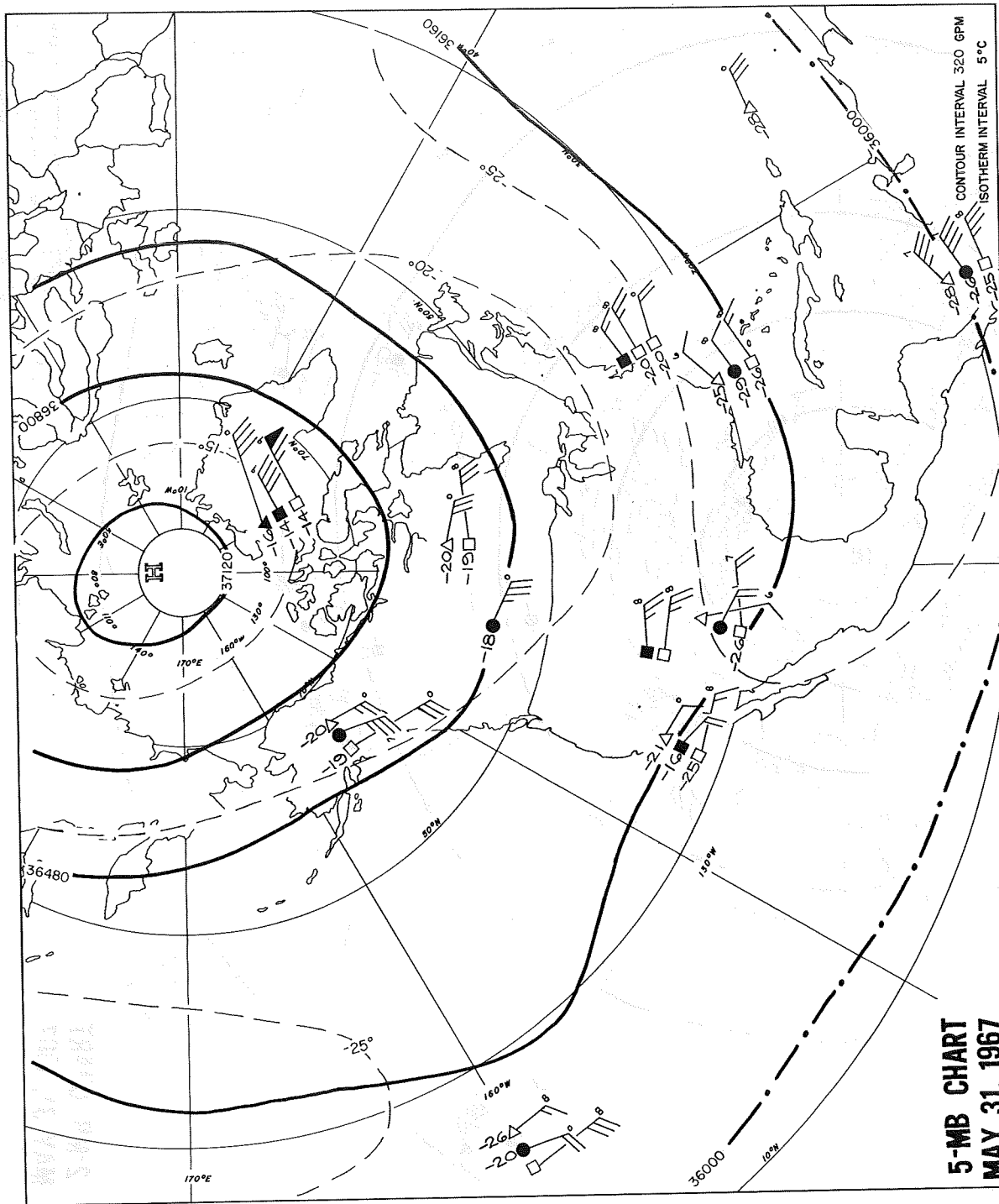
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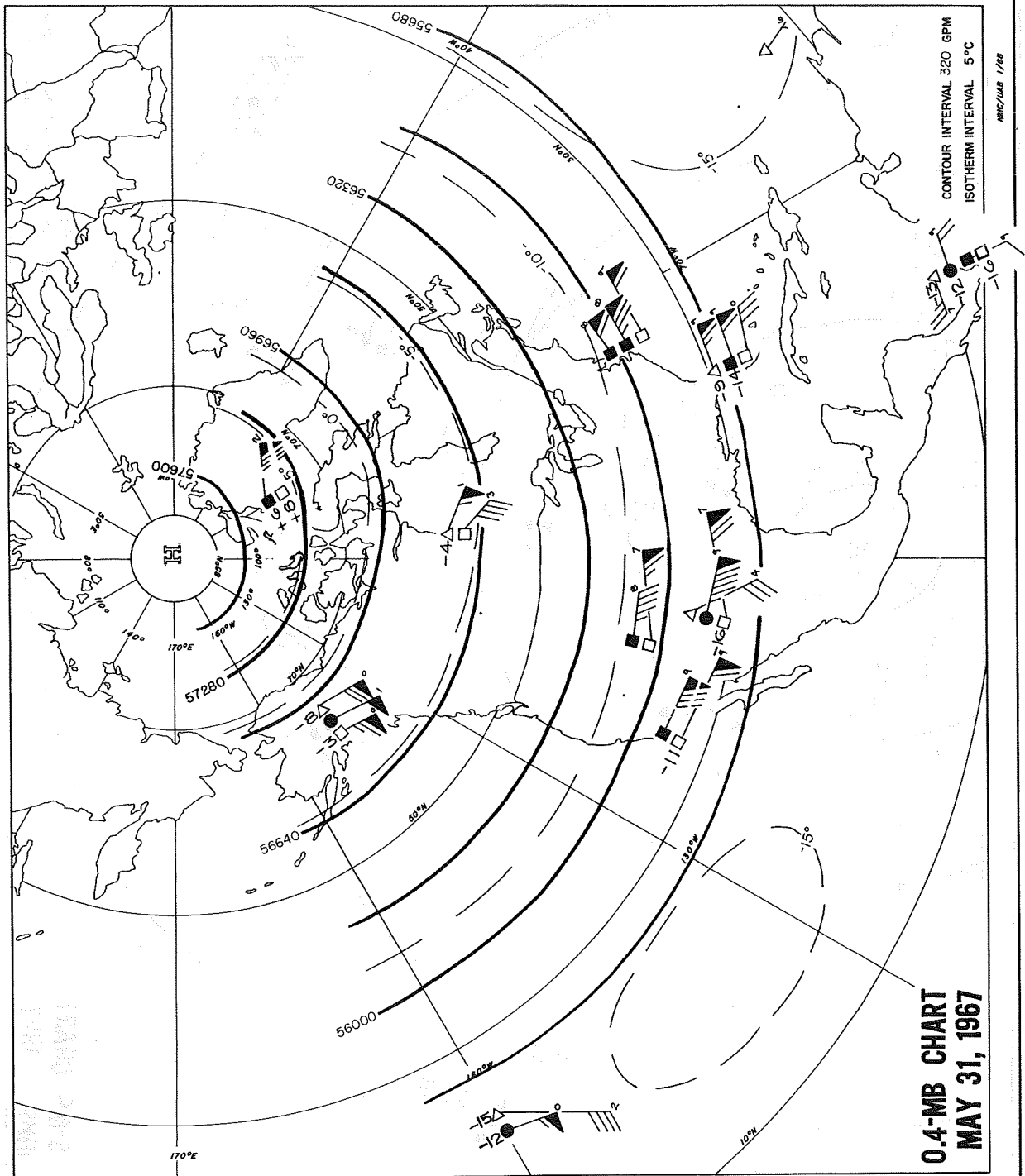
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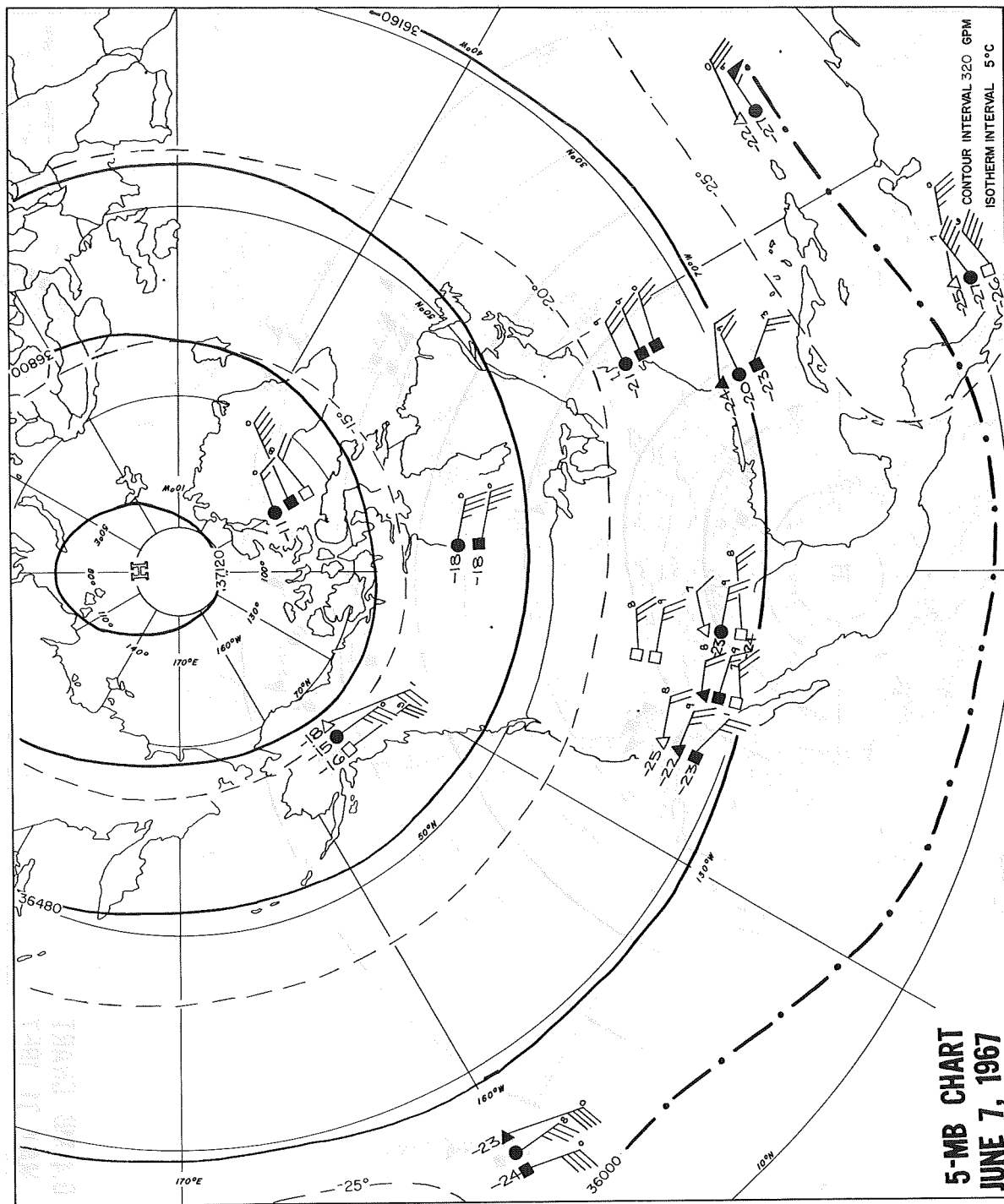




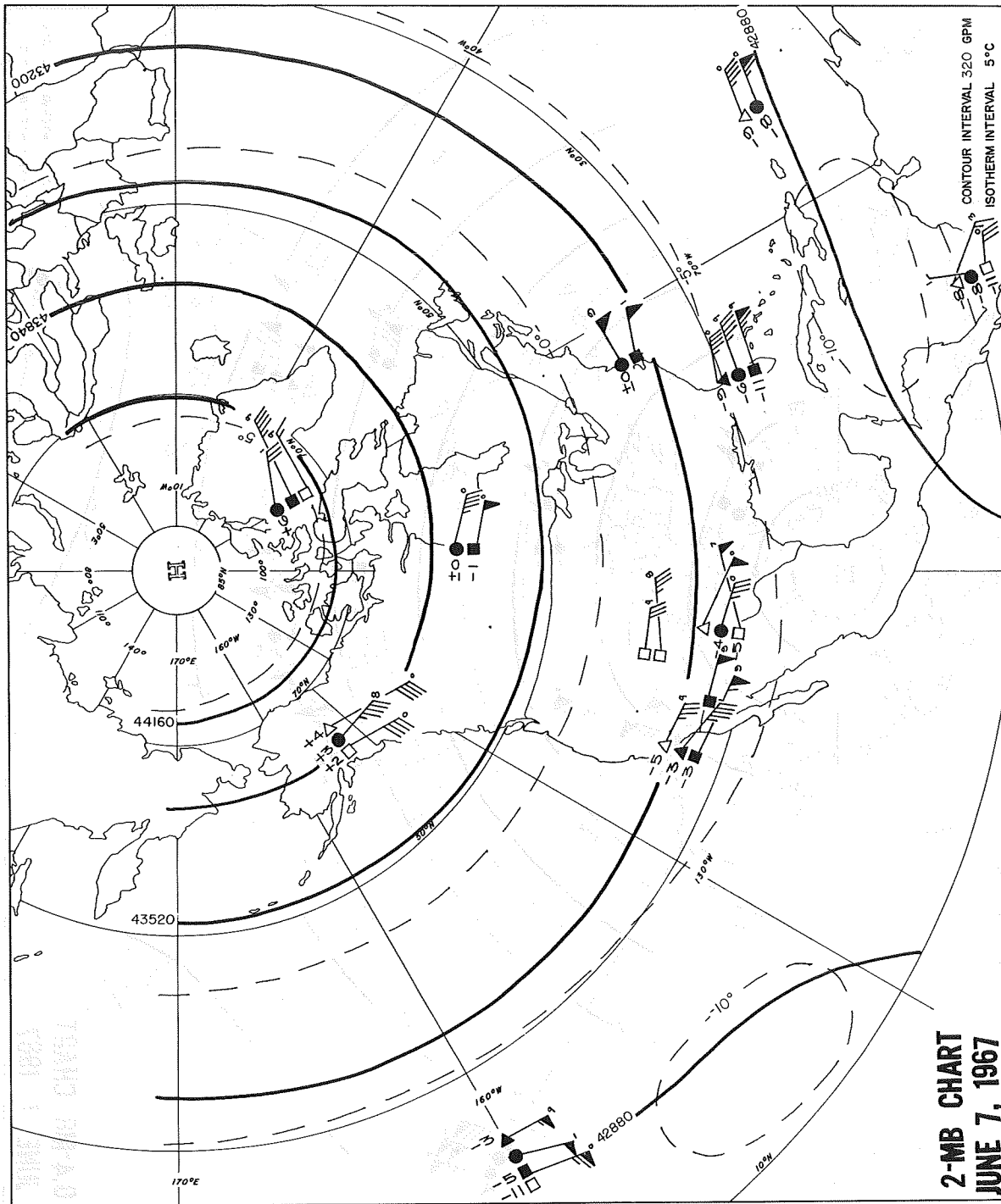
AMC/US 1/68

USCOMM ESSA-DC WB 1017





USCOMM ESSA-DC WB 1017

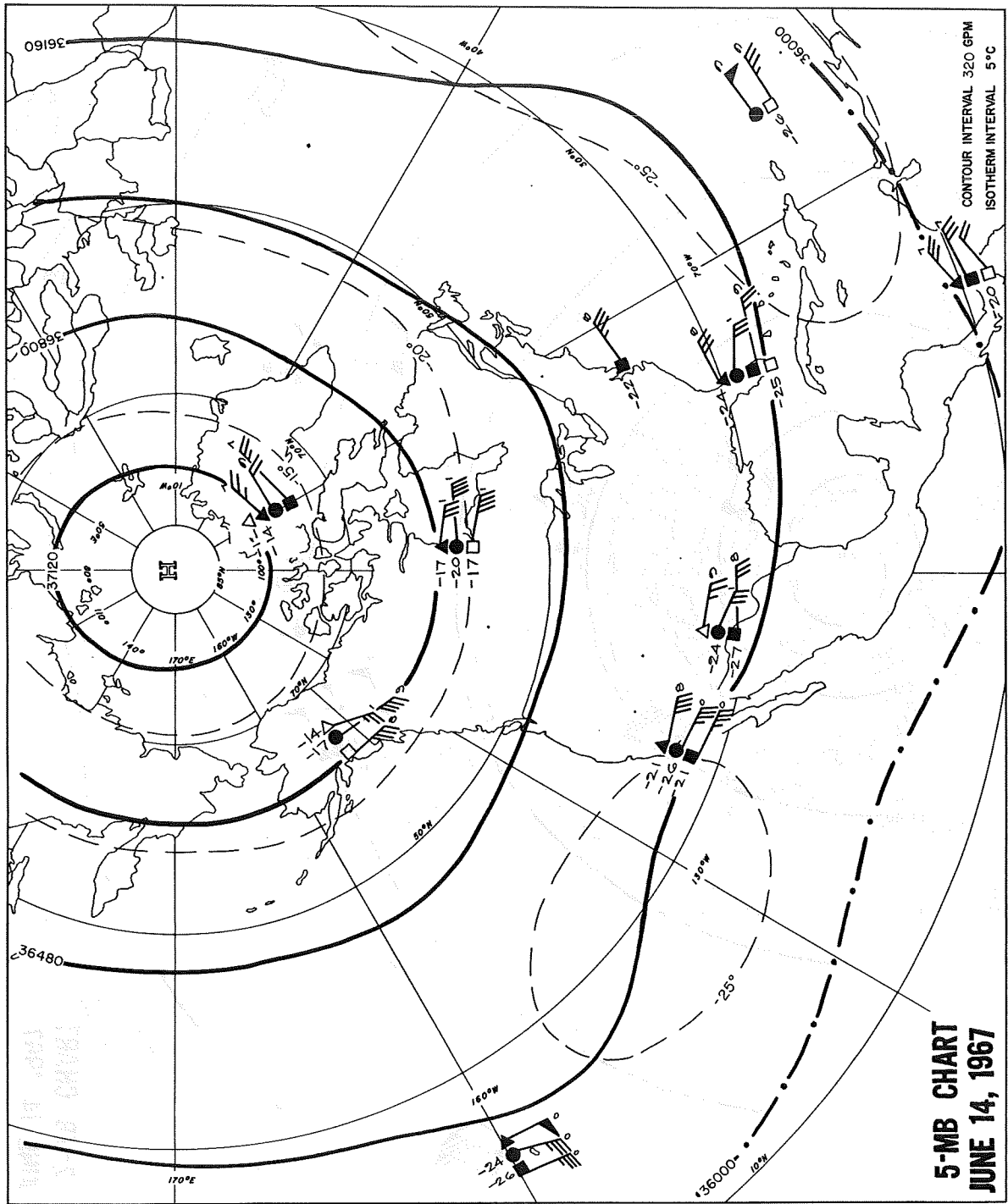


**2-MB CHART
JUNE 7, 1967**

CONTOUR INTERVAL 320 GPM
ISOTHERM INTERVAL 5°C

NMC/US 1/68

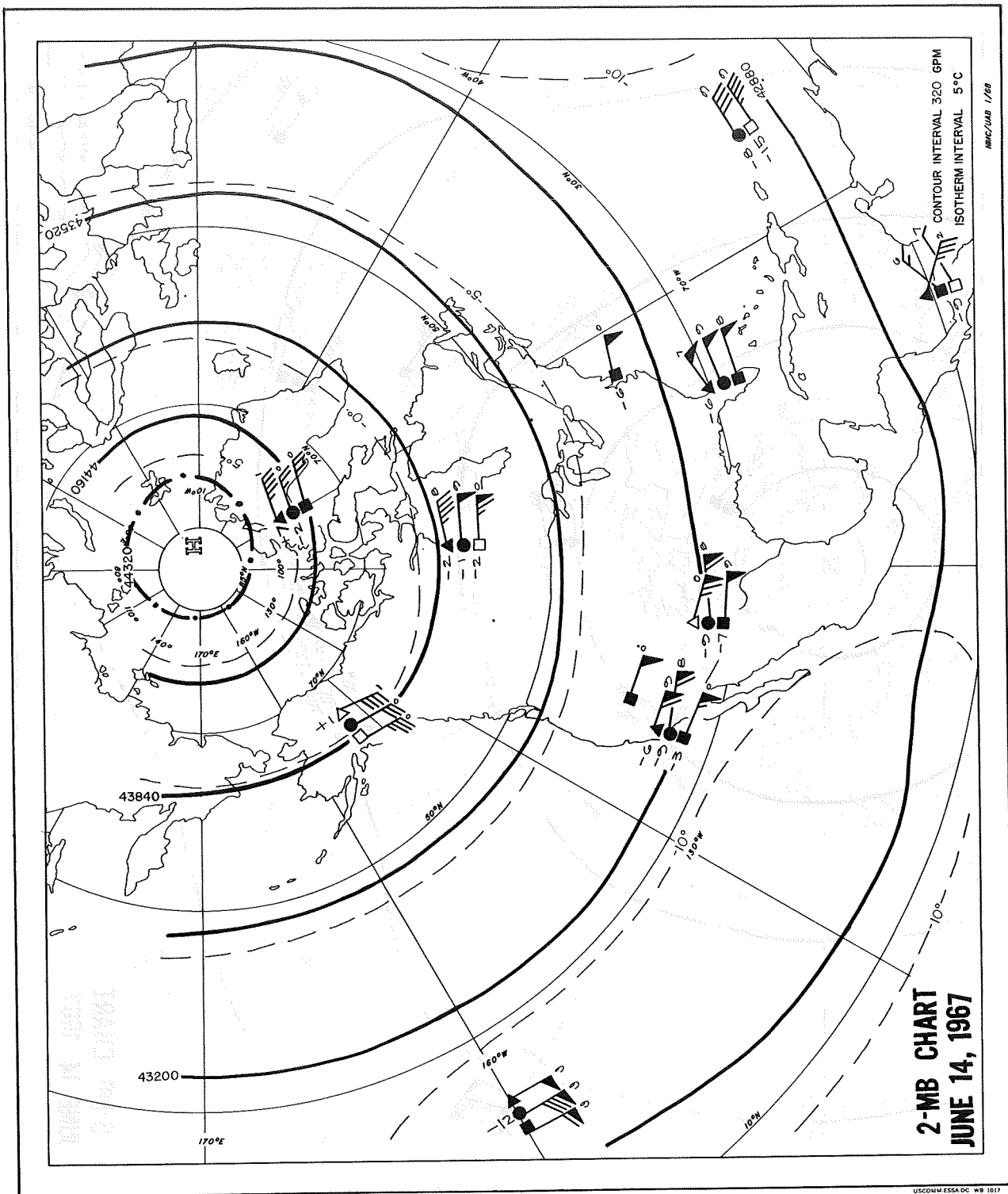
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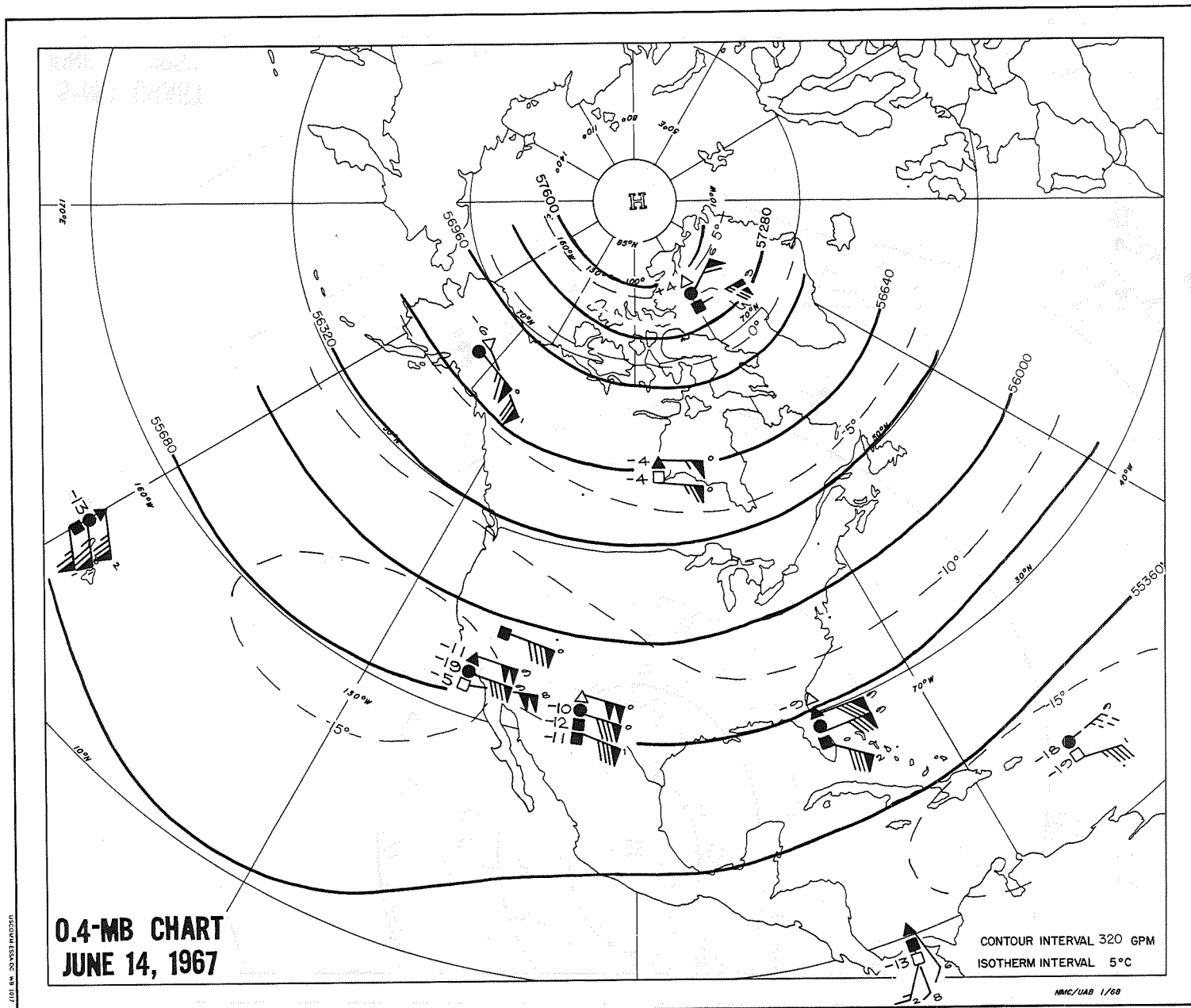


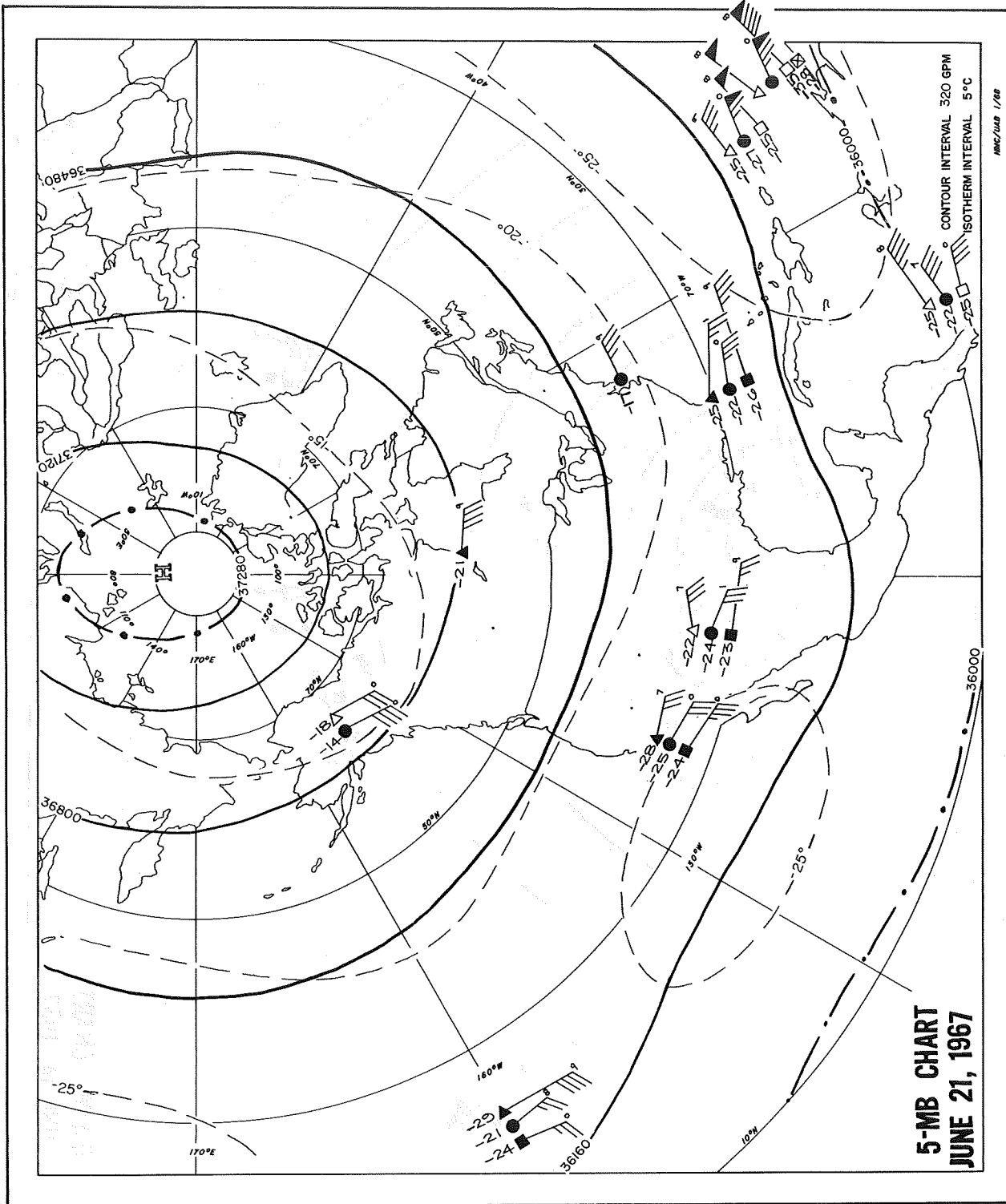
CONTOUR INTERVAL 320 GPM
ISOTHERM INTERVAL 5°C

5-MB CHART
JUNE 14, 1967

USCOMM ESSA DC WB 1017

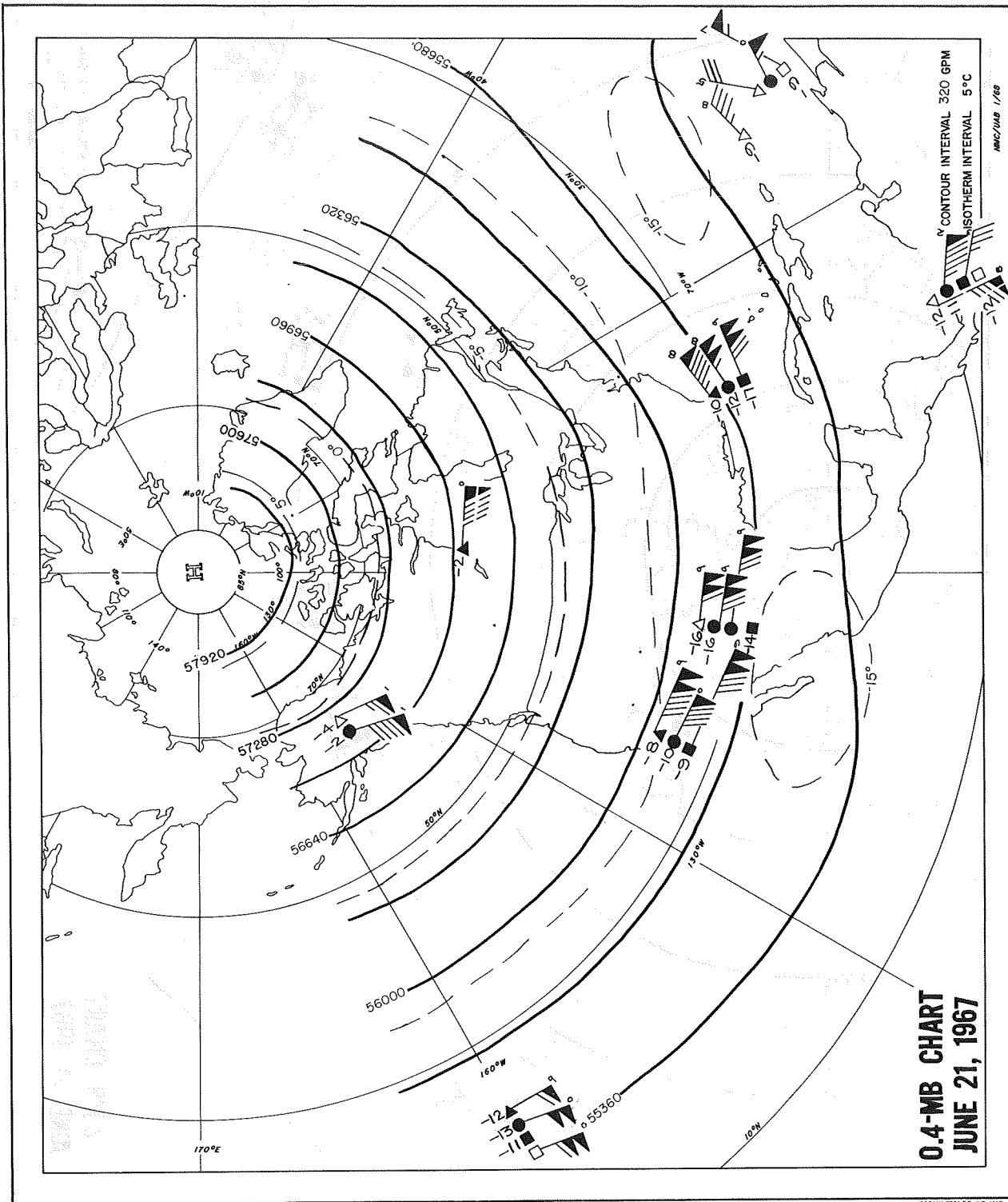




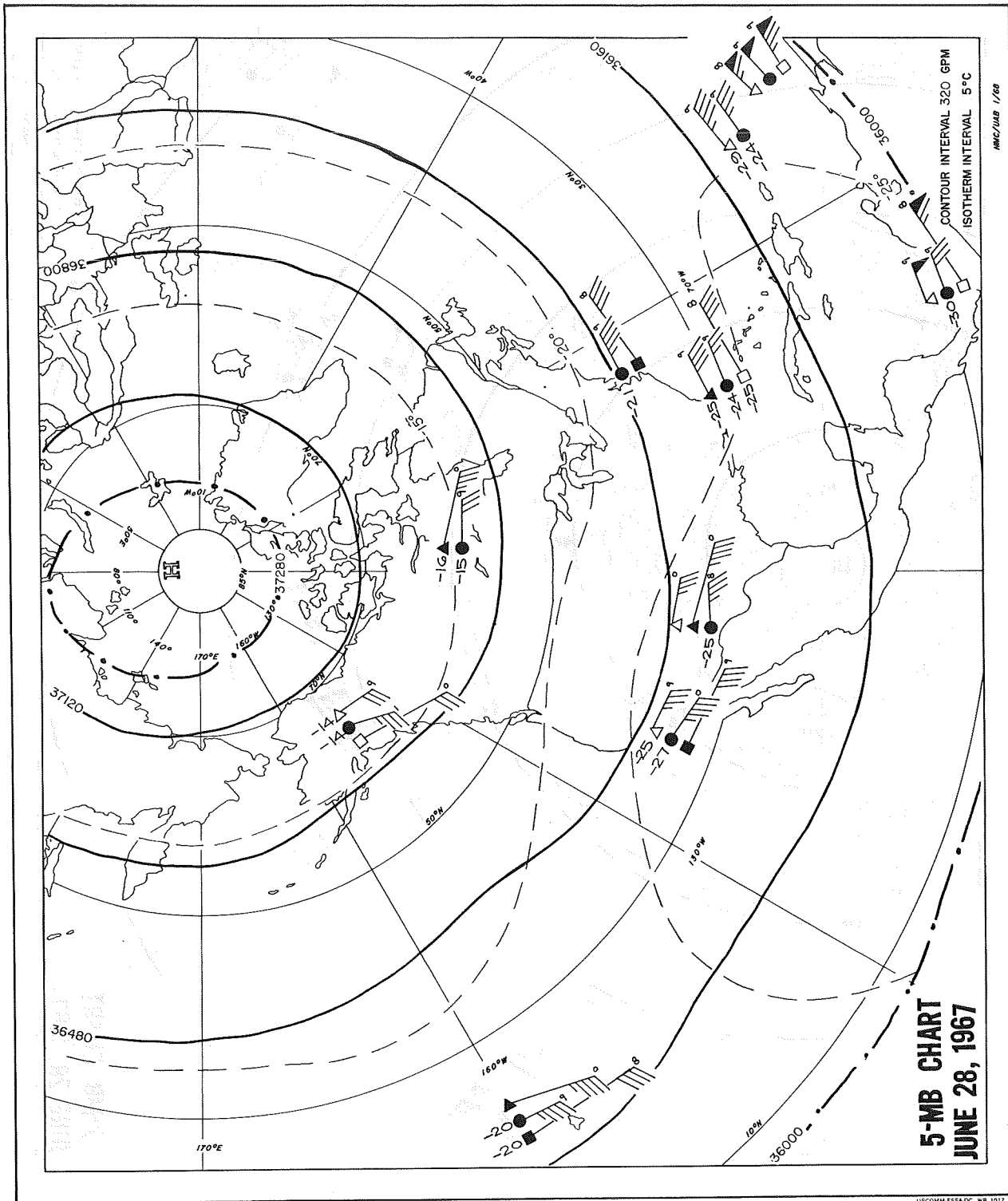


**5-MB CHART
JUNE 21, 1967**

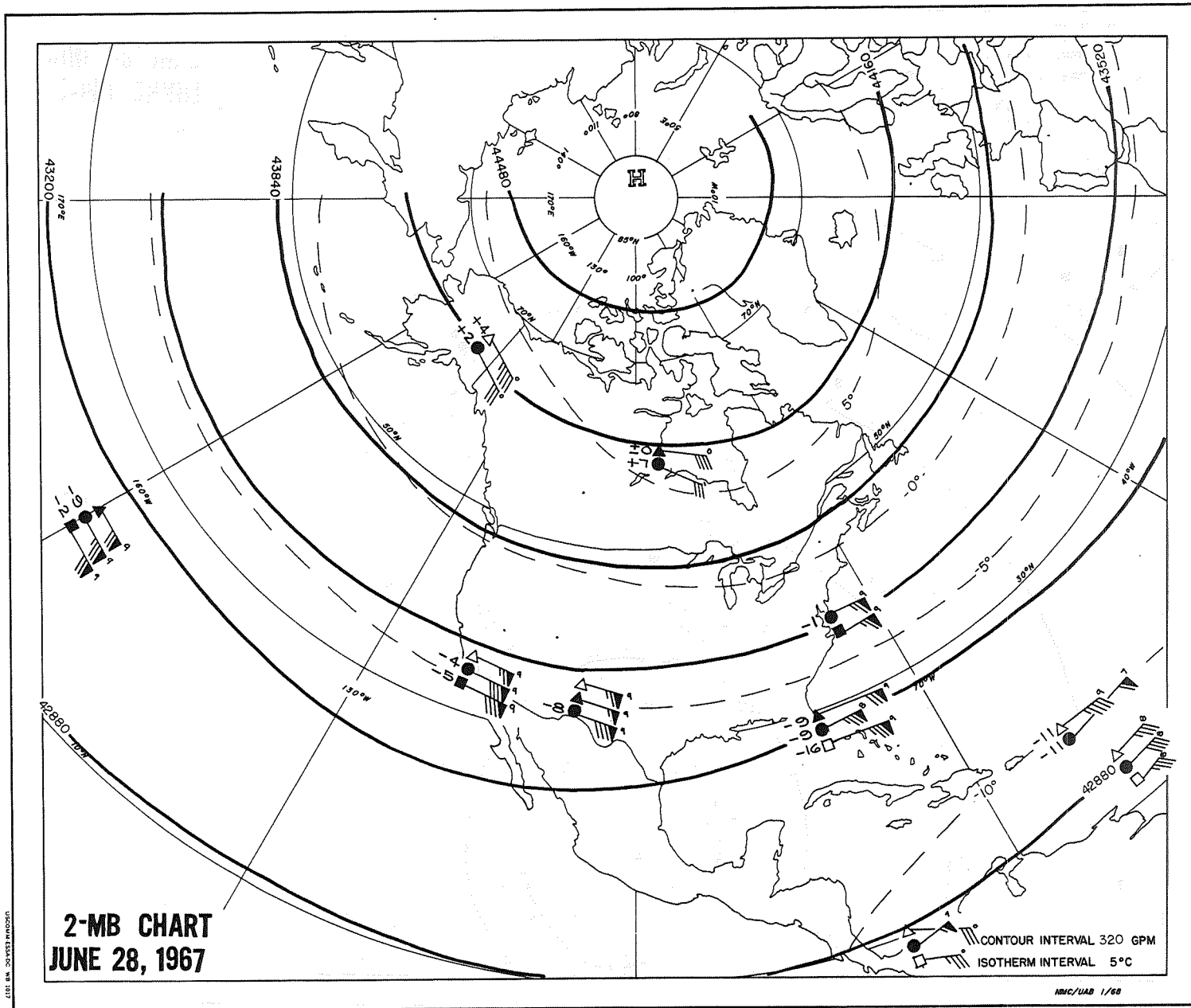
USCOMM-ESSA-DC WB 1017

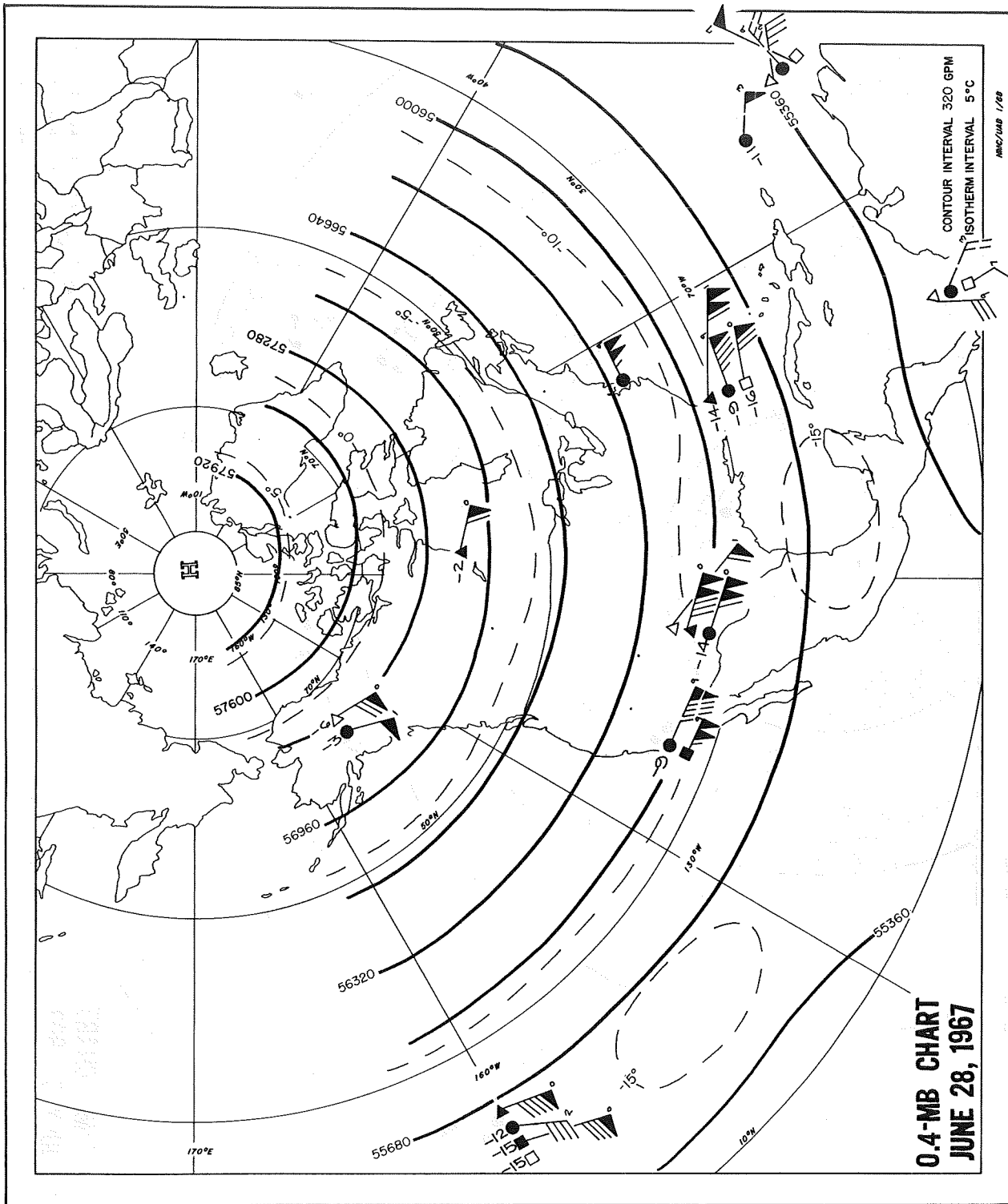


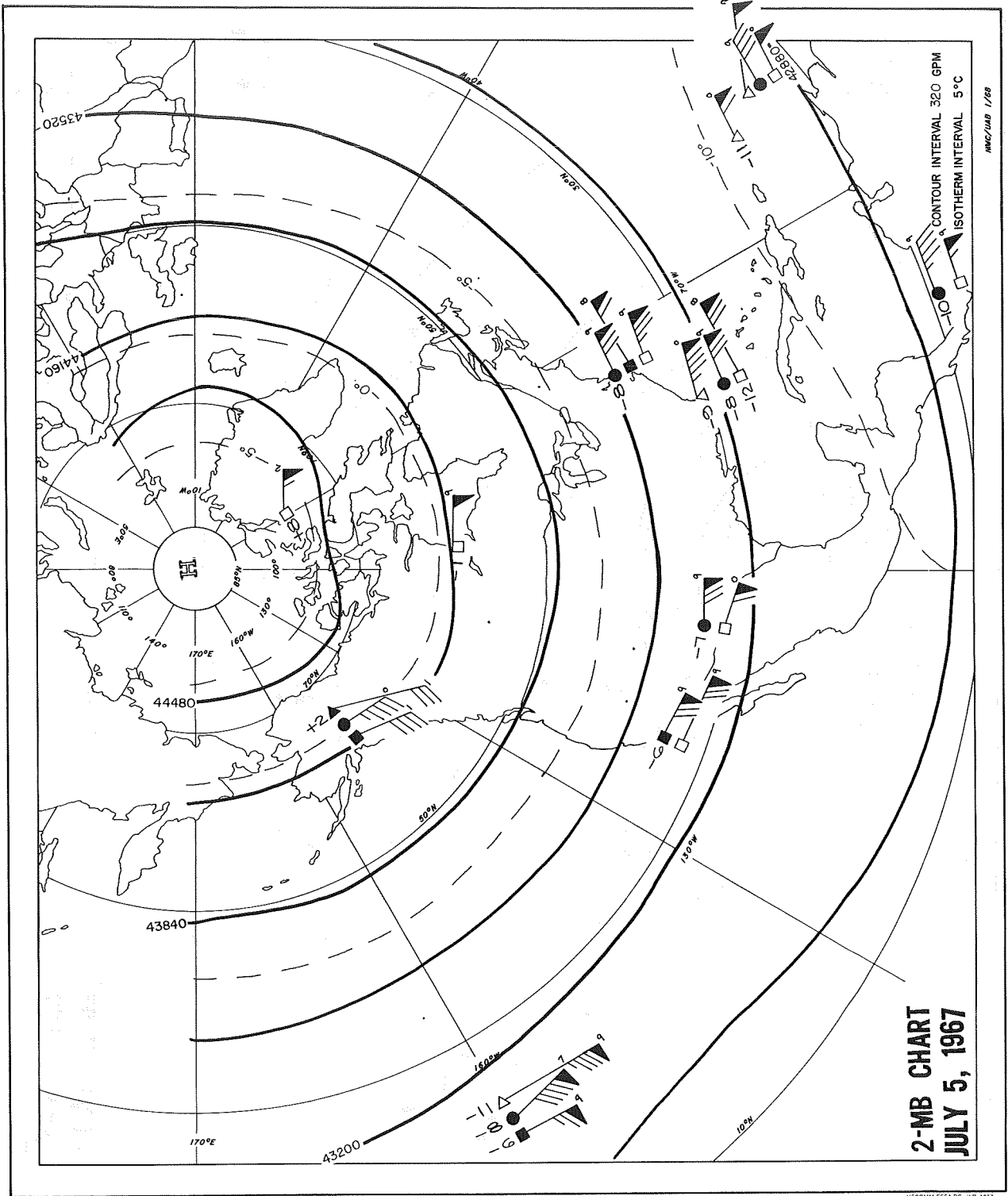
USCOMM-ESSA-DC WB 1017



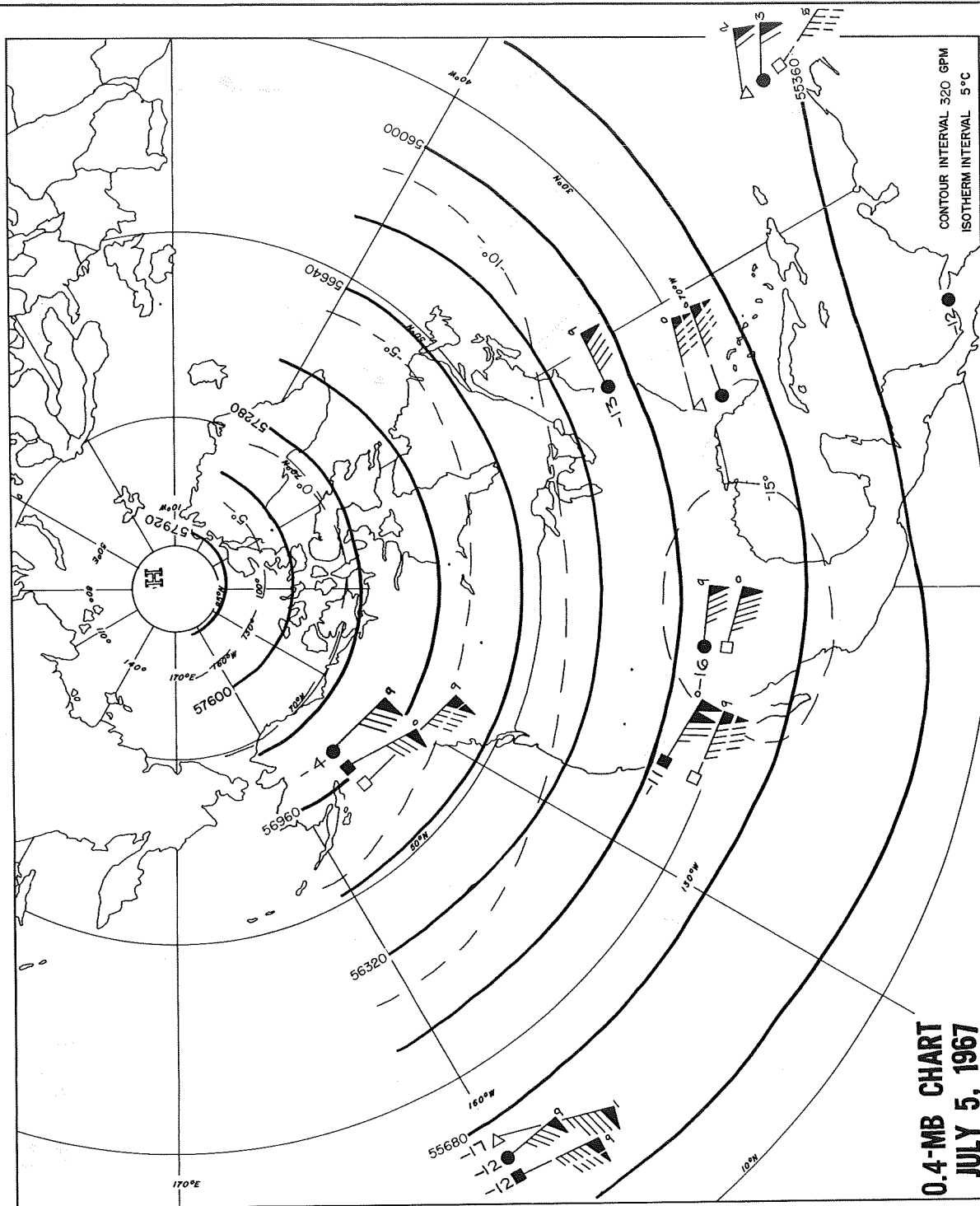
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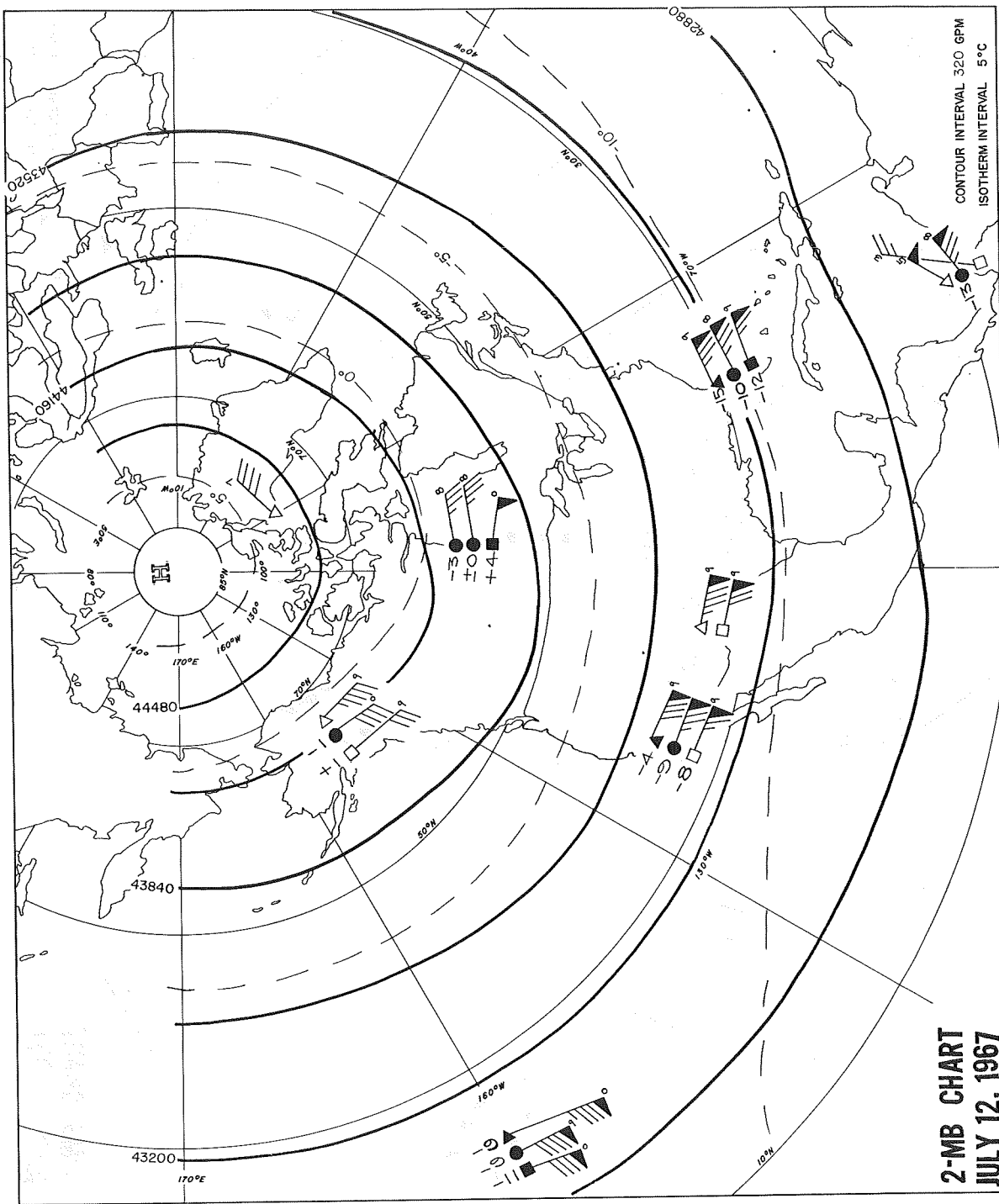






USCOWM C55A-DC WB 1017



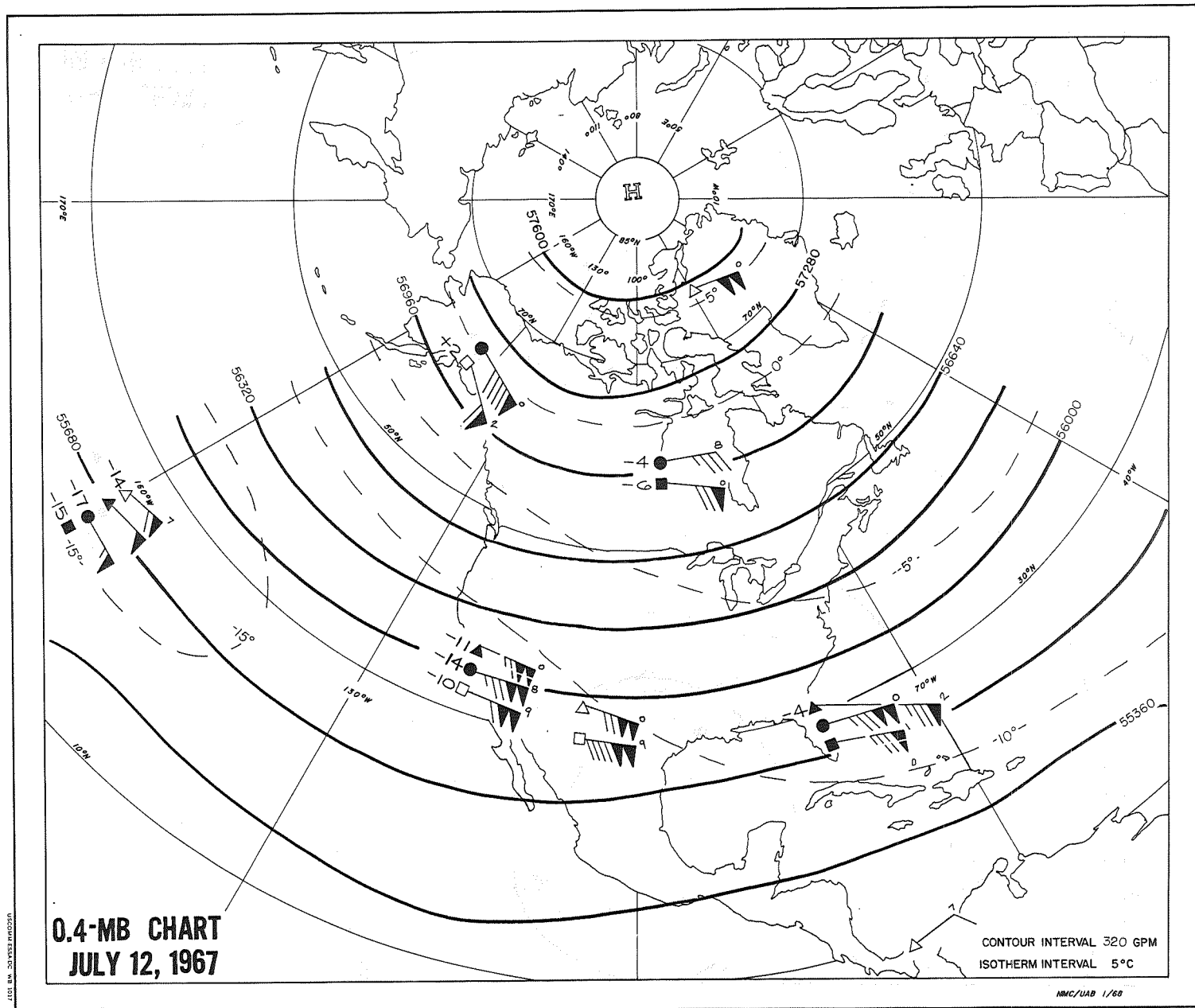


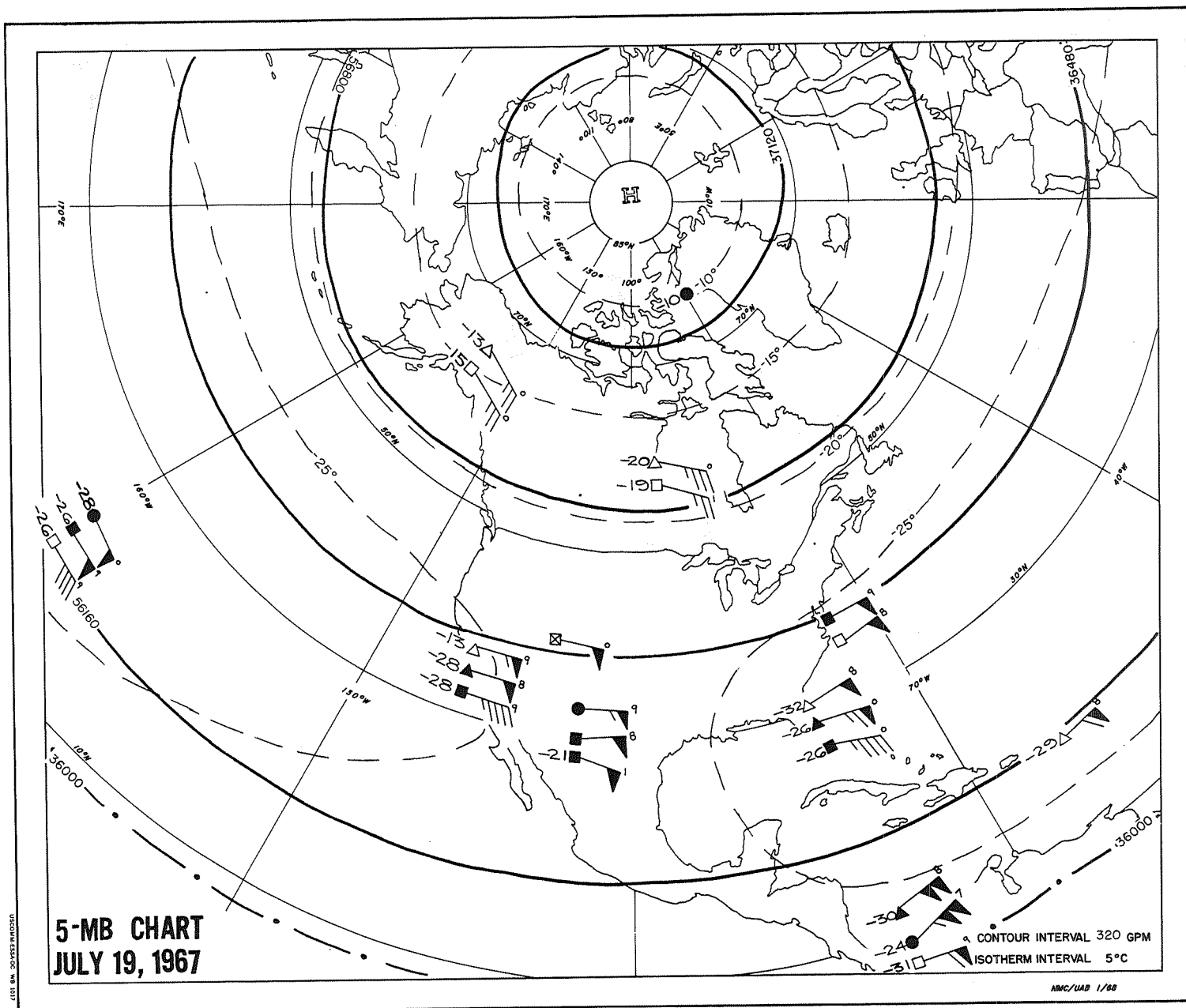
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JULY 12, 1967**

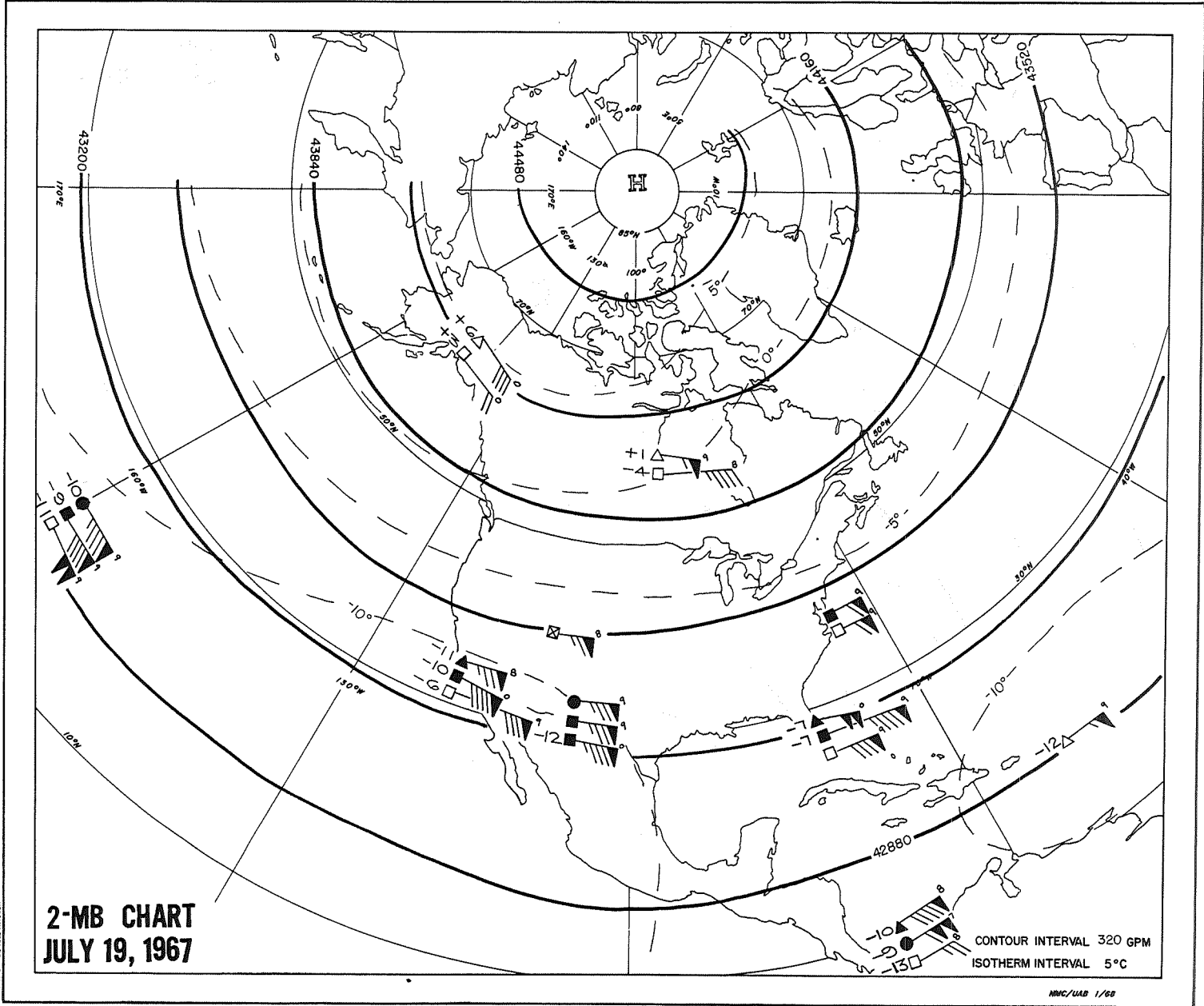
CONTOUR INTERVAL 320 GPM
ISOTHERM INTERVAL 5°C

NWS/NOAA 1/68

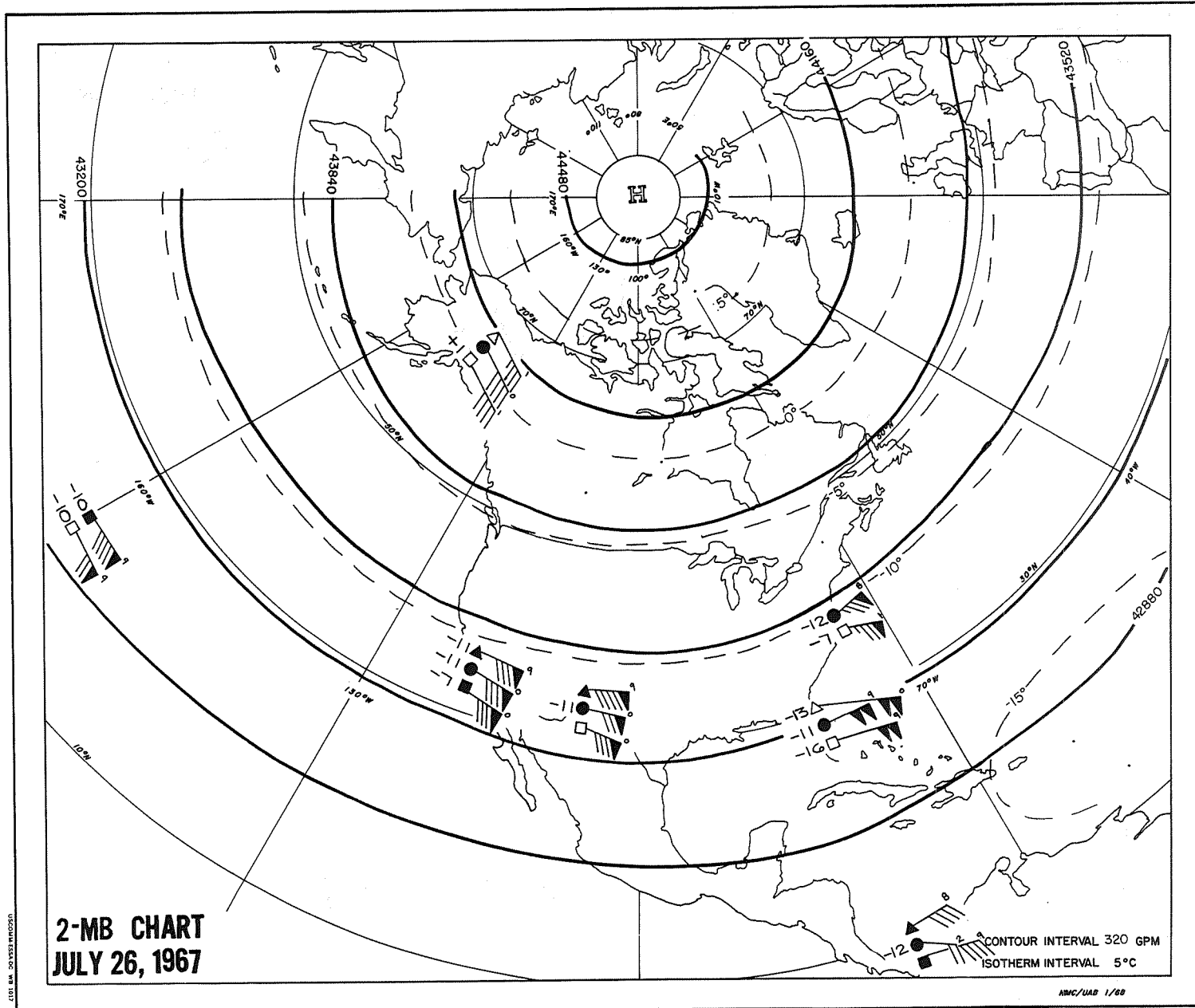
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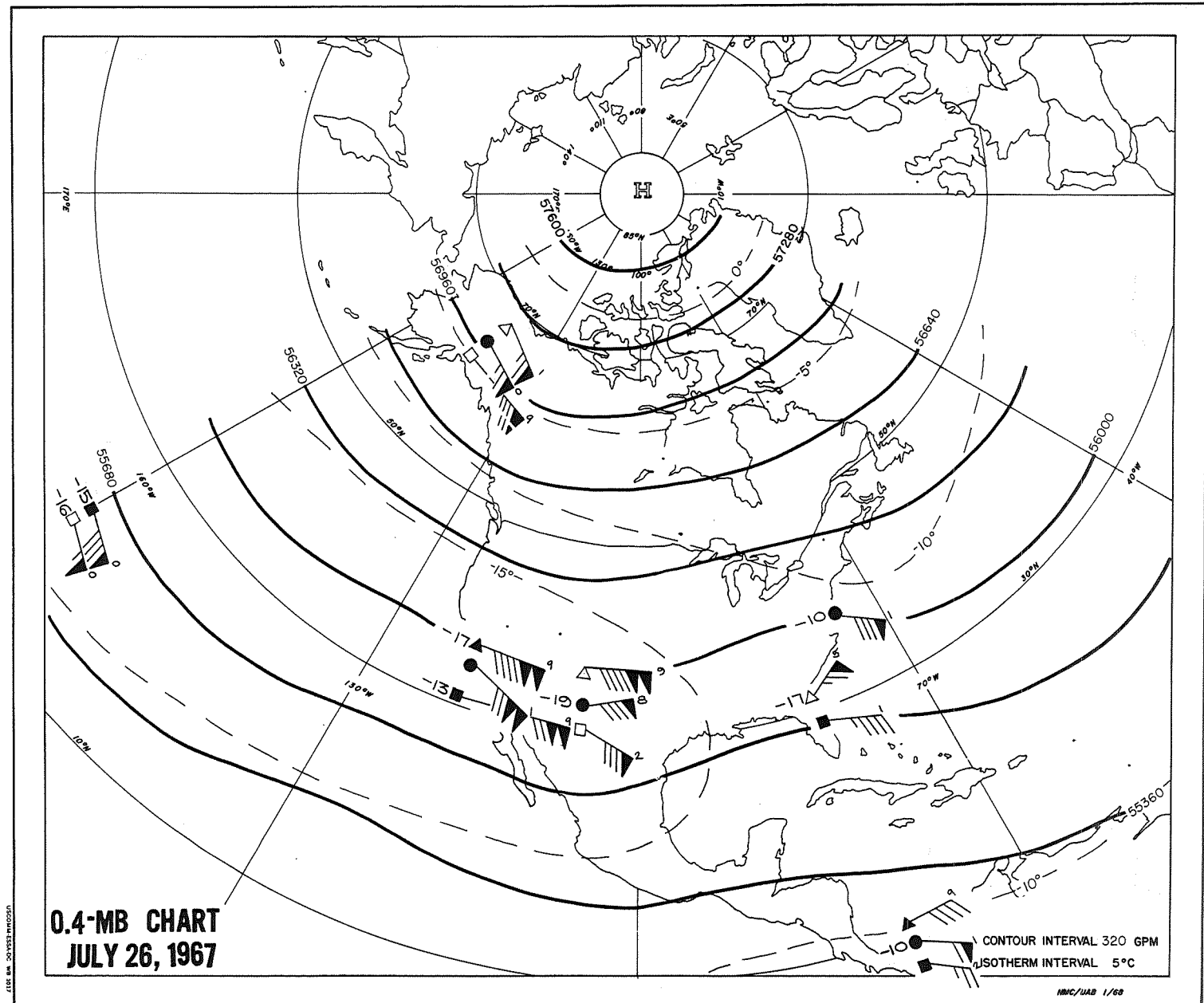


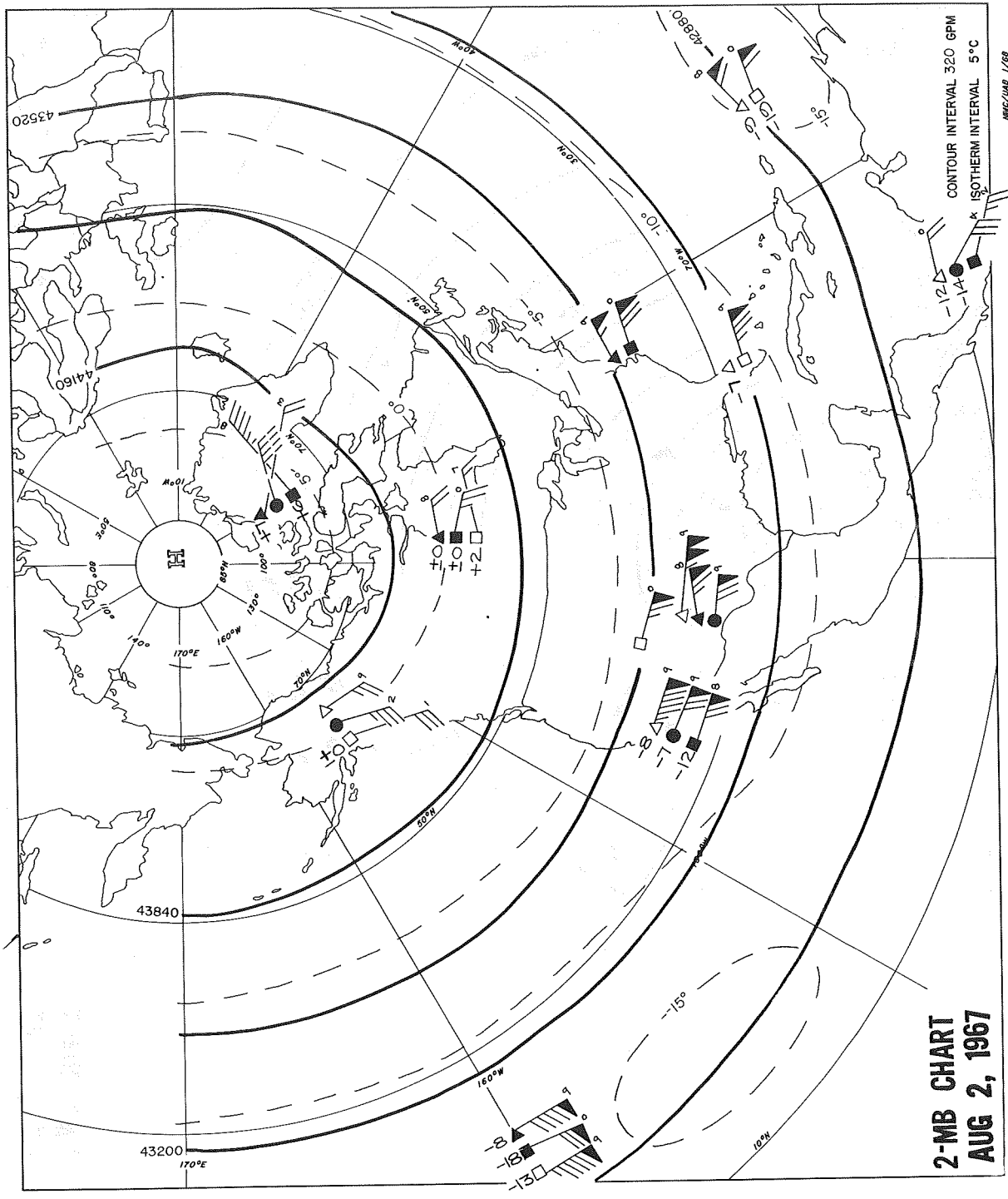




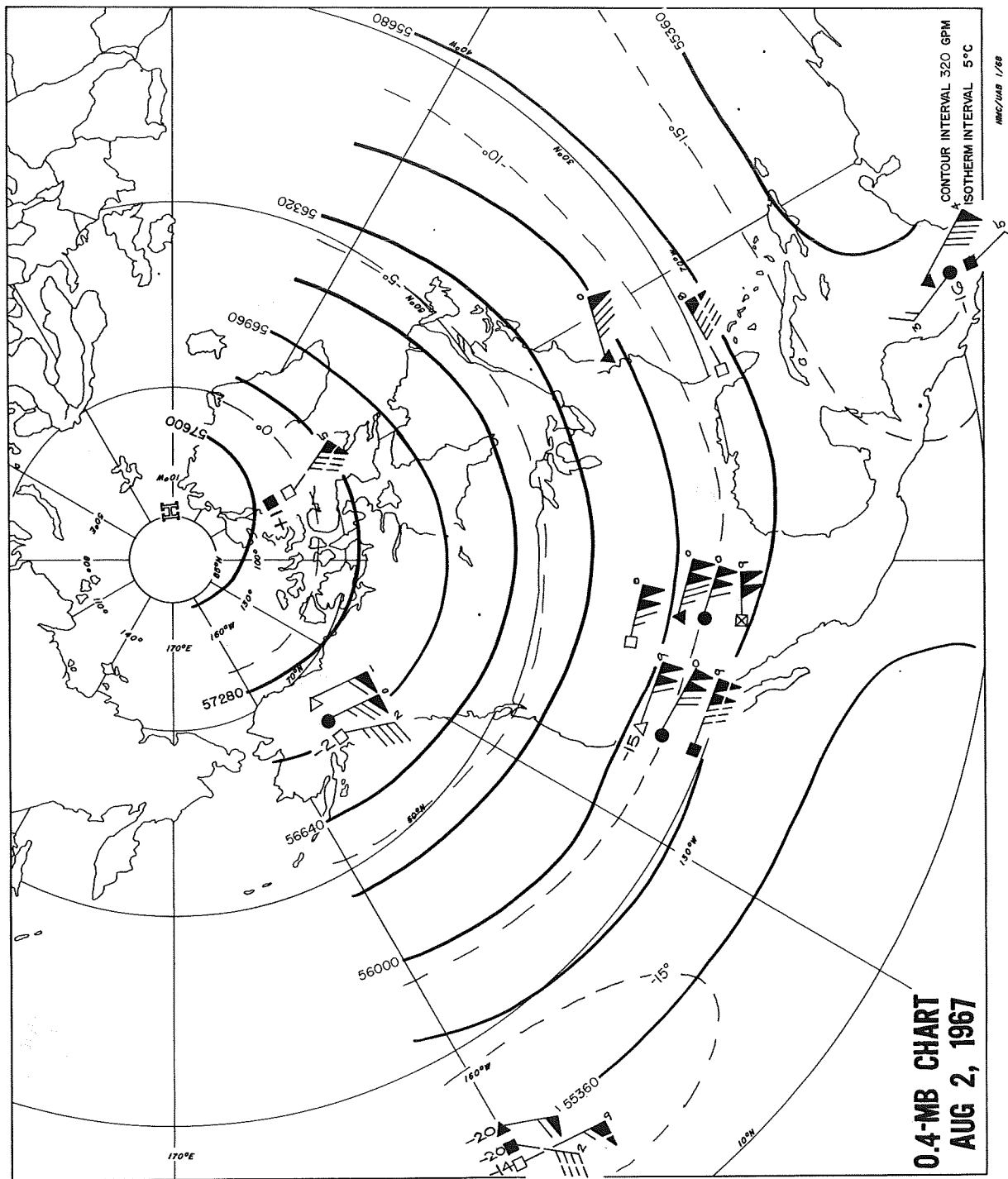






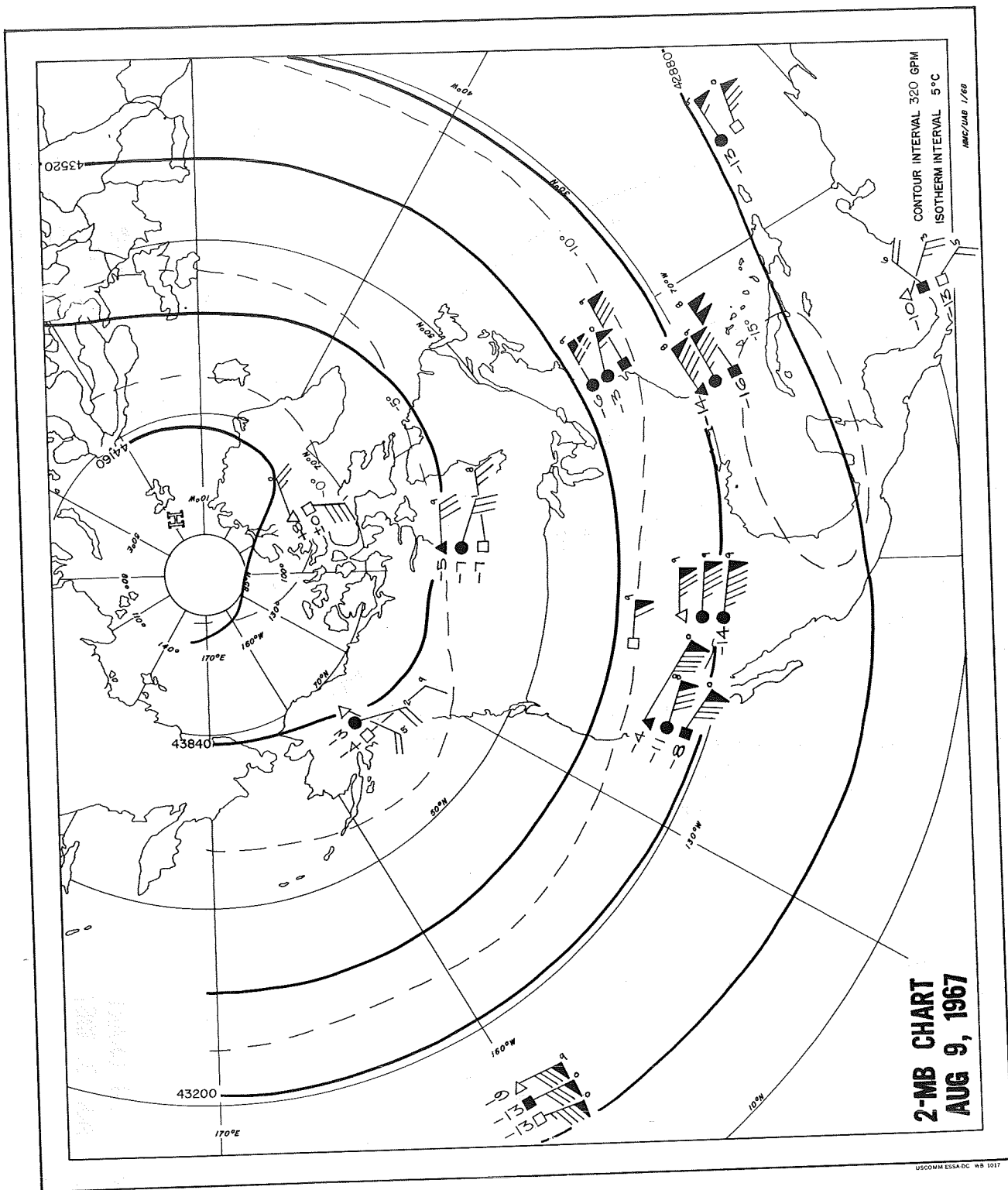


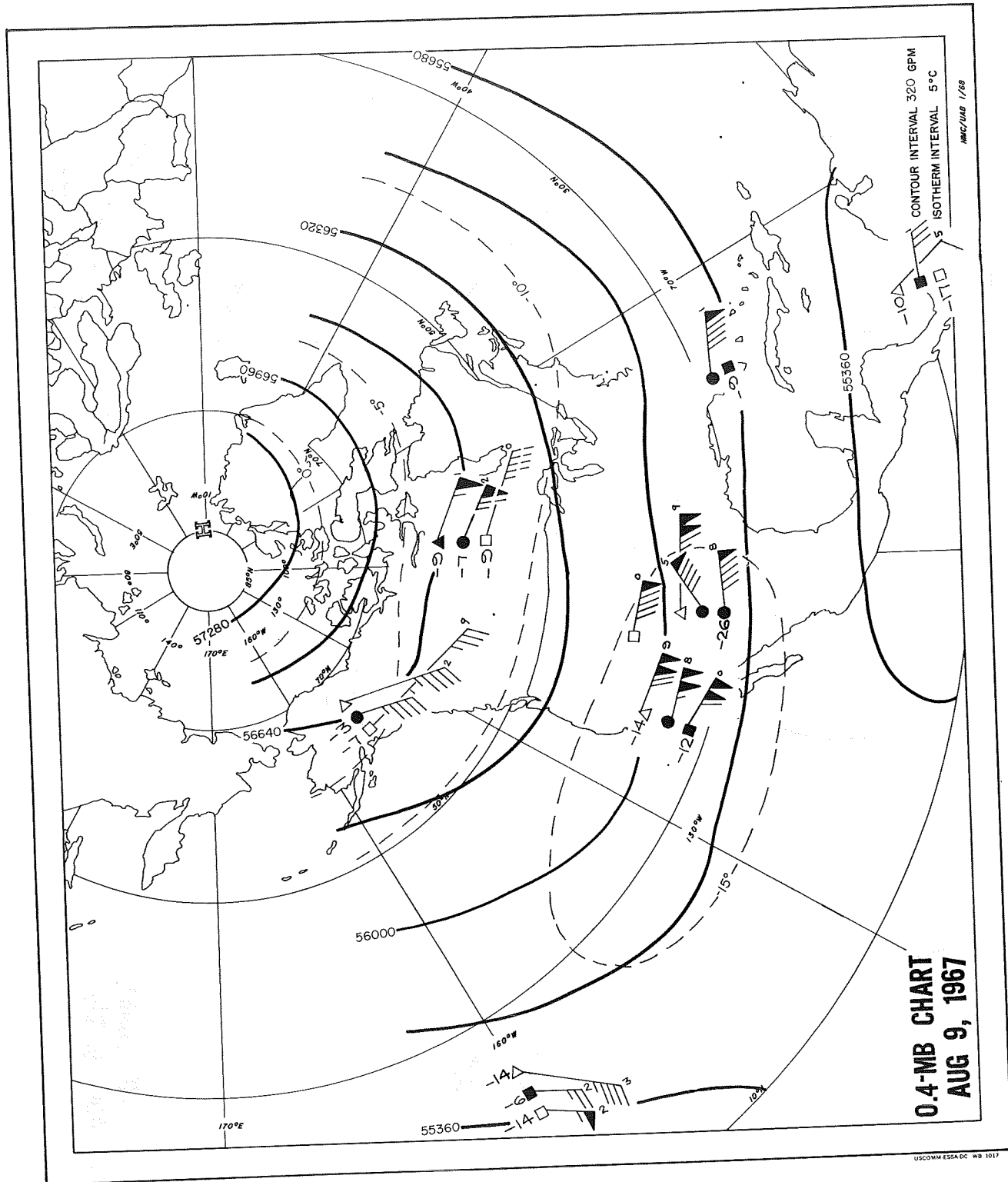
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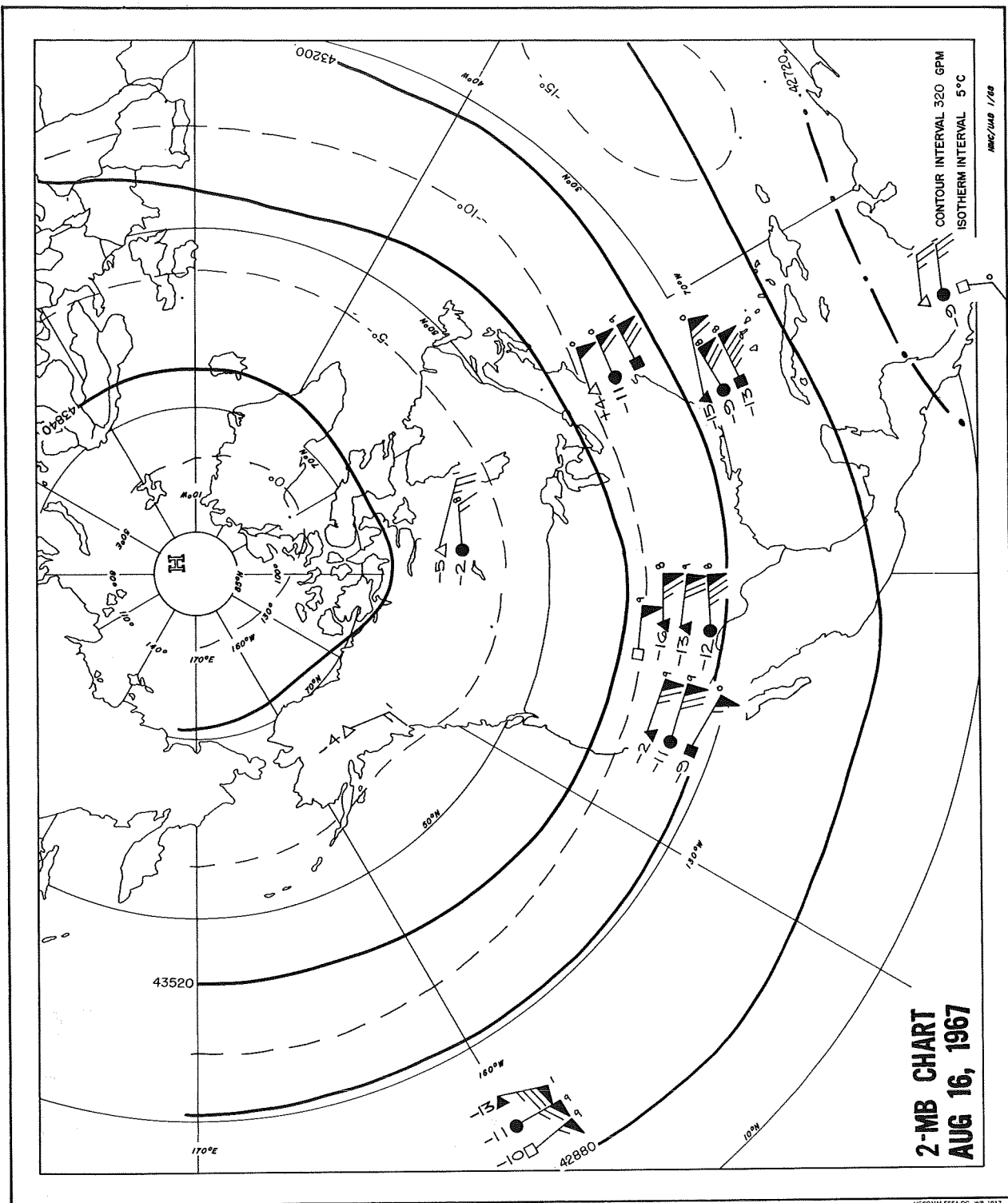


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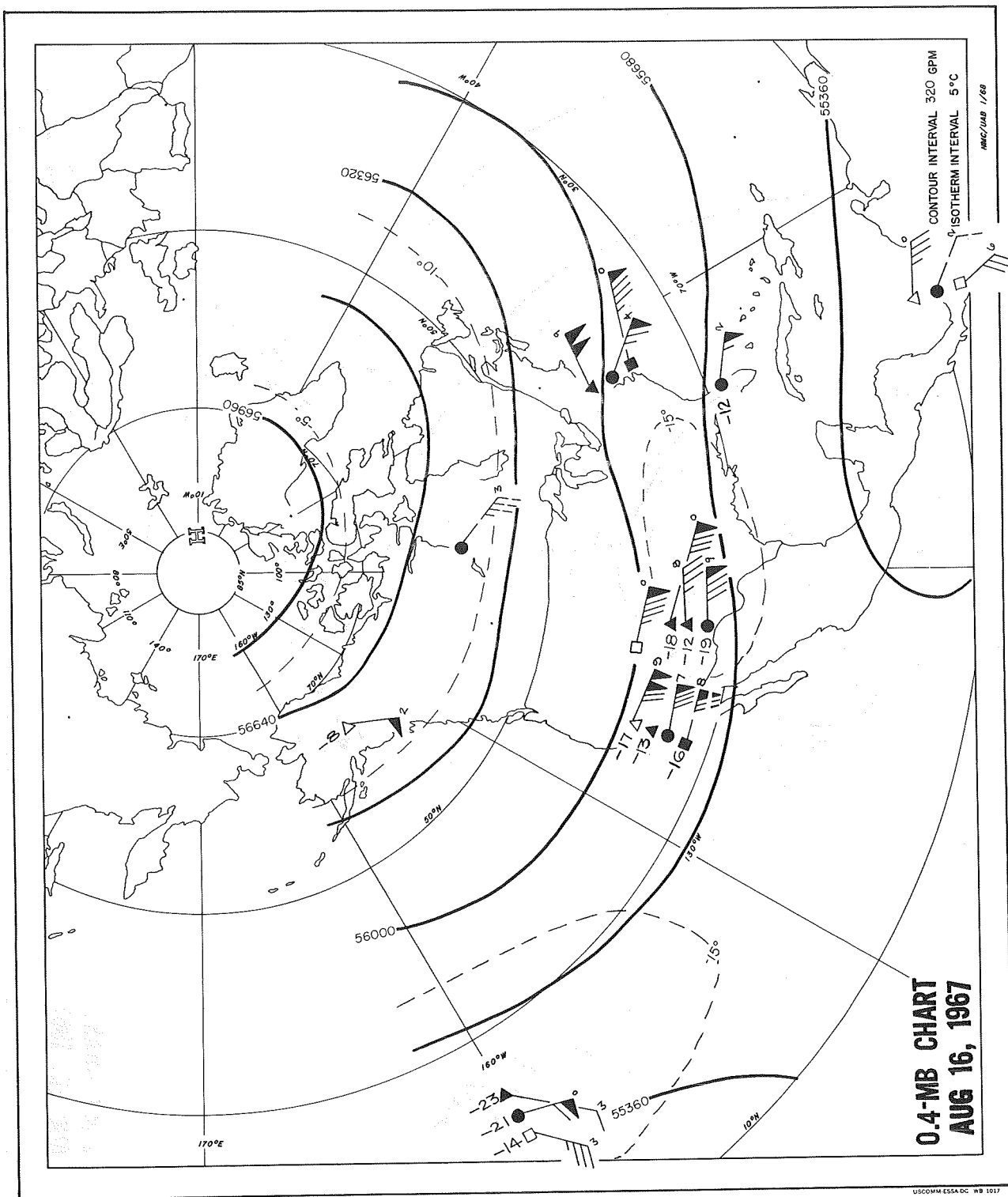


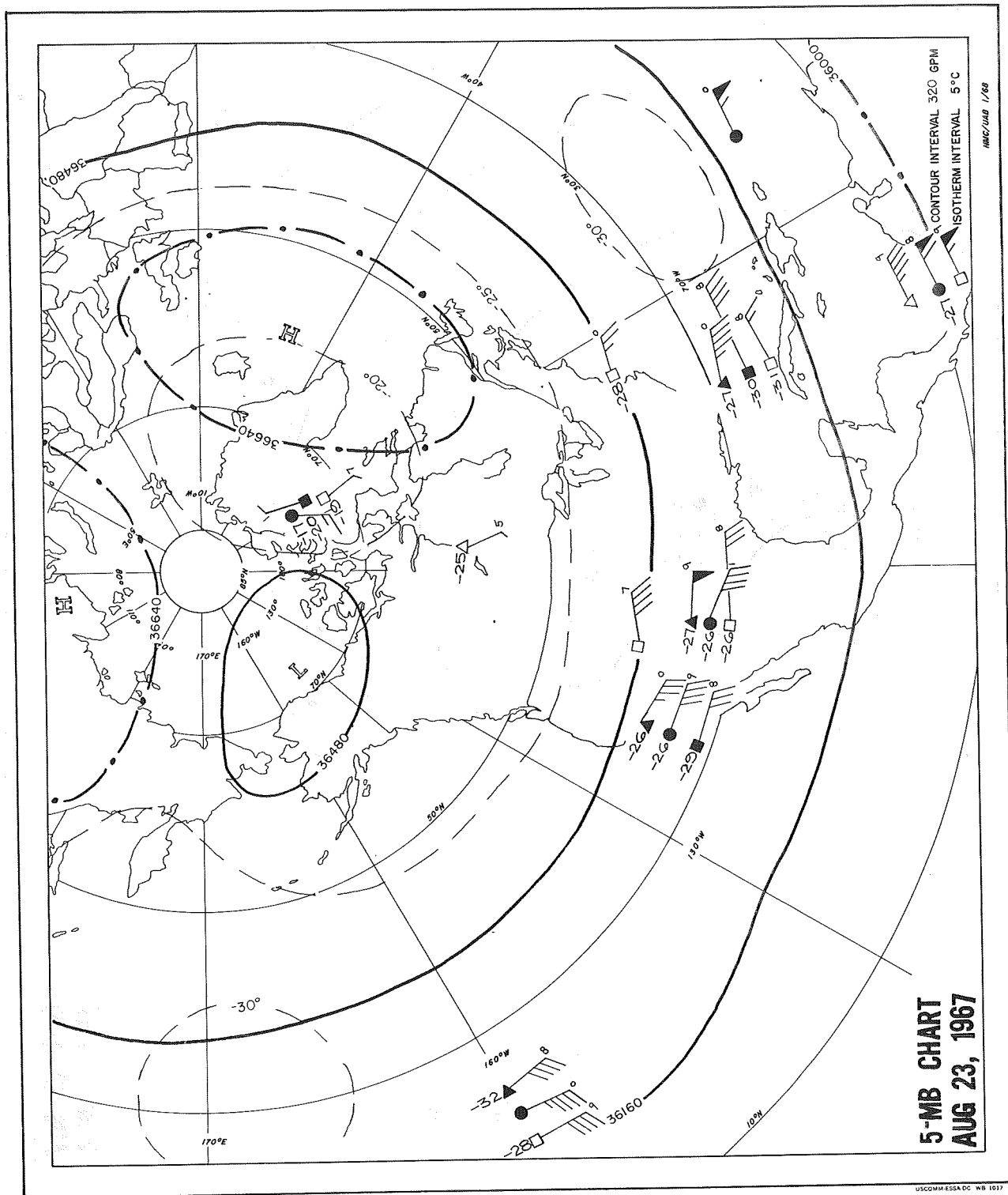


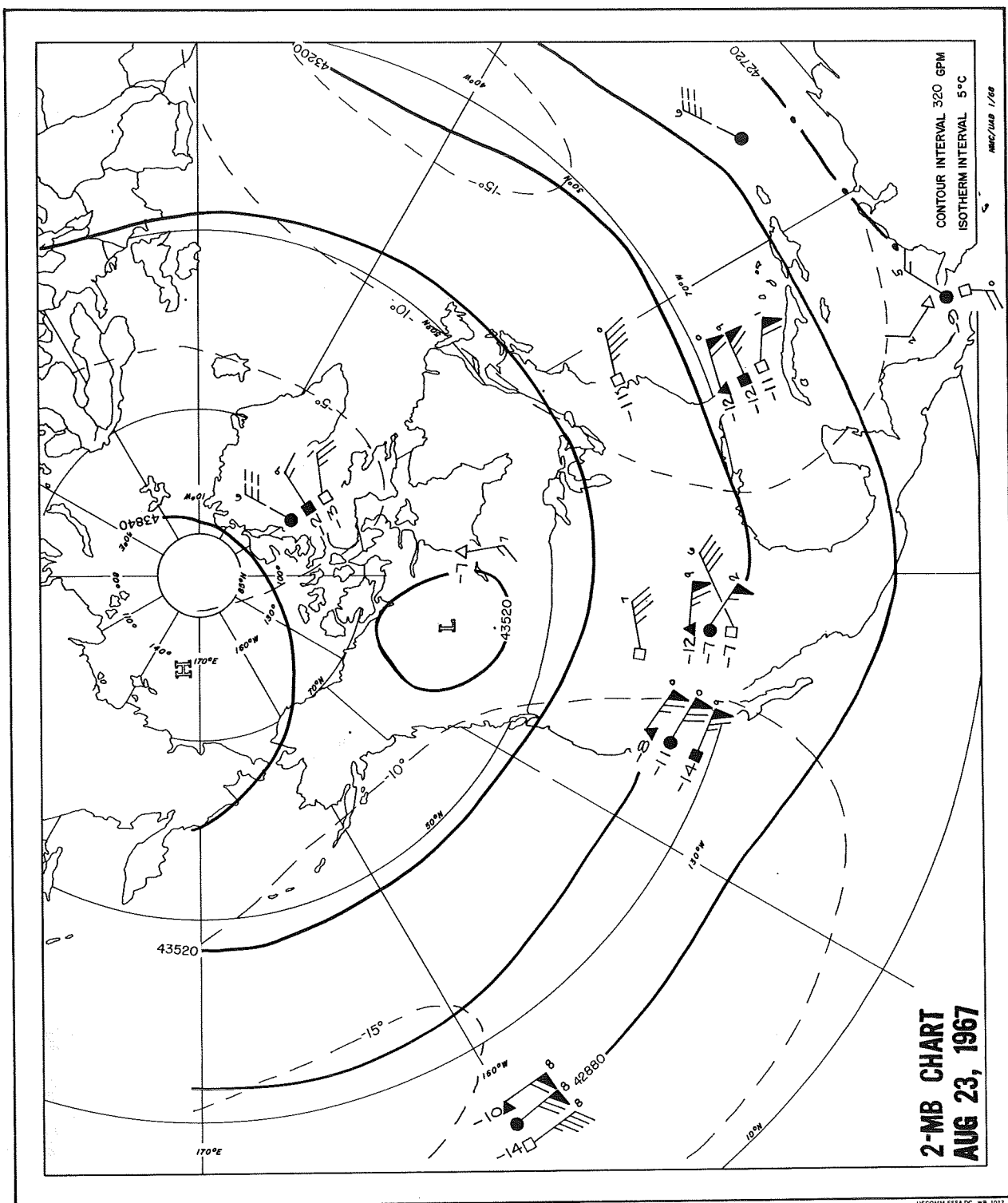


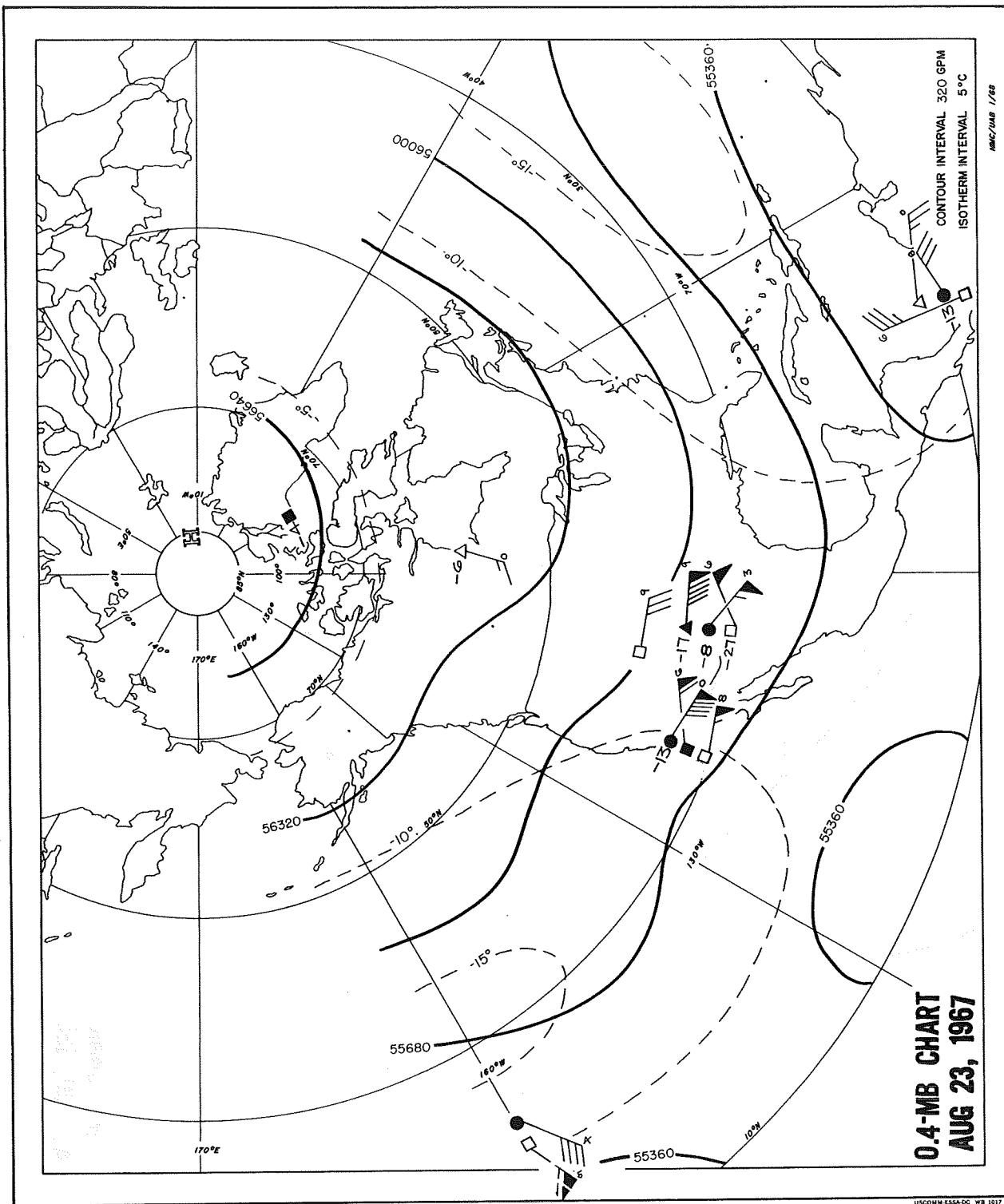


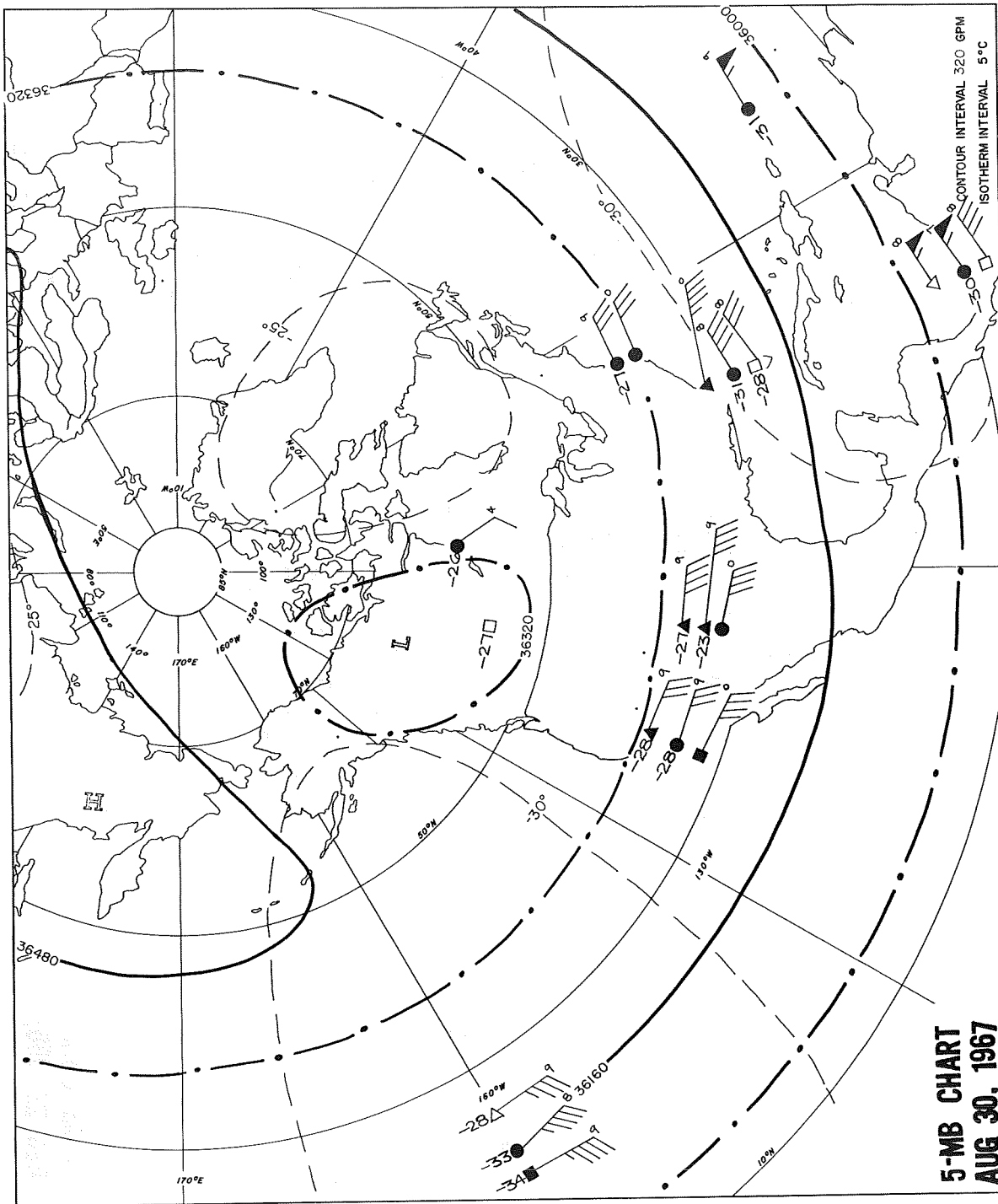
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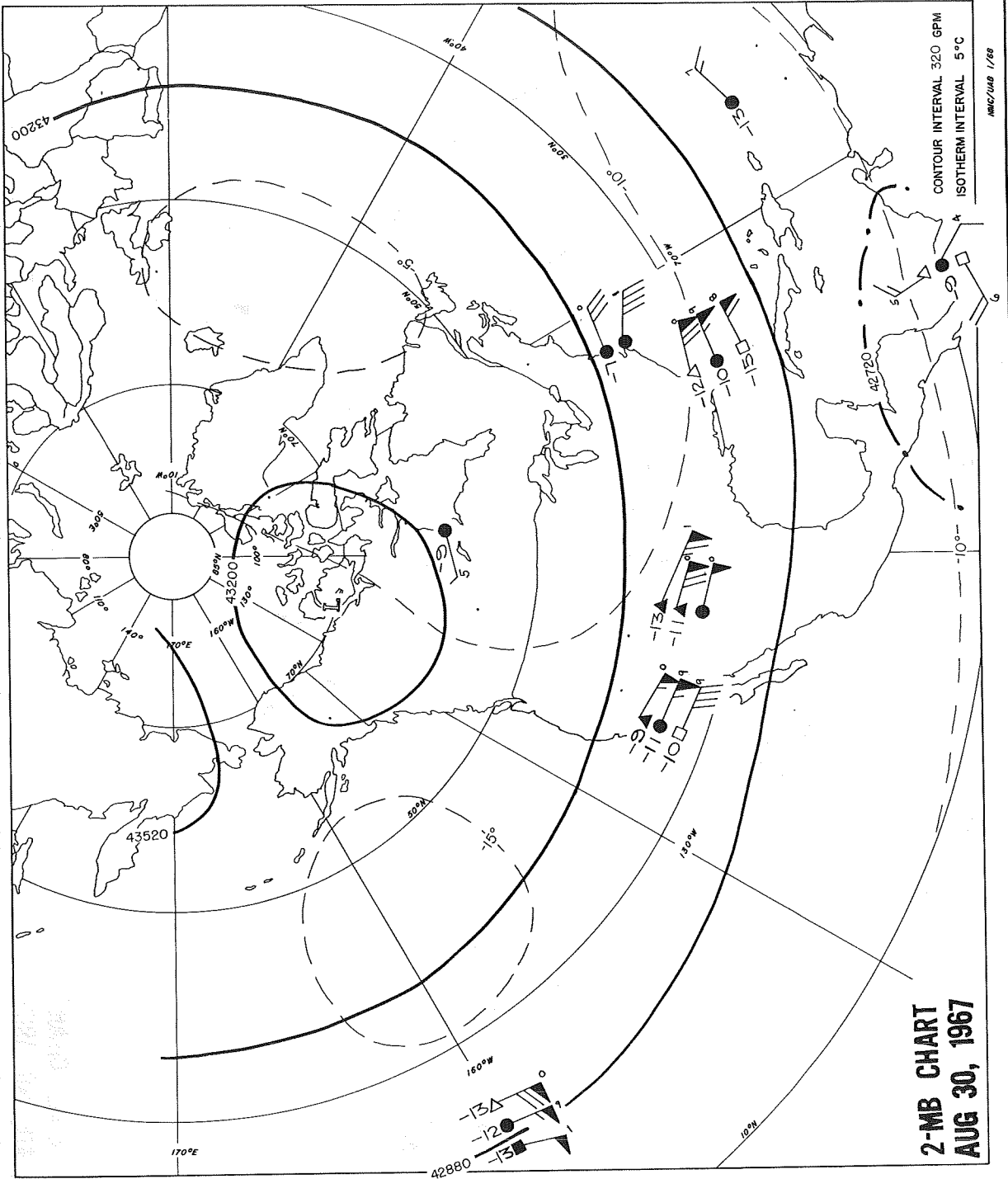




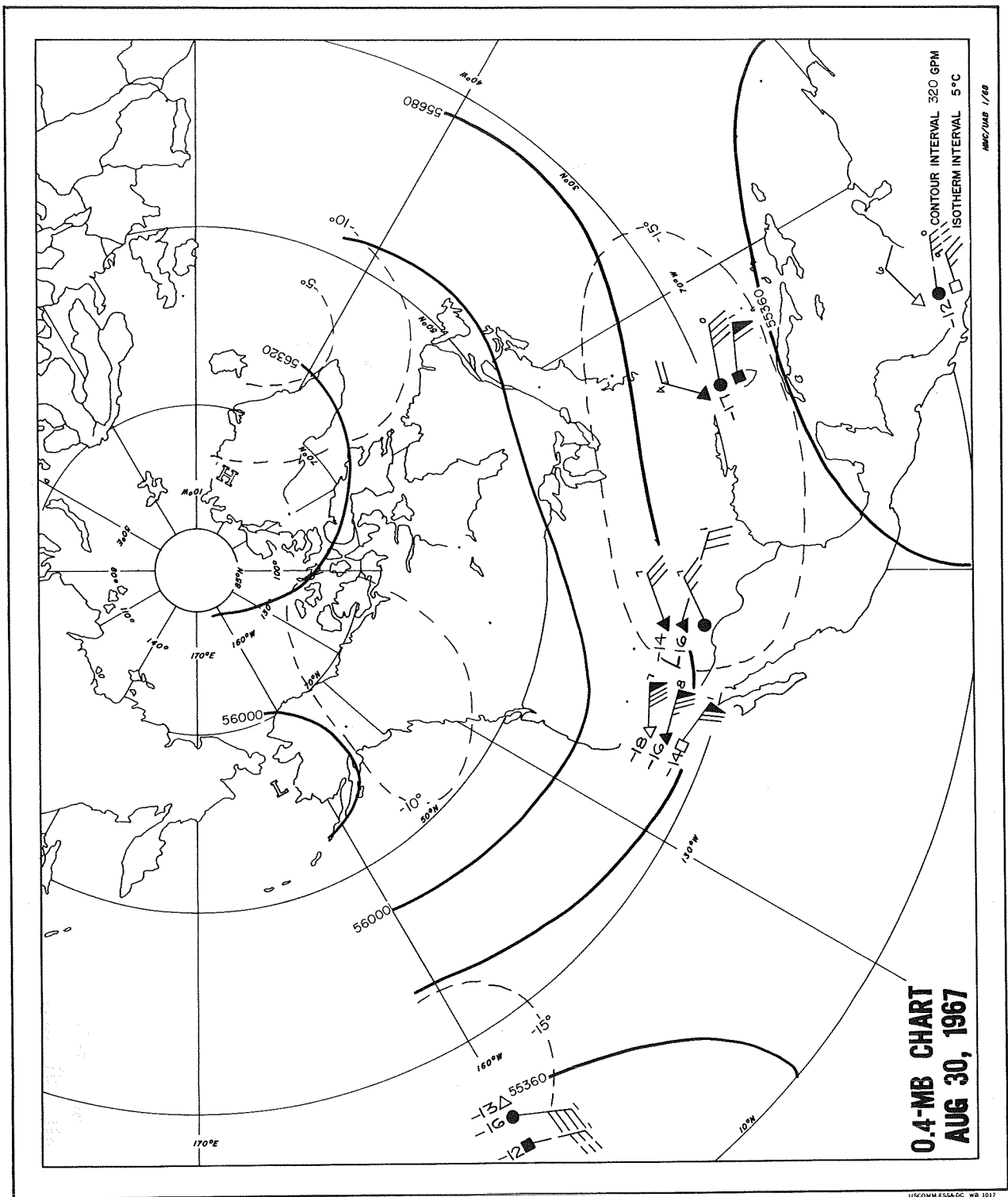
5-MB CHART
AUG 30, 1967

USCOMM ESSA DC WB 1017

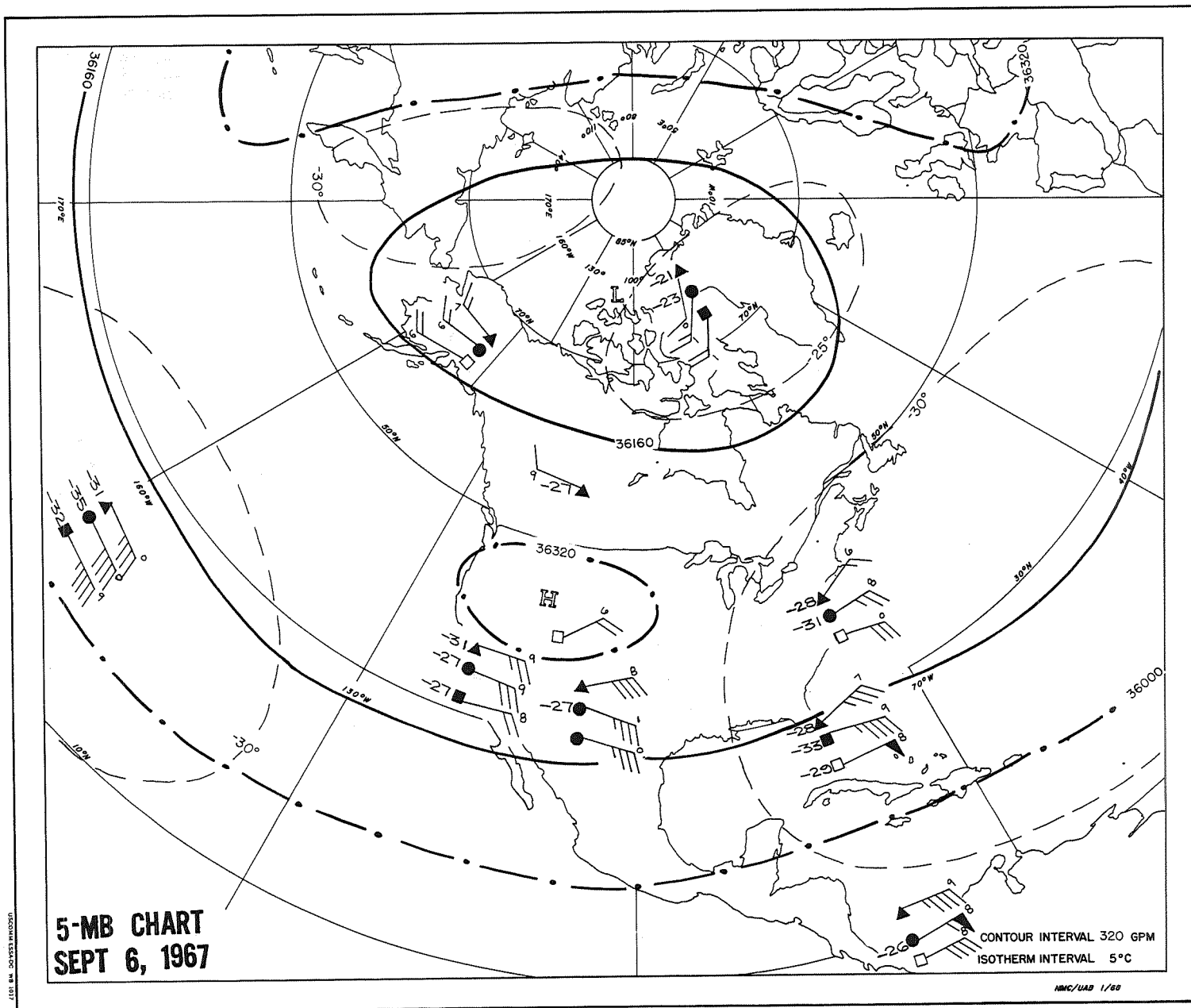
1/60

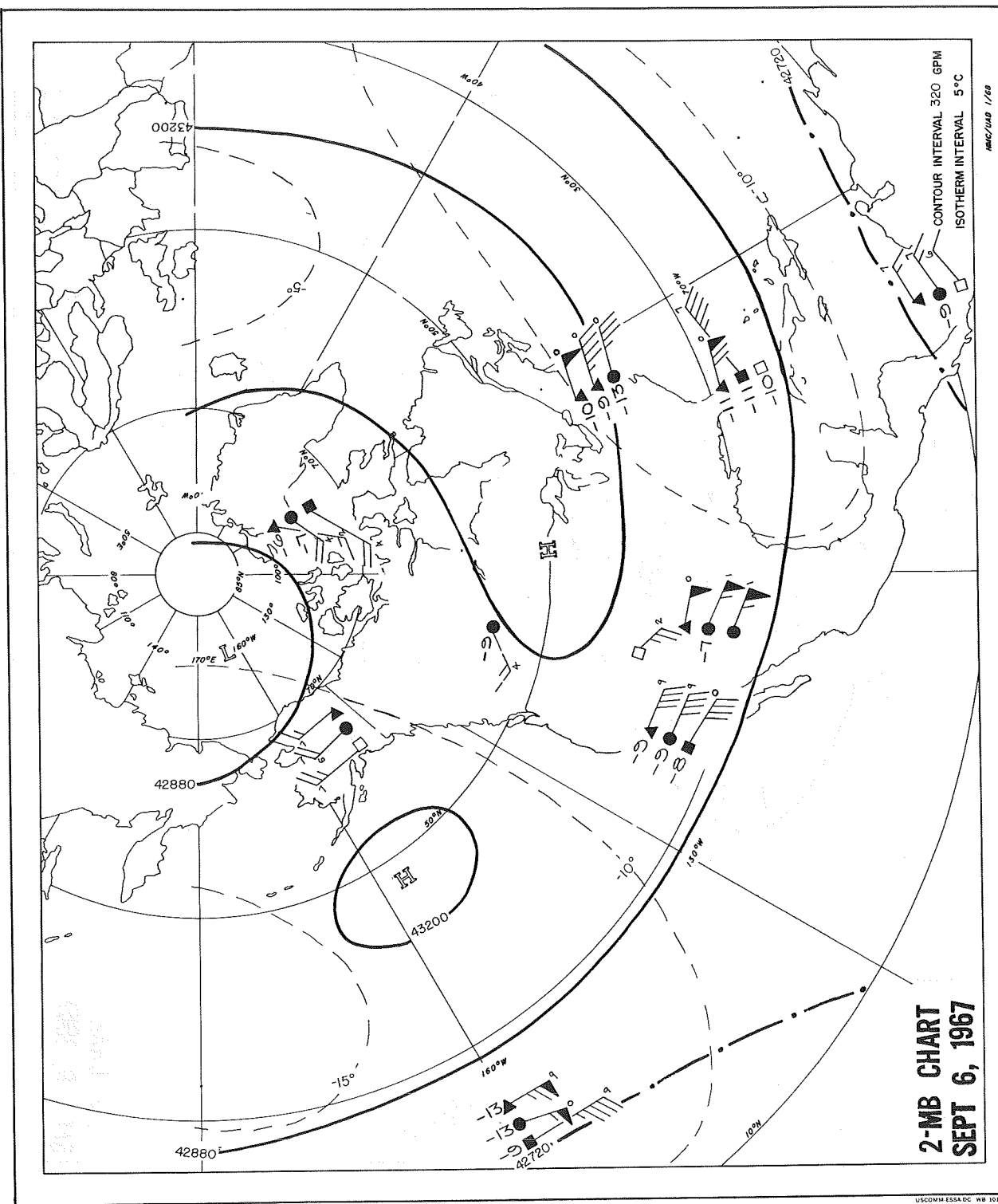


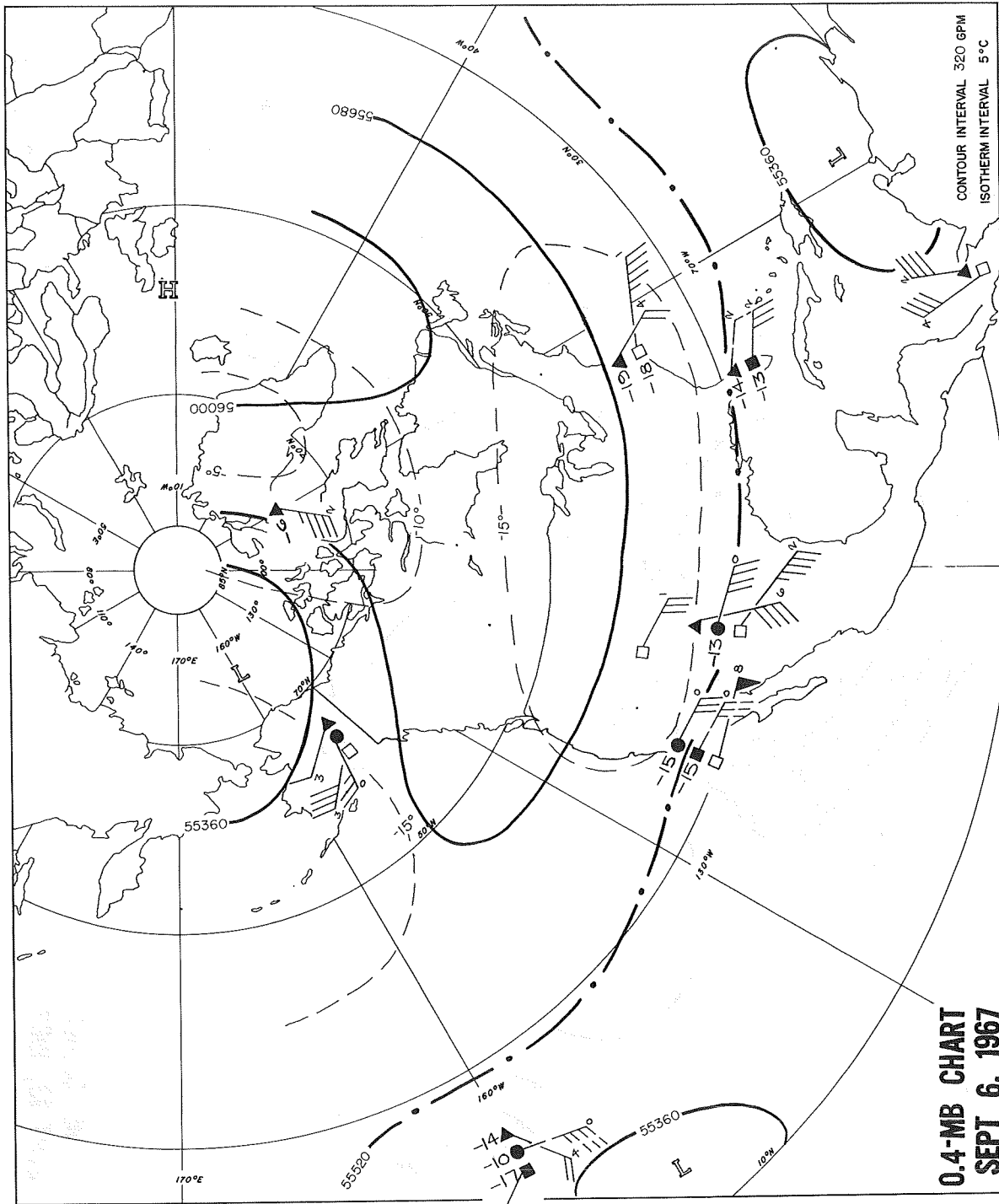
USCOMH ESSA DC WB 1017



USCOMM ESSA-DC WB 1017



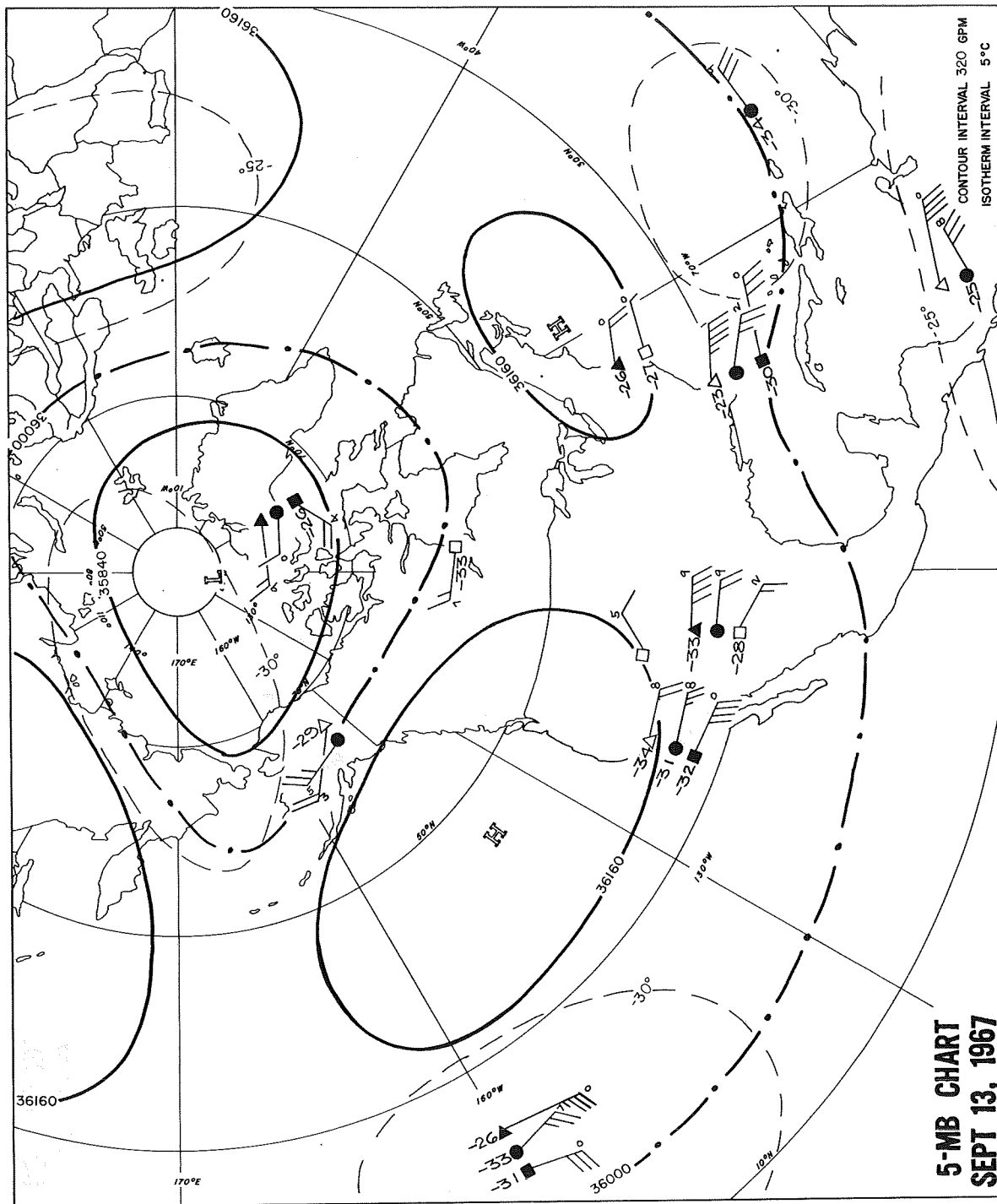




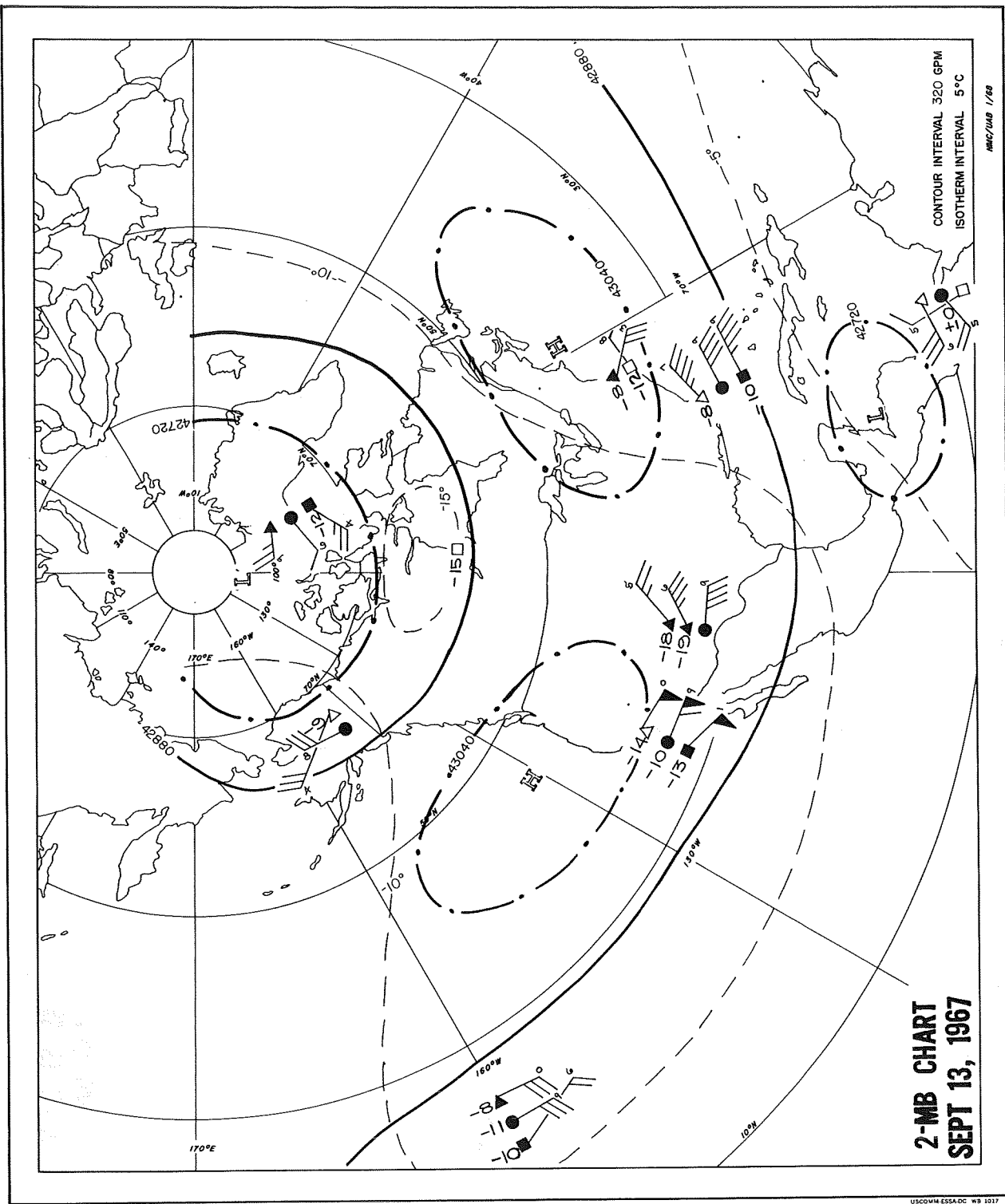
NMC/US 1/68

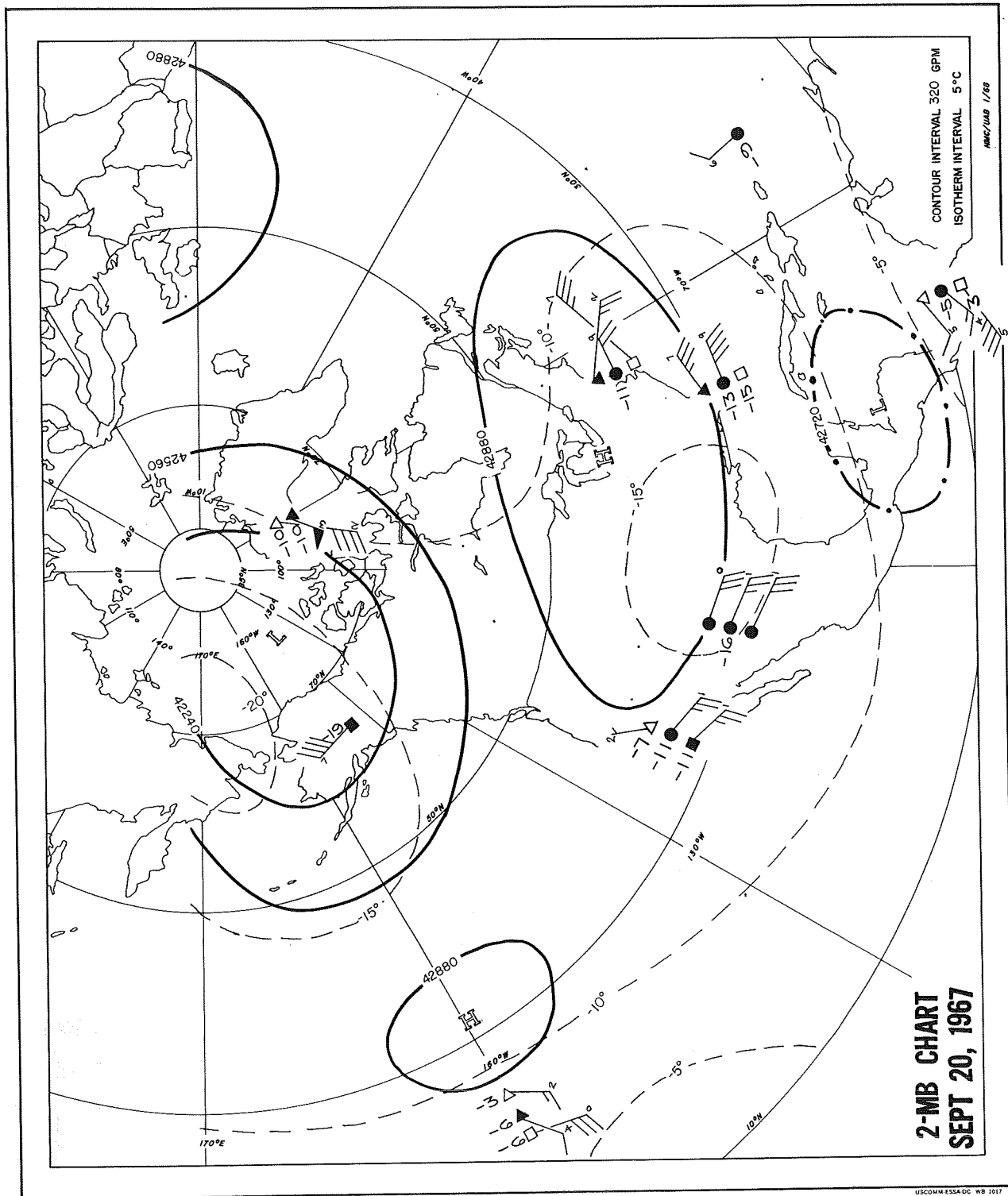
0.4-MB CHART
SEPT 6, 1967

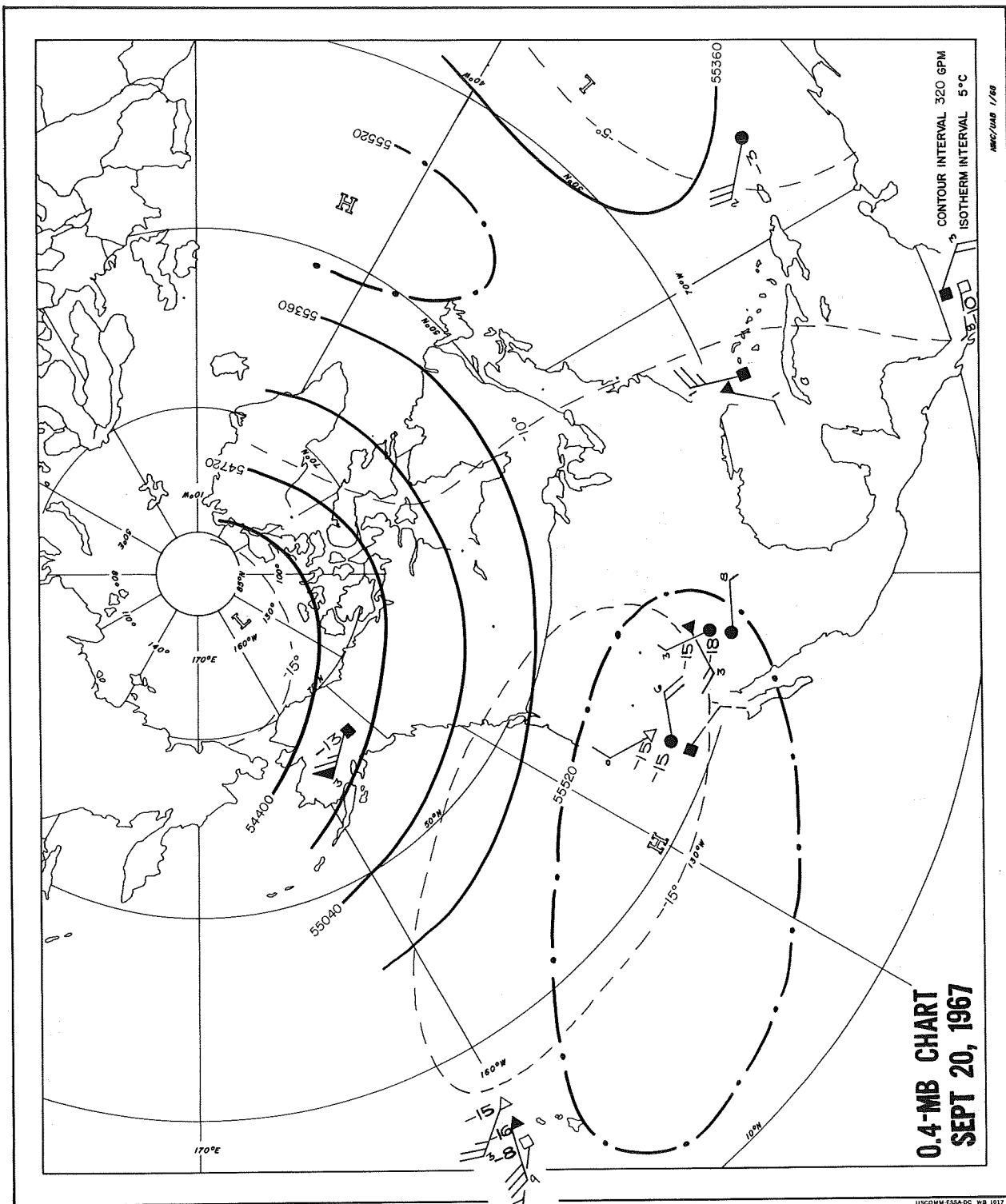
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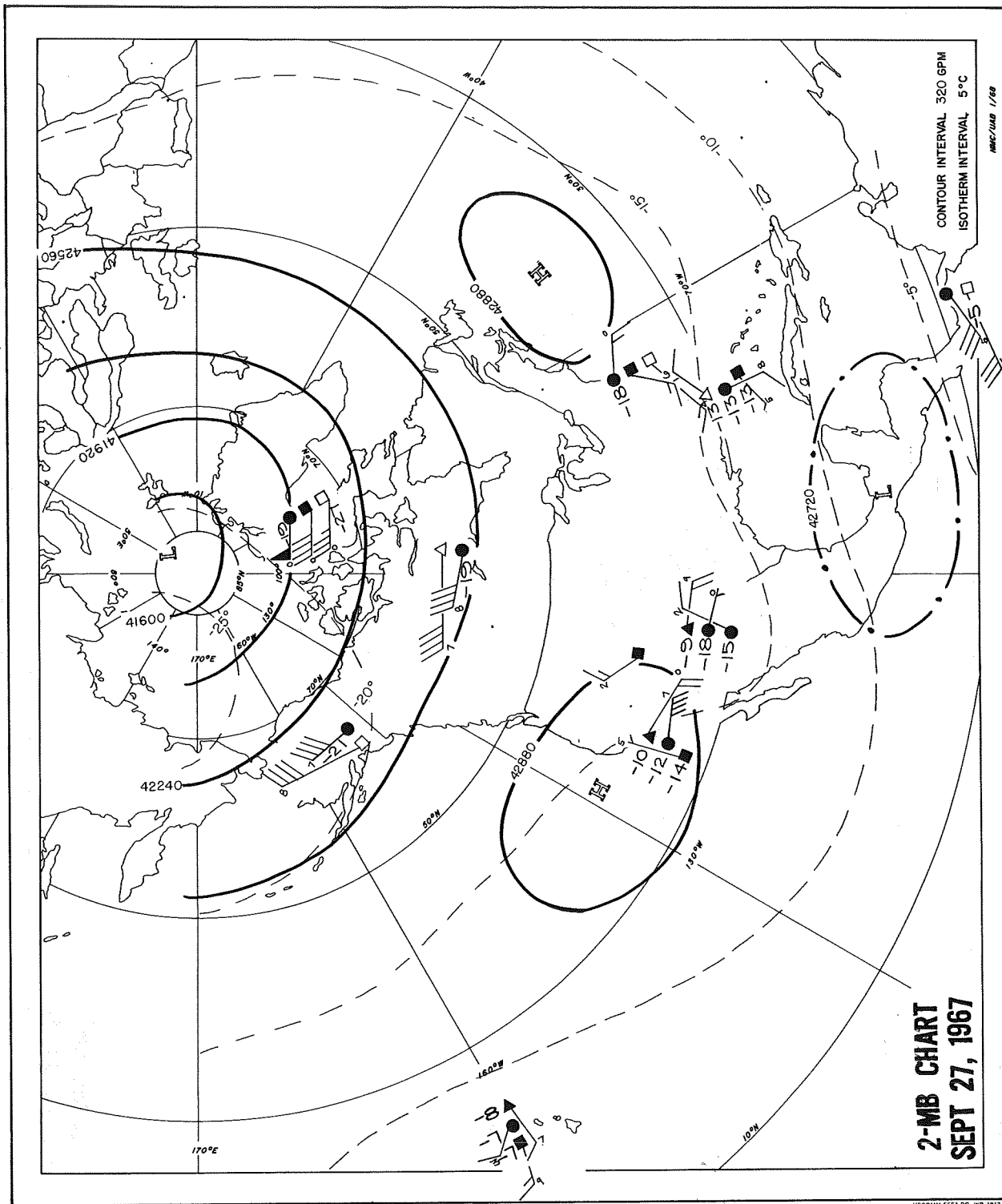


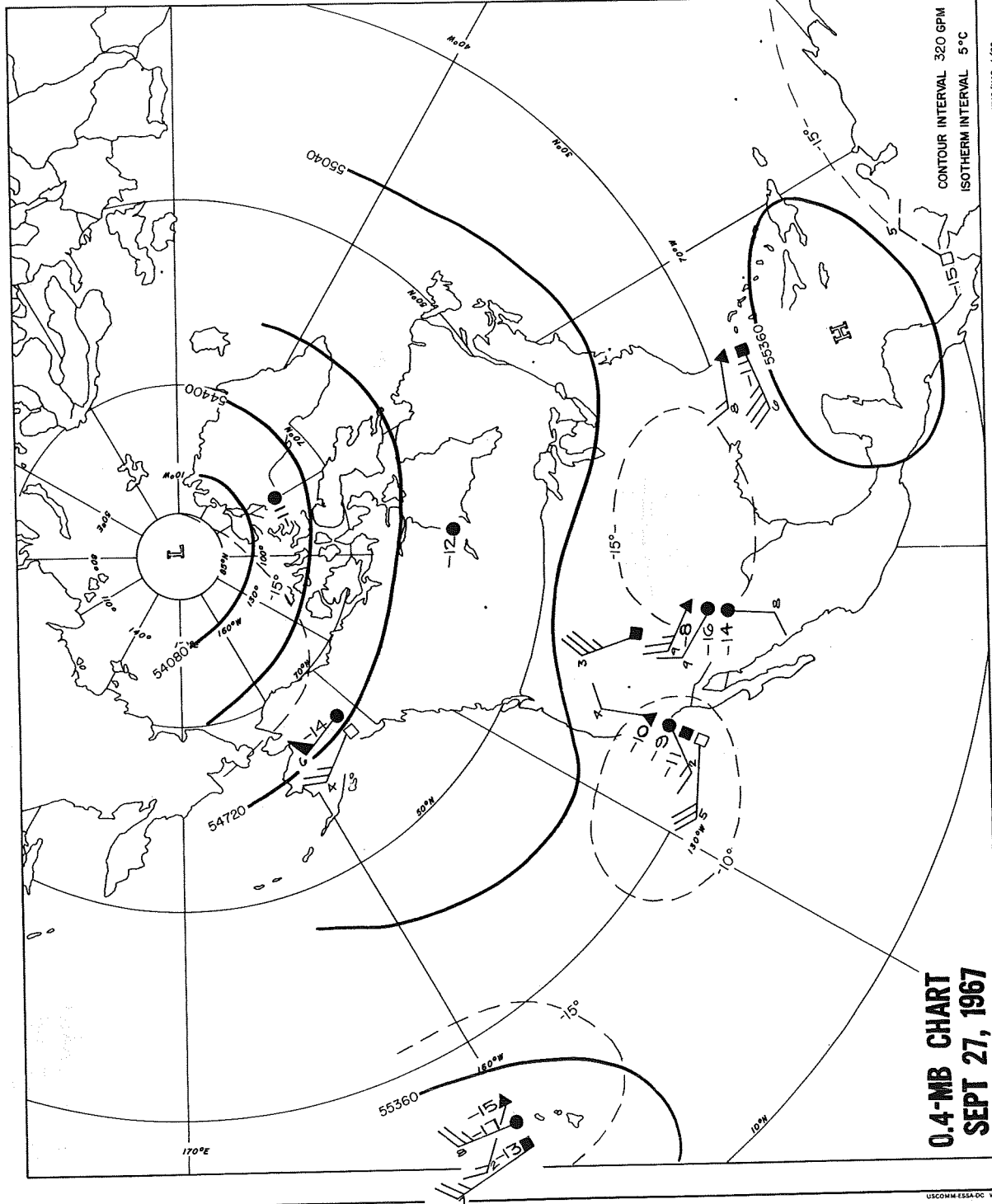
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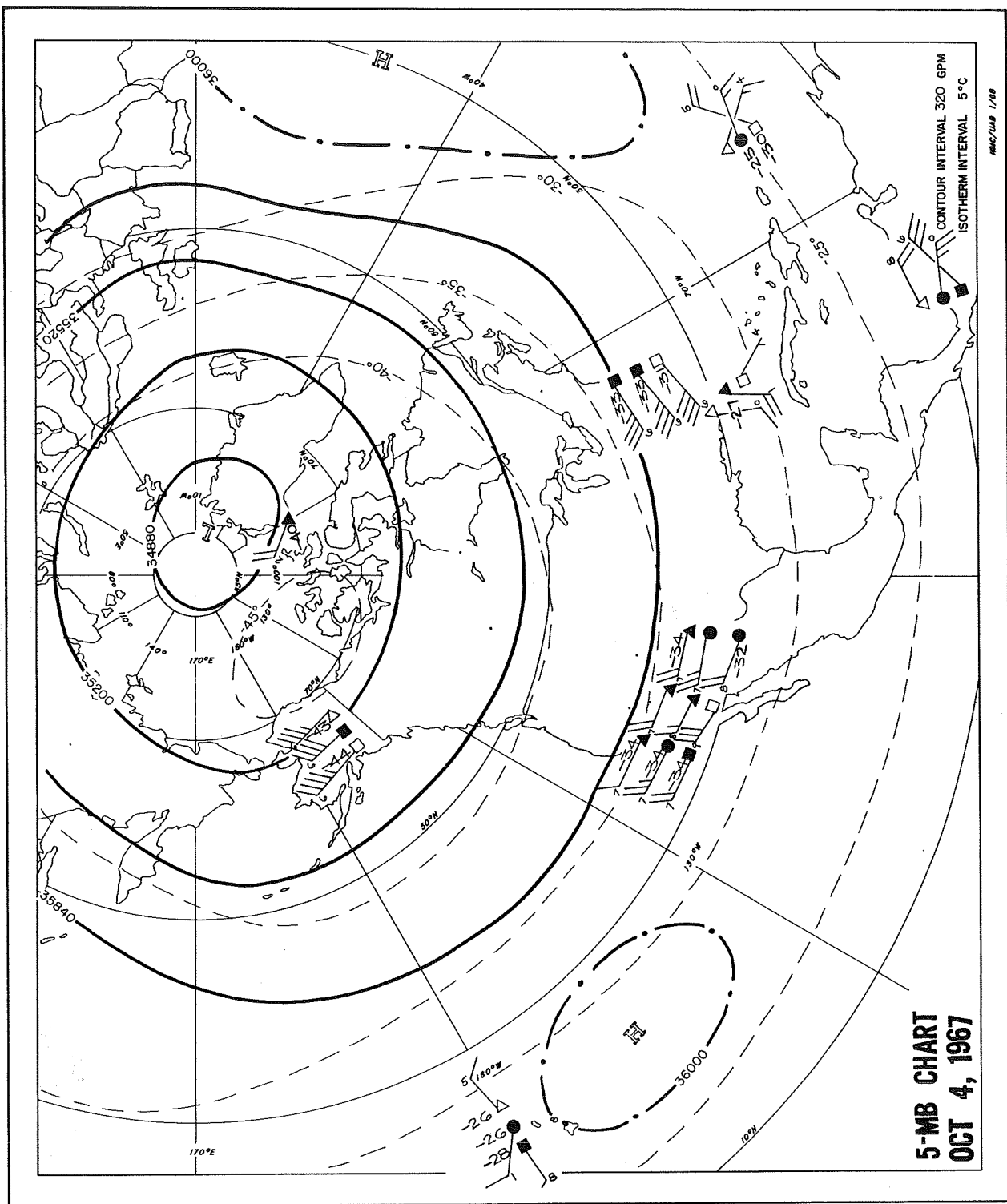


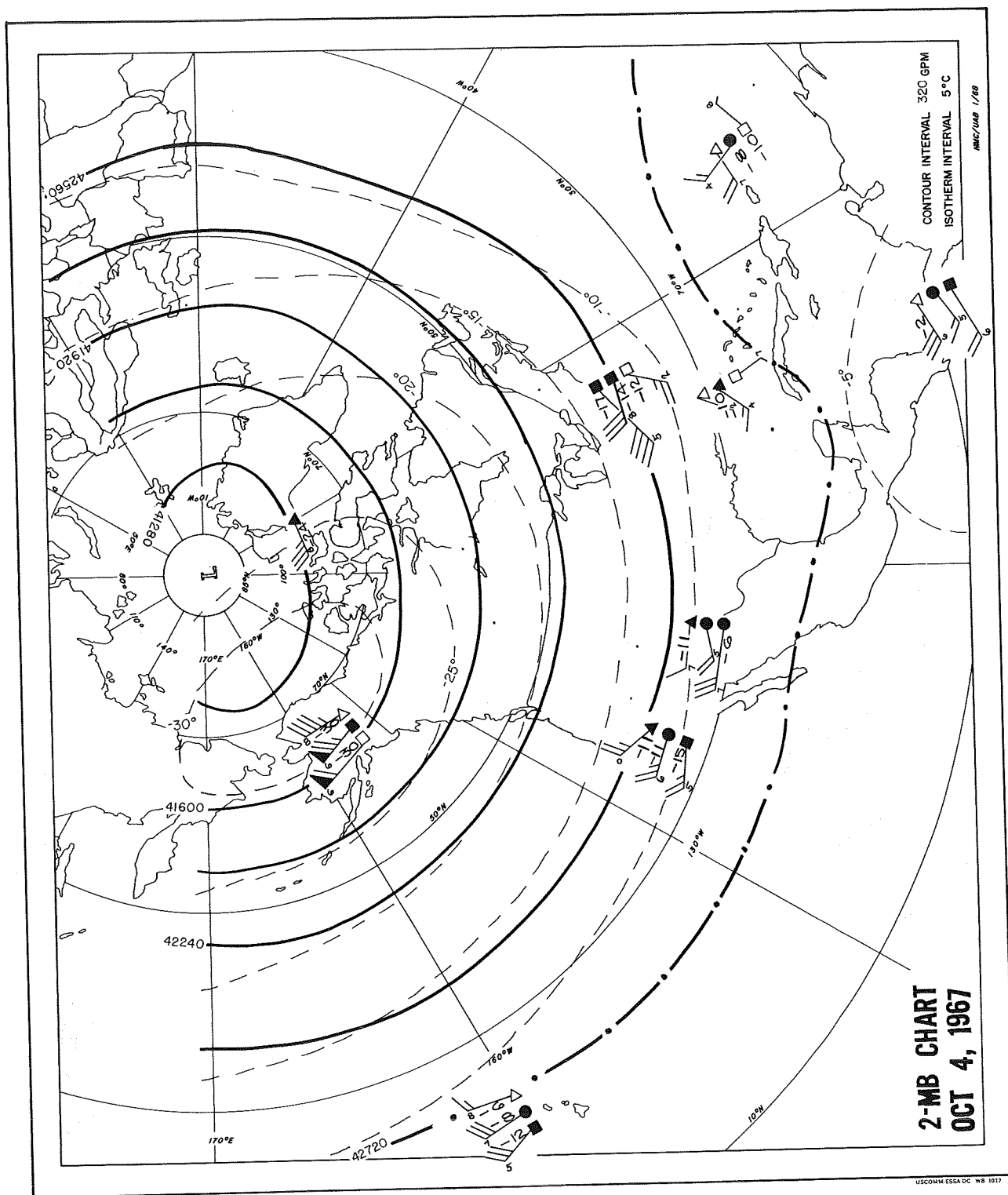


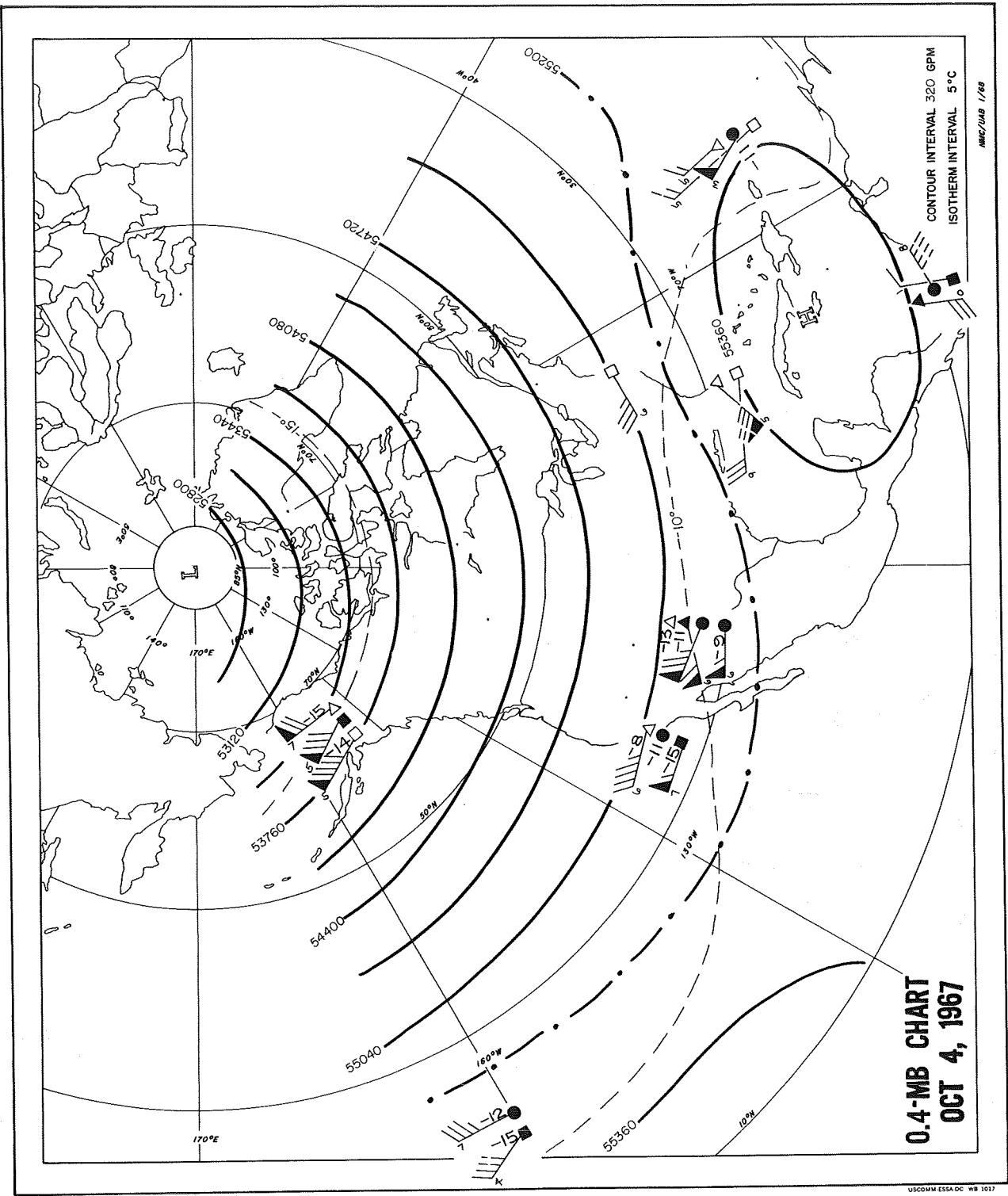




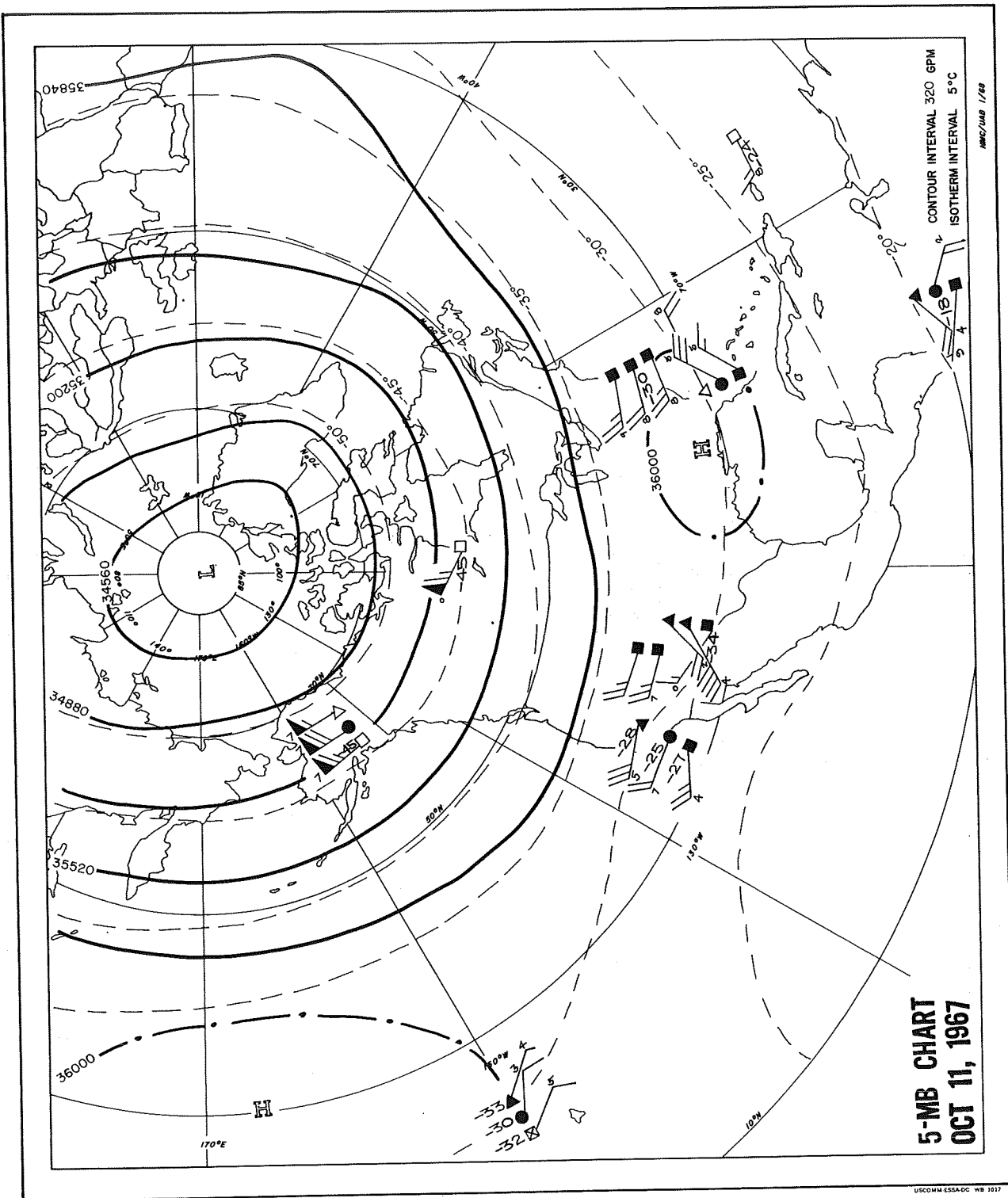
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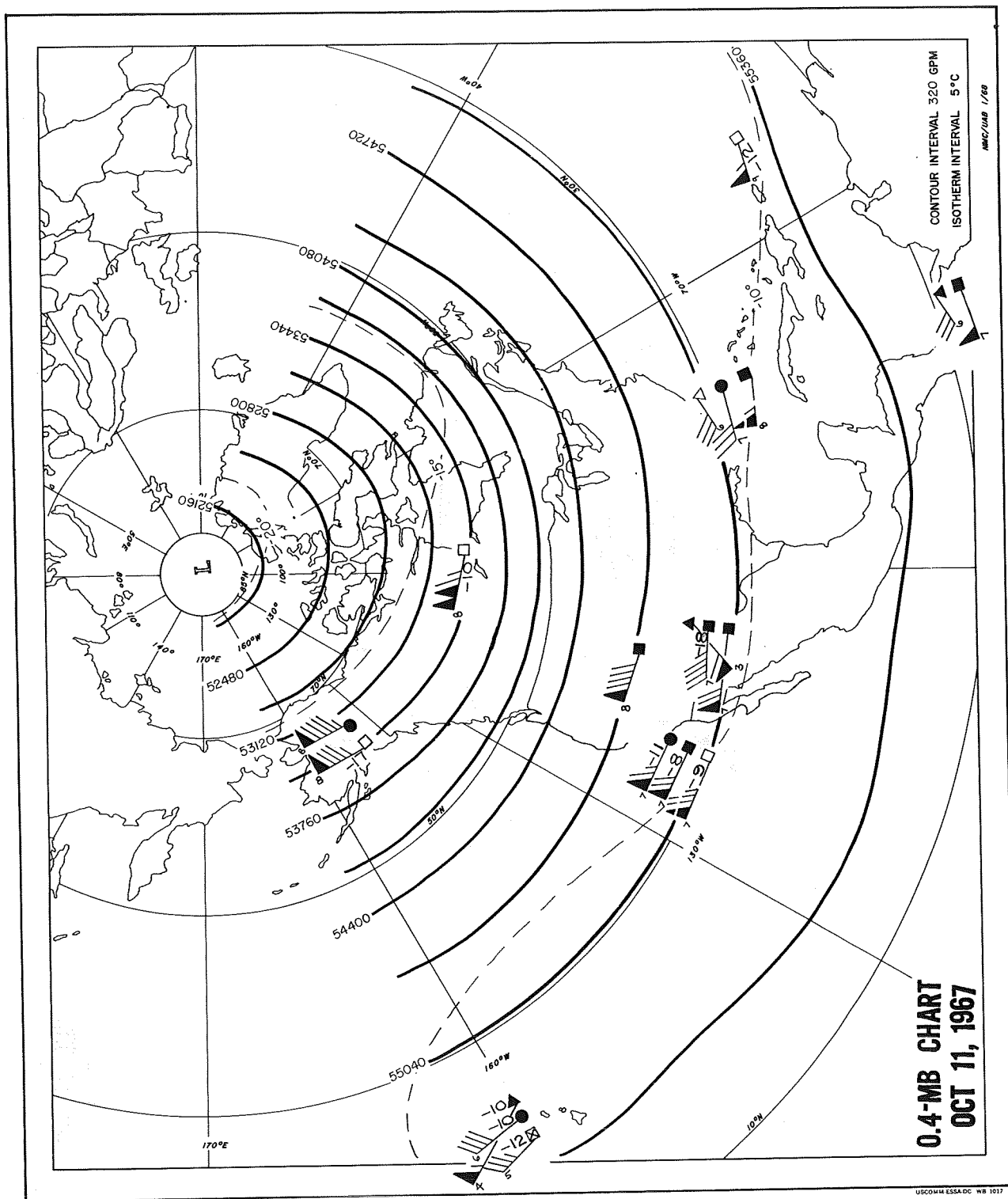


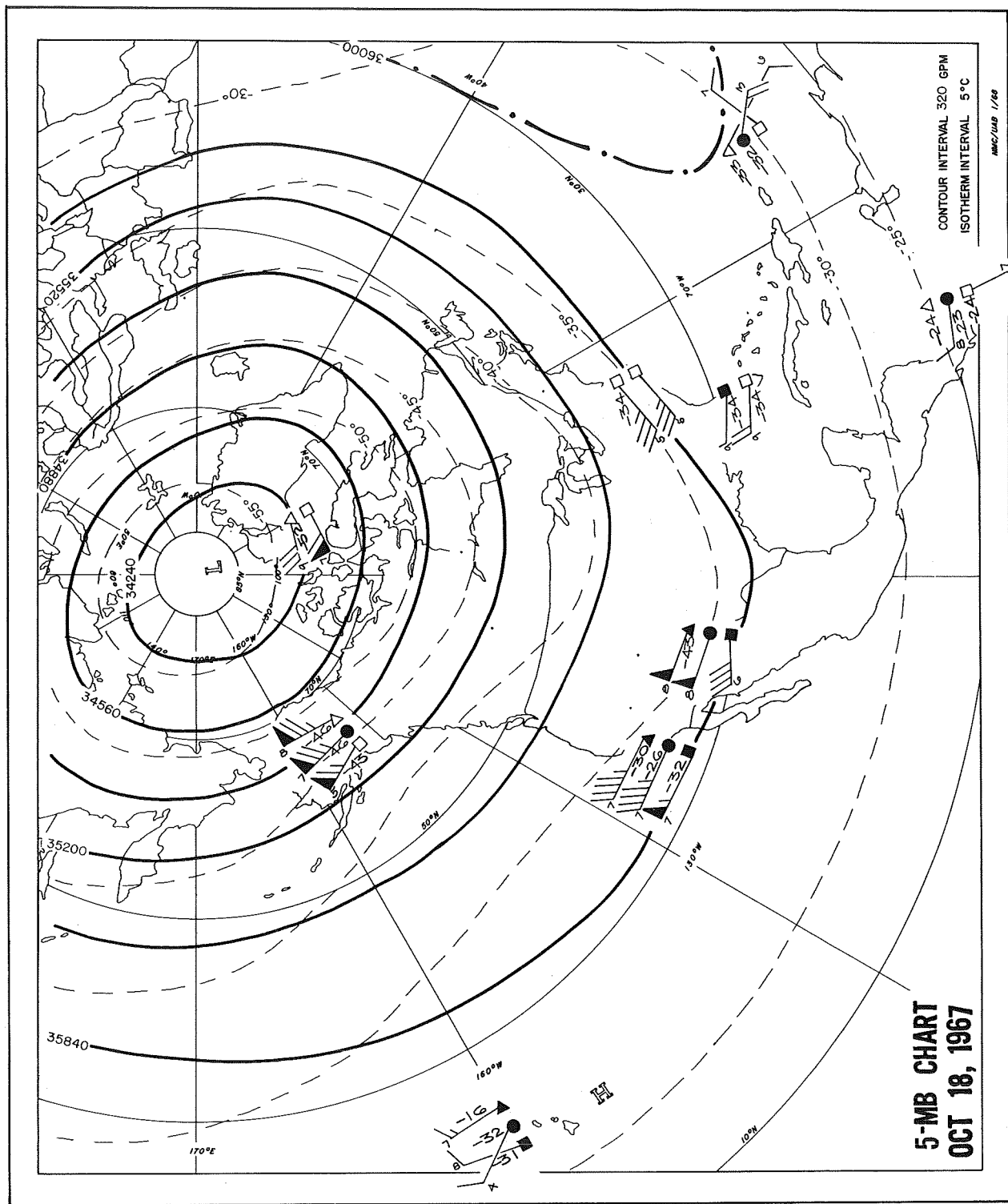




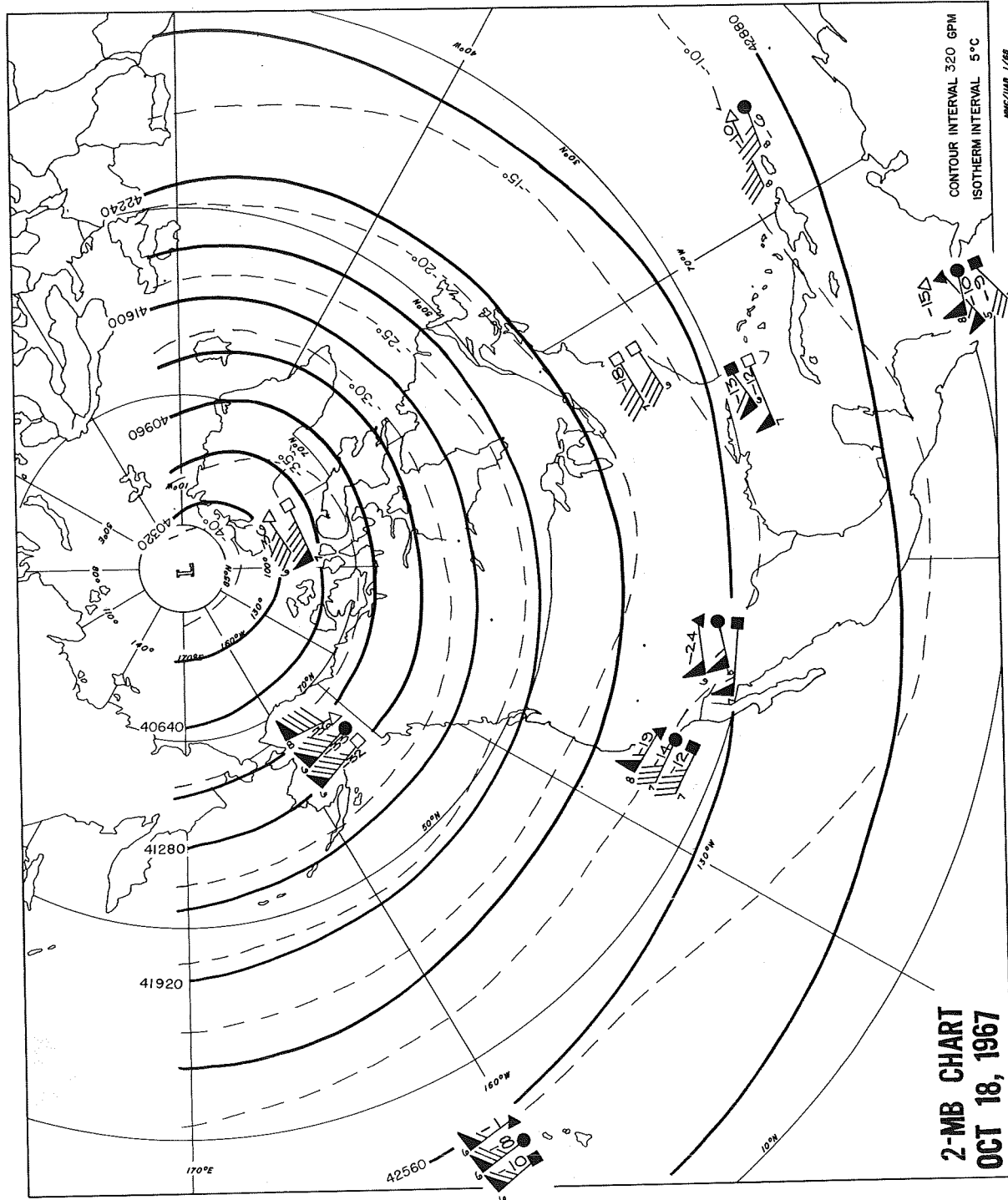
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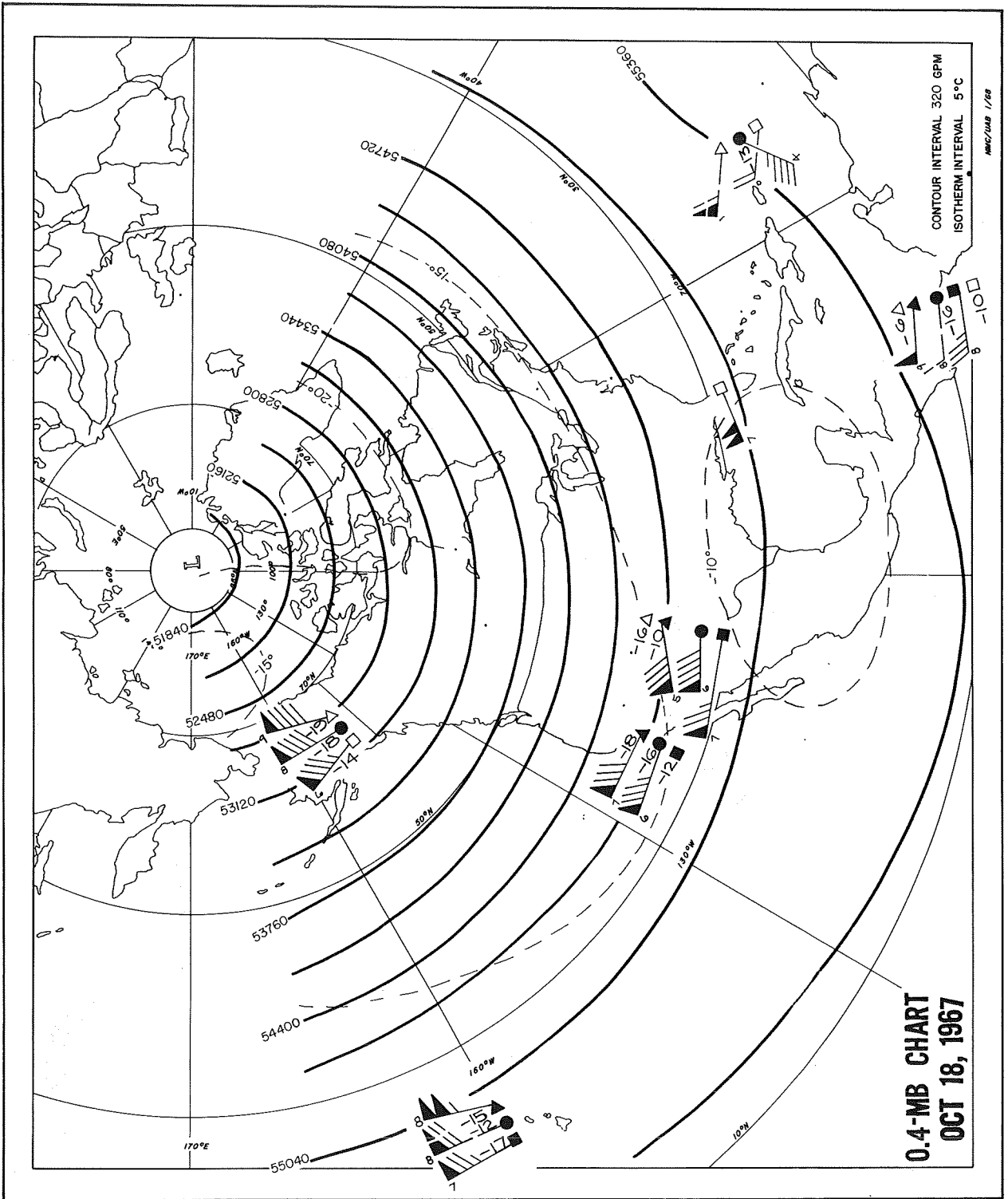




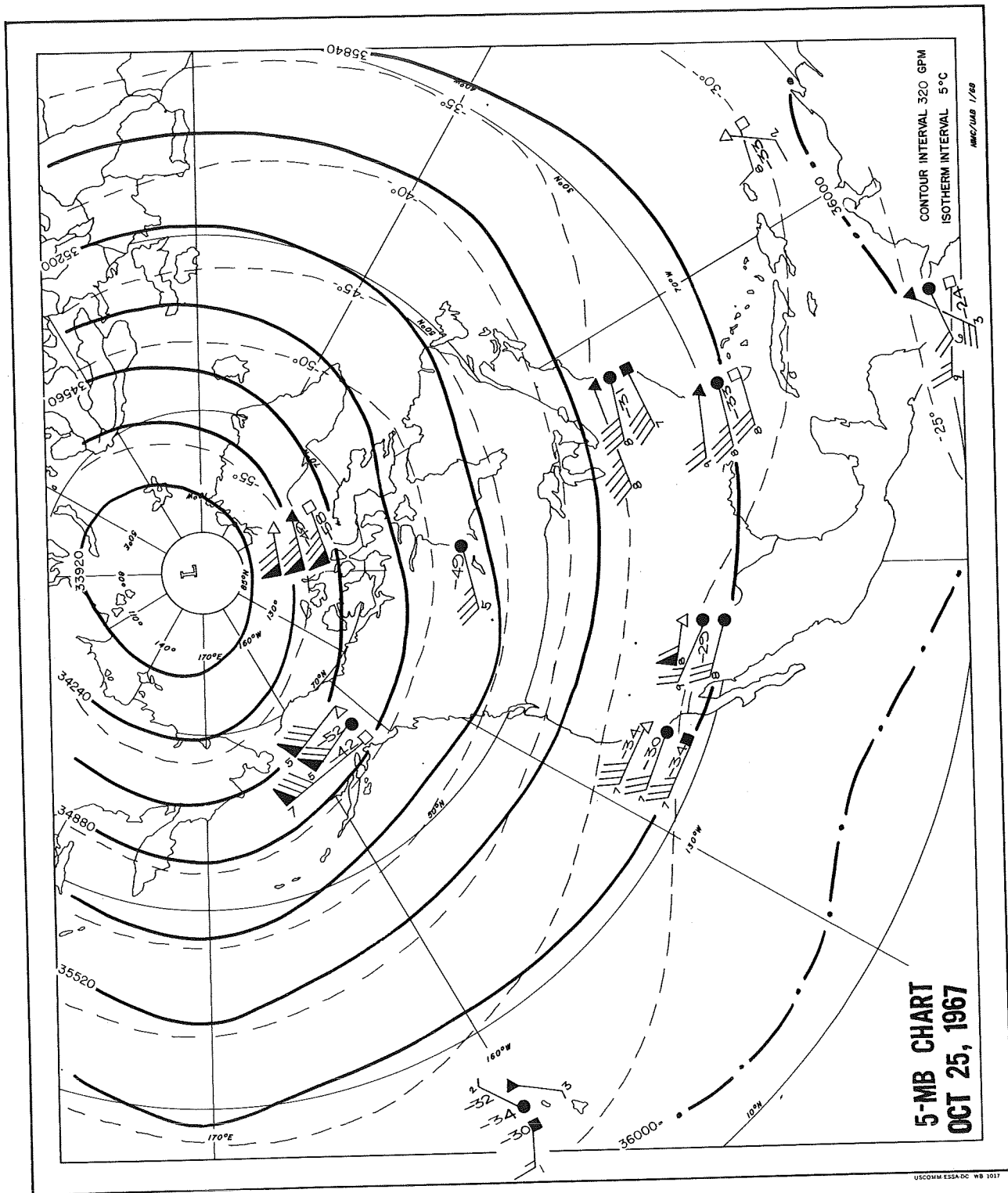
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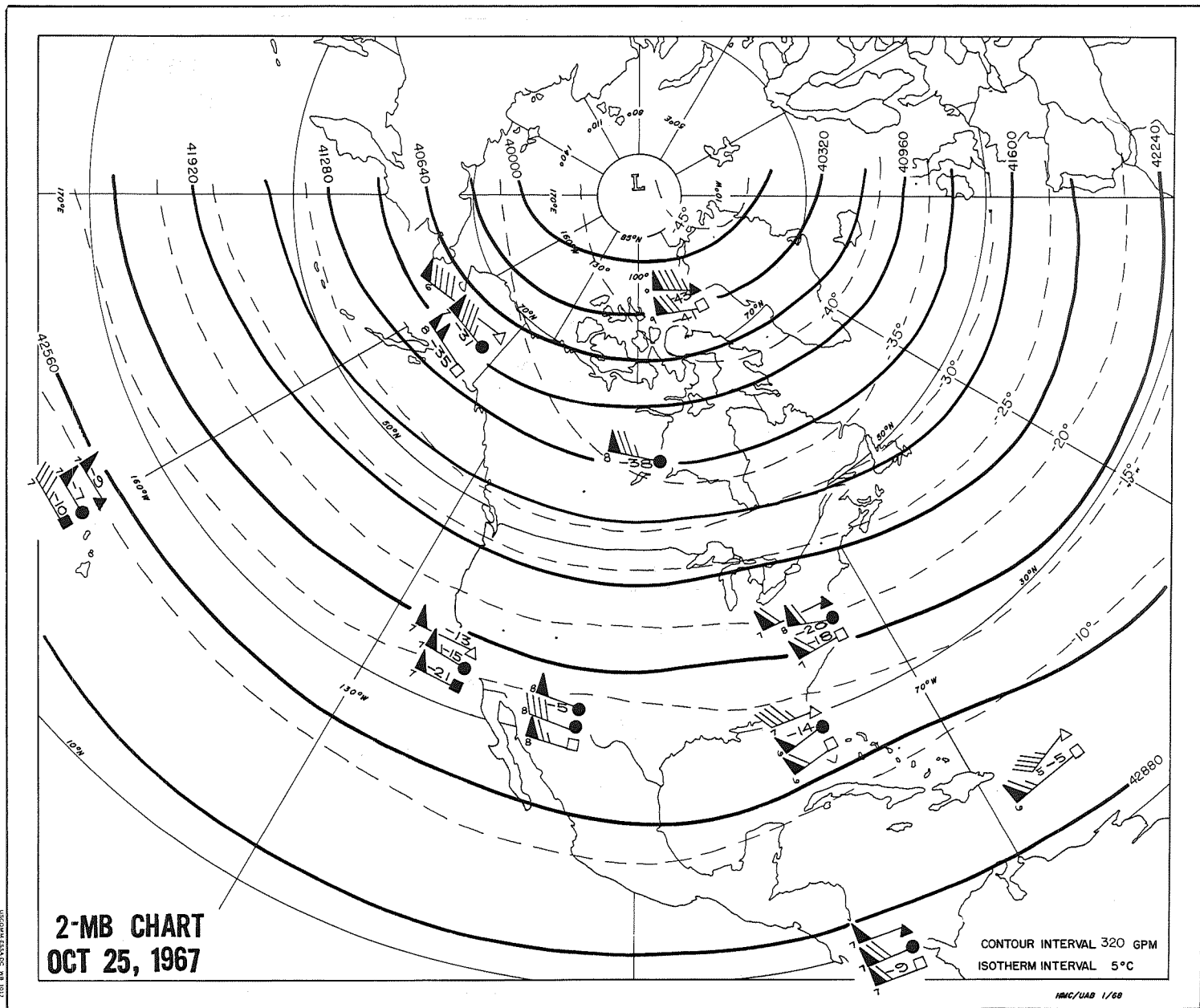


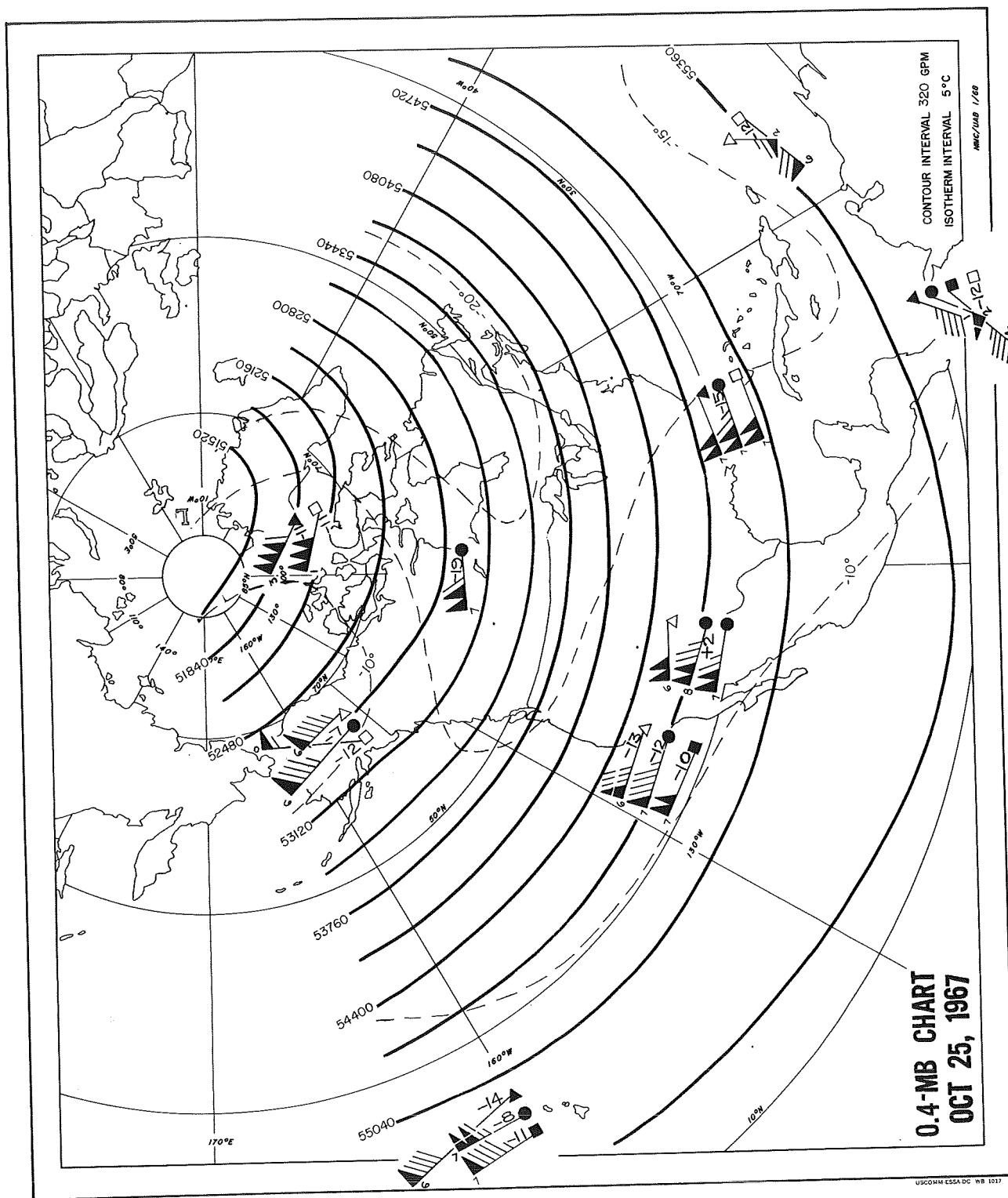
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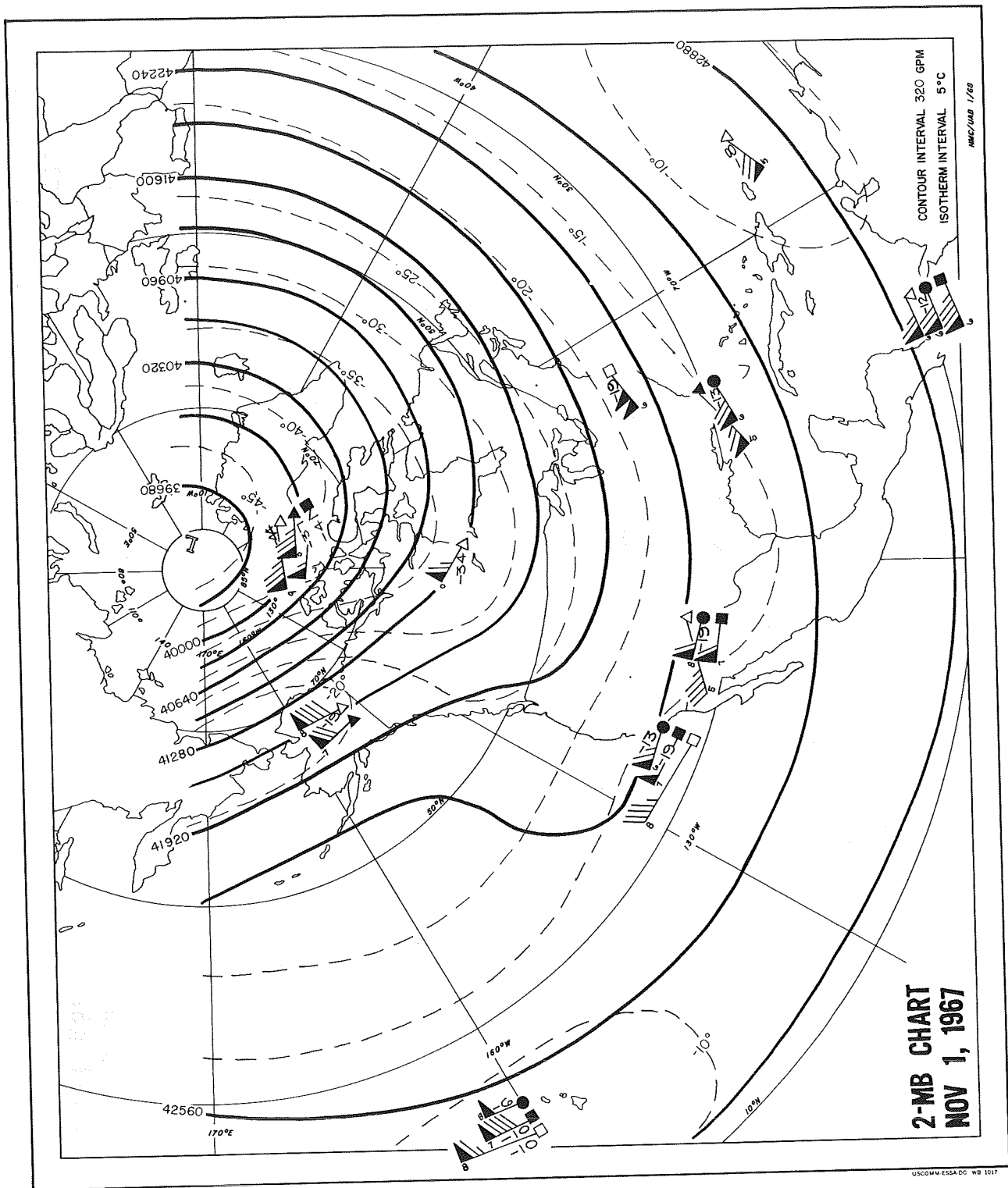


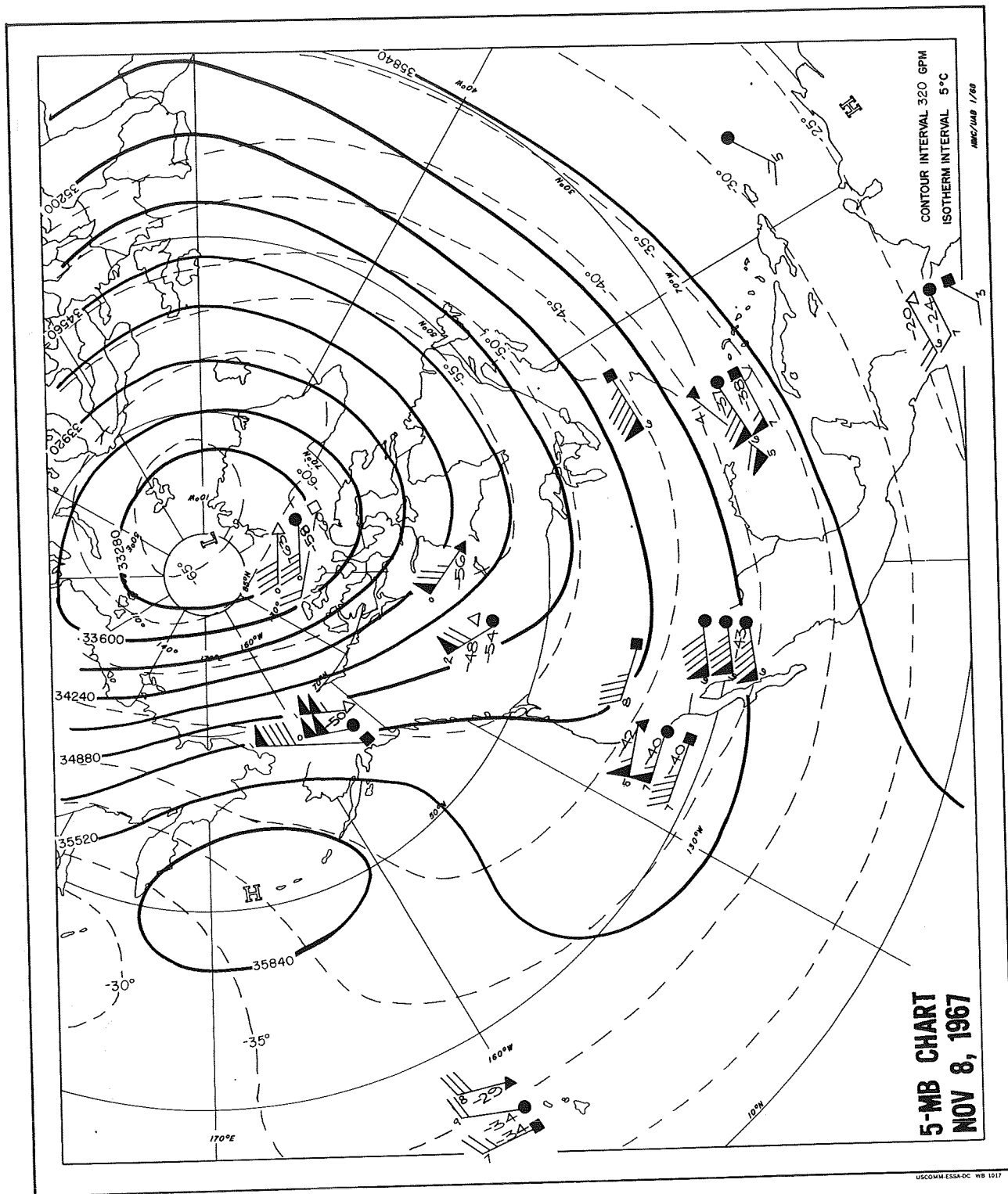
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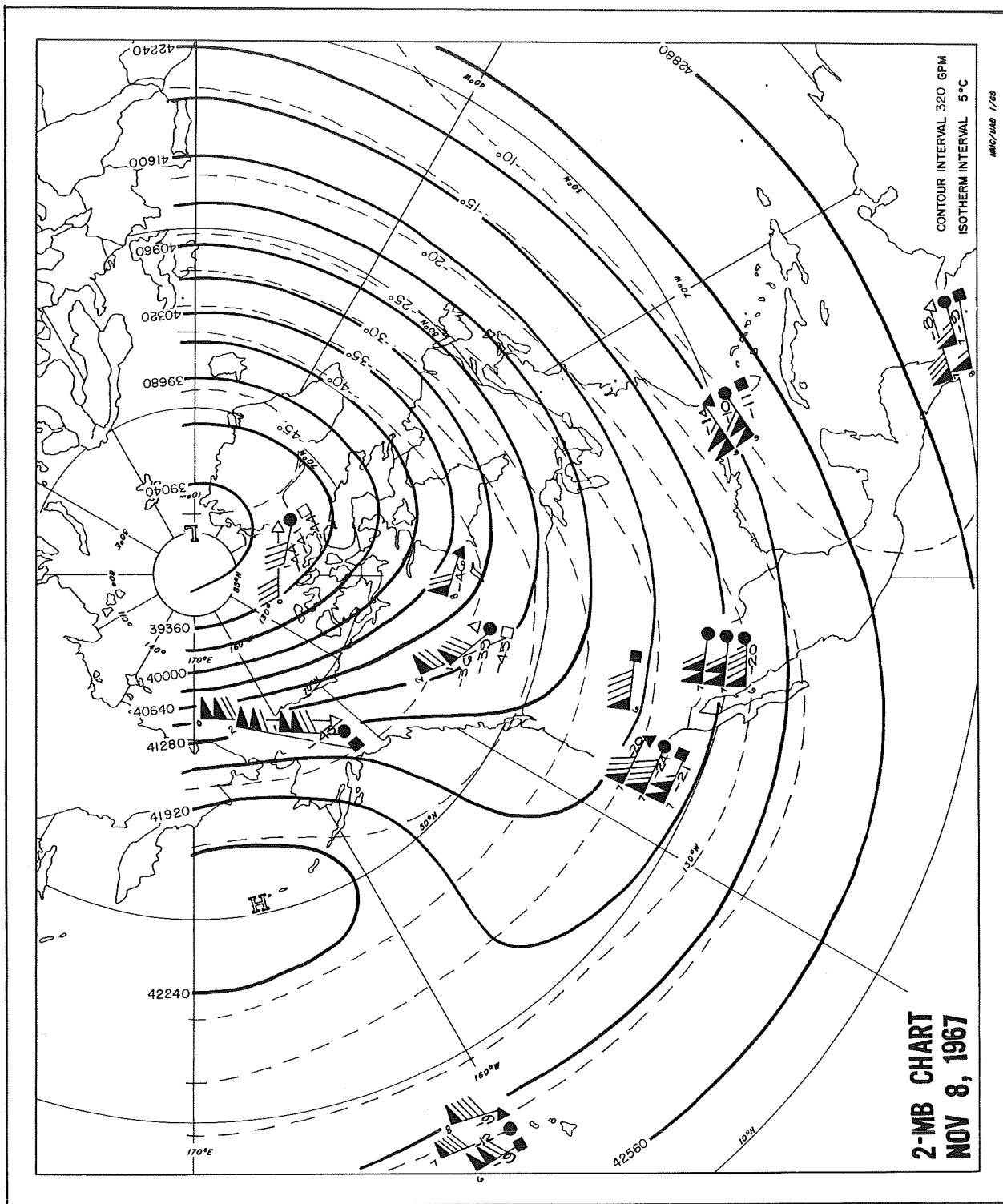


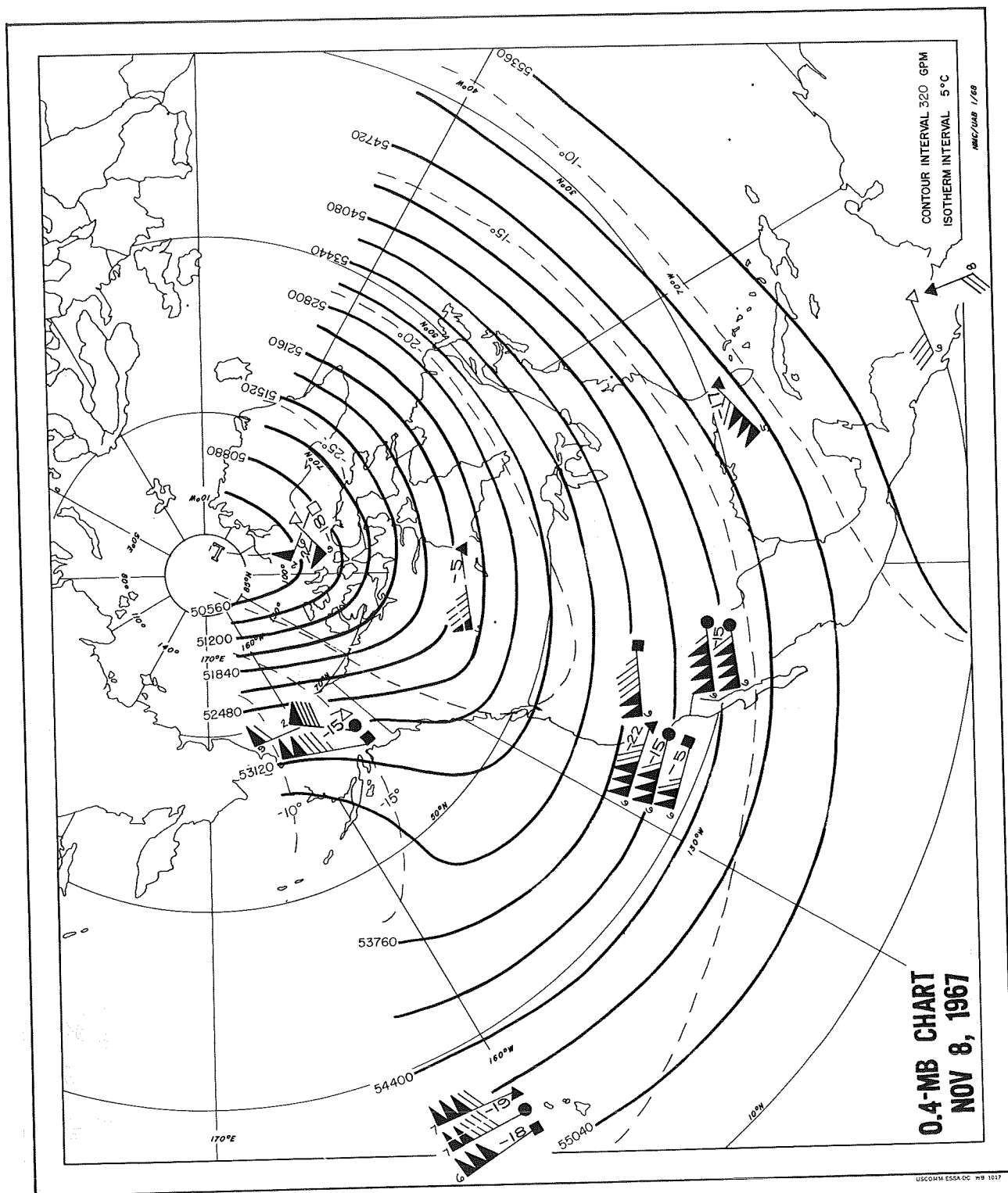


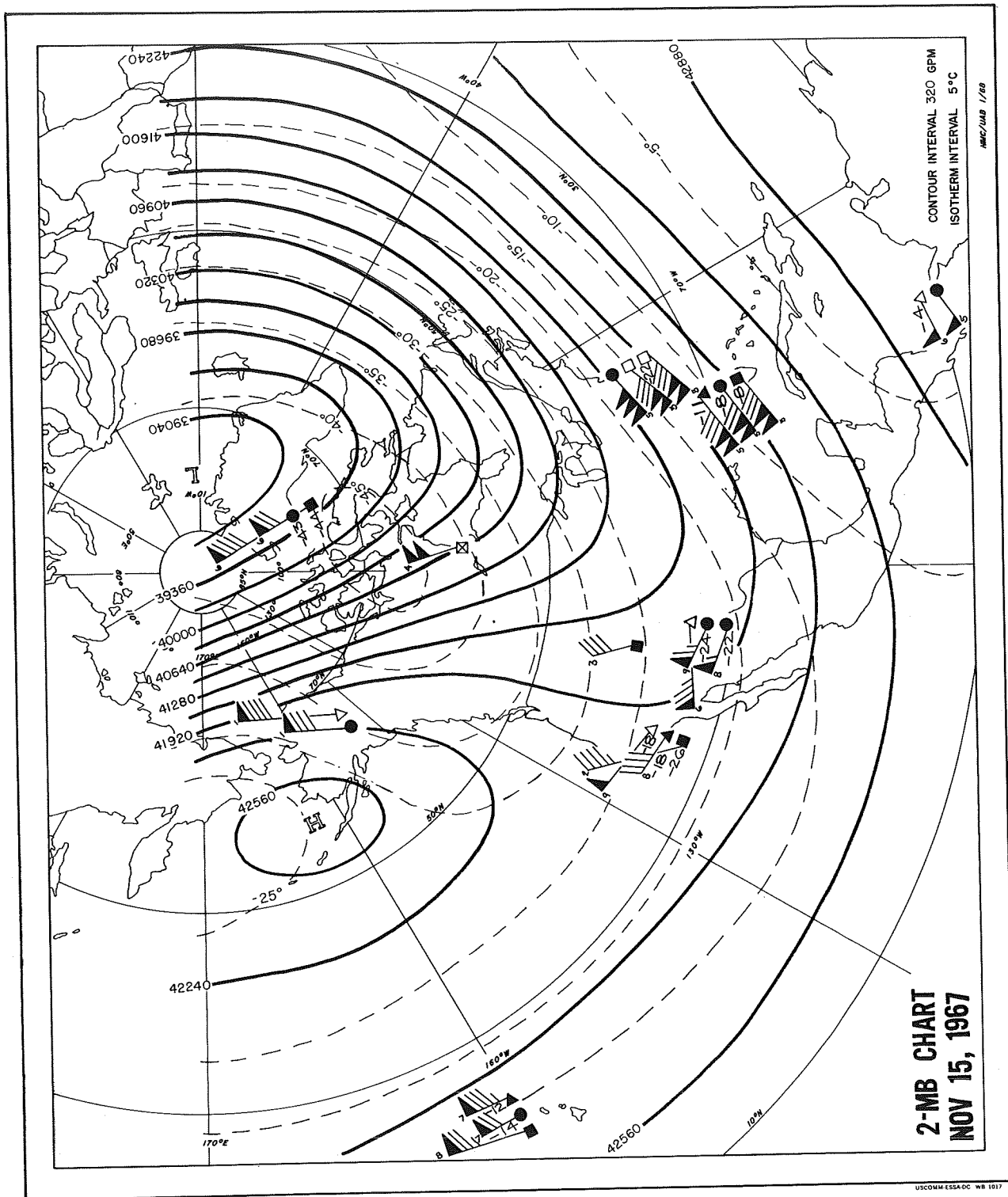


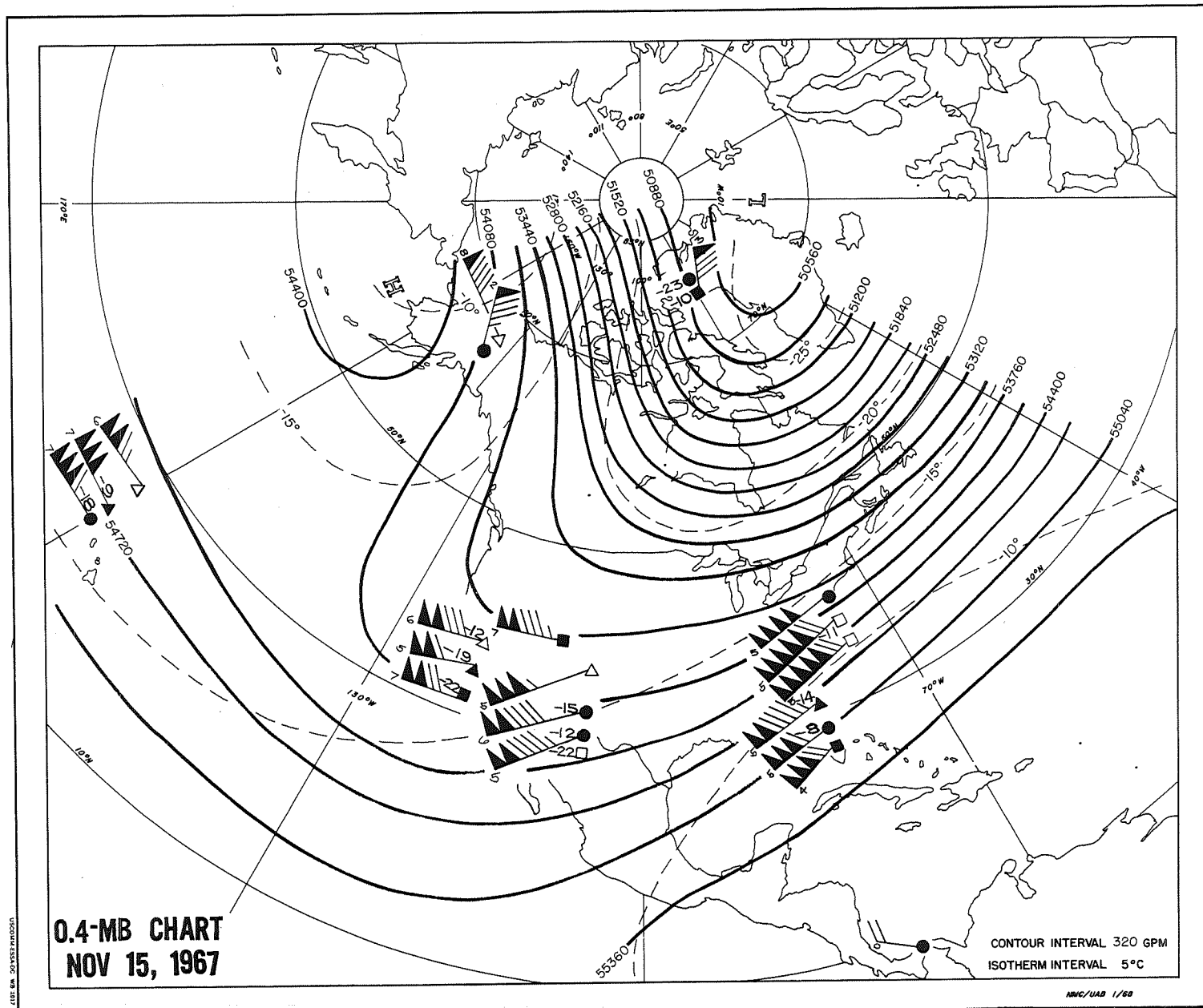


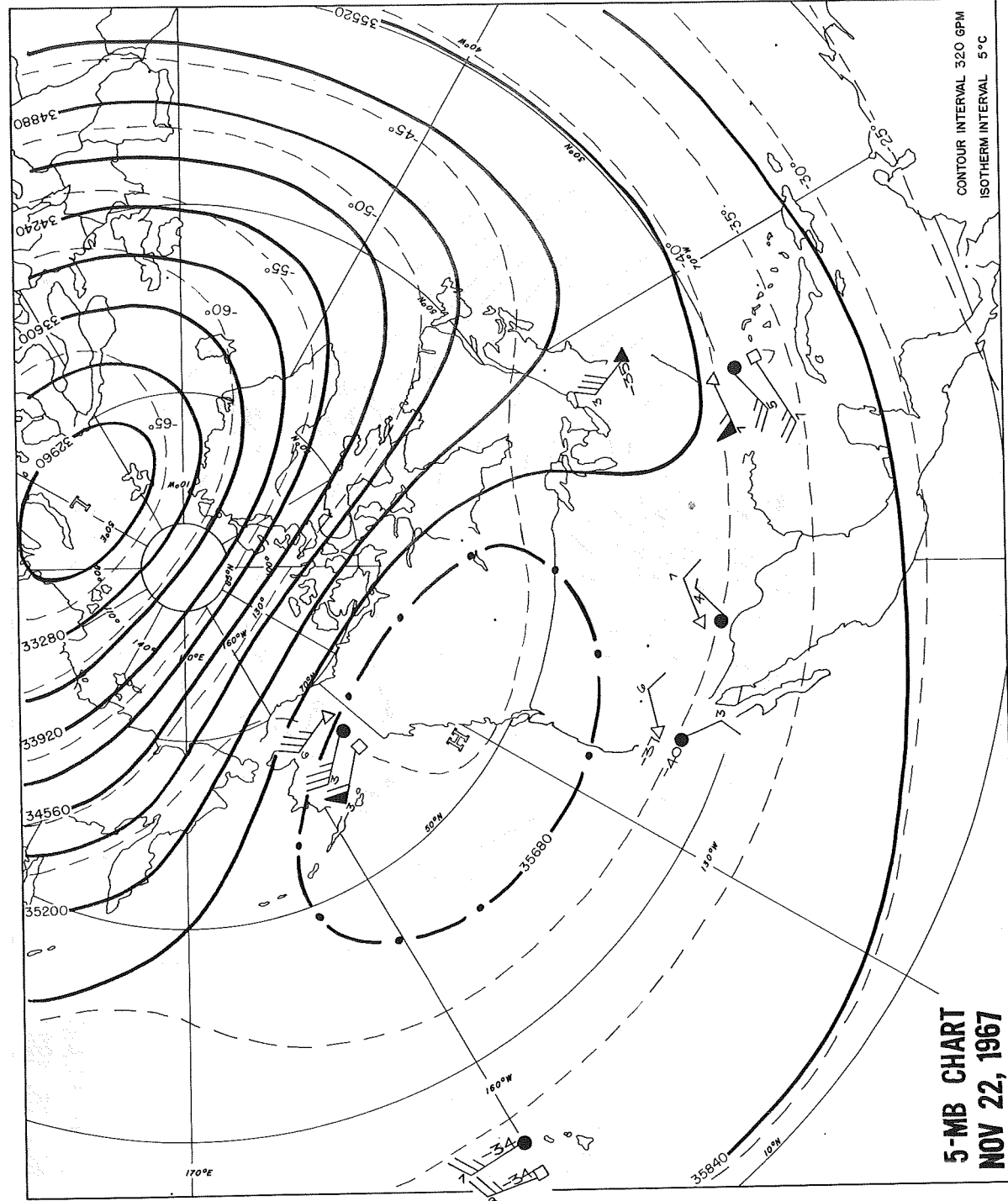




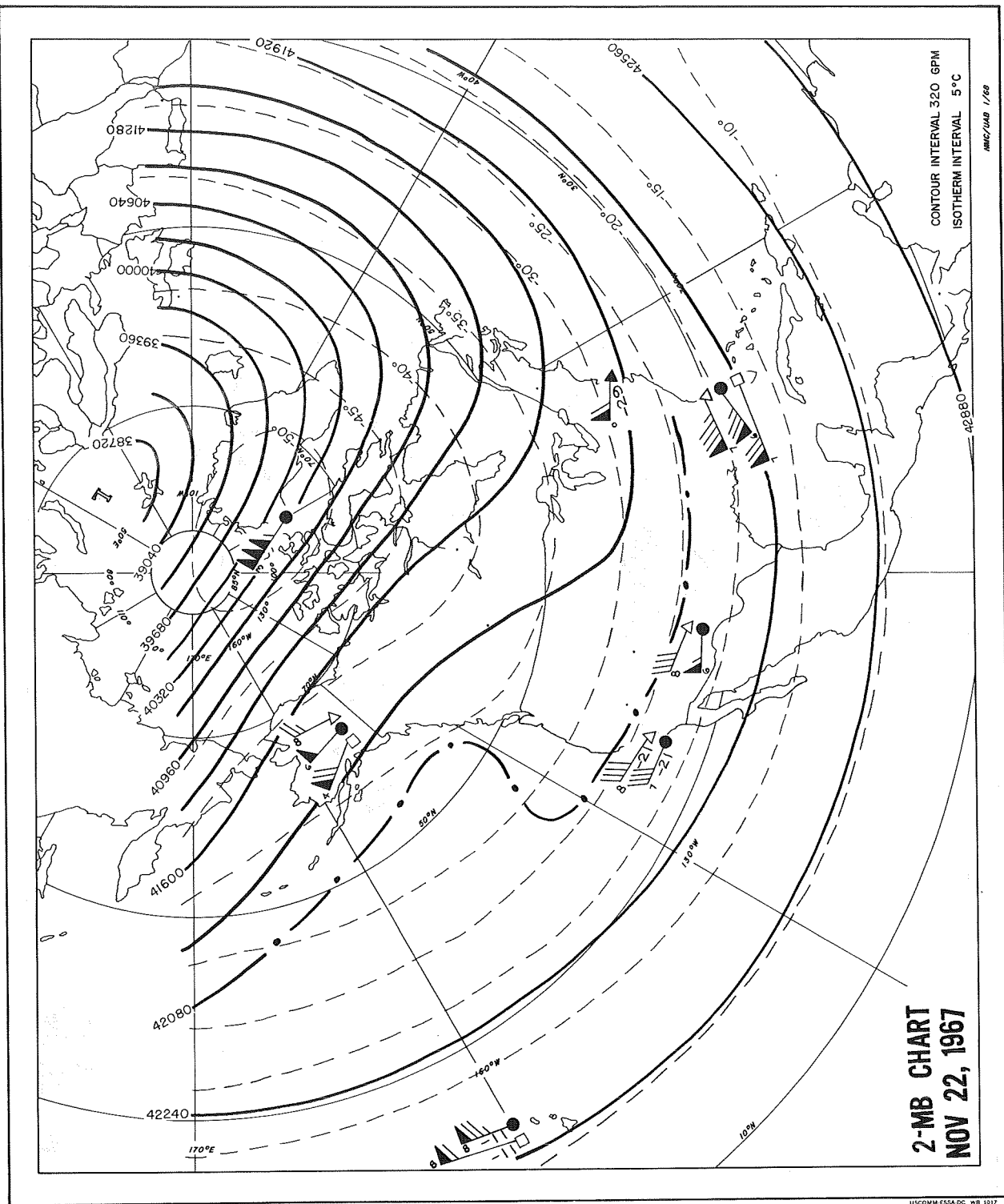


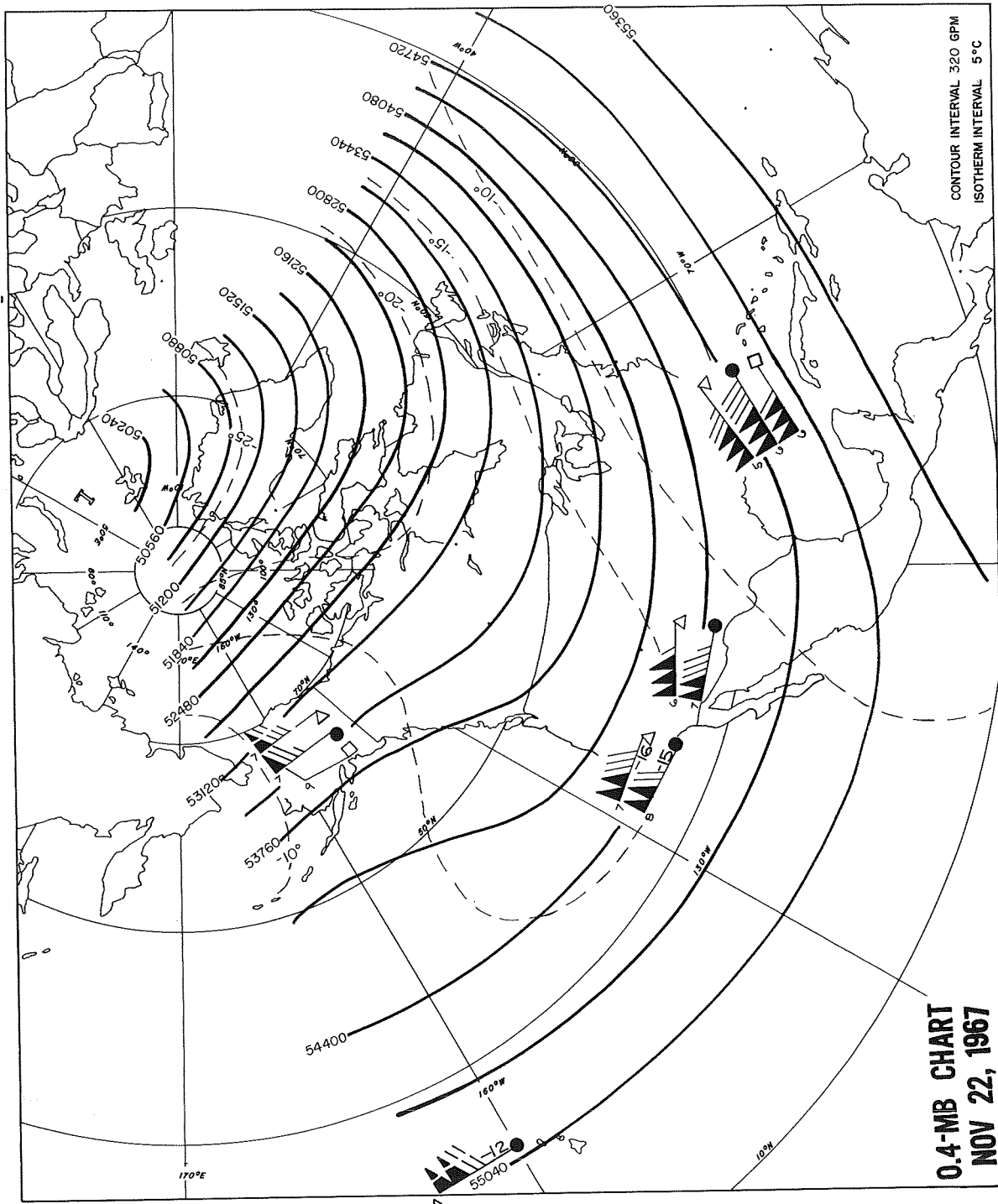






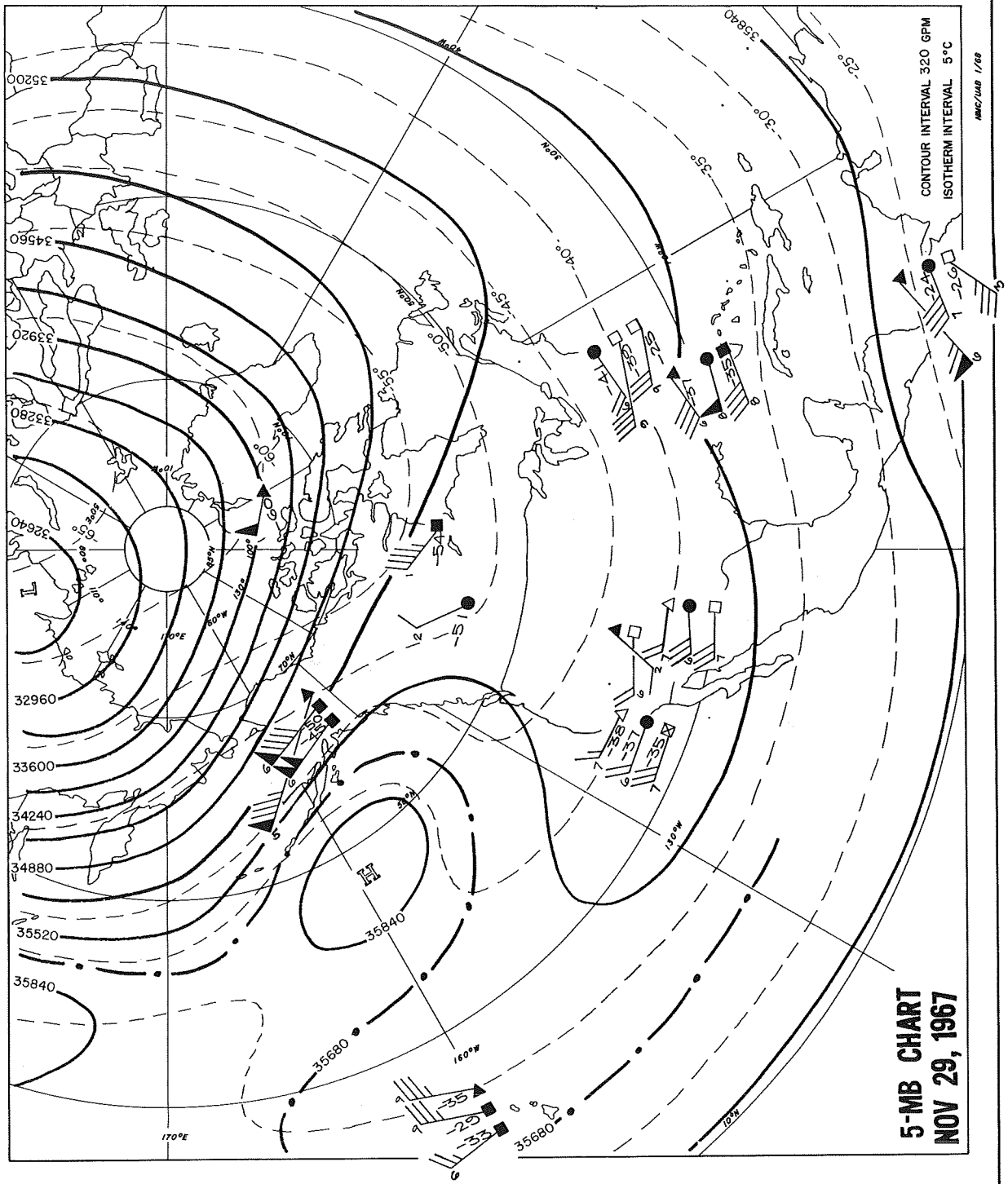
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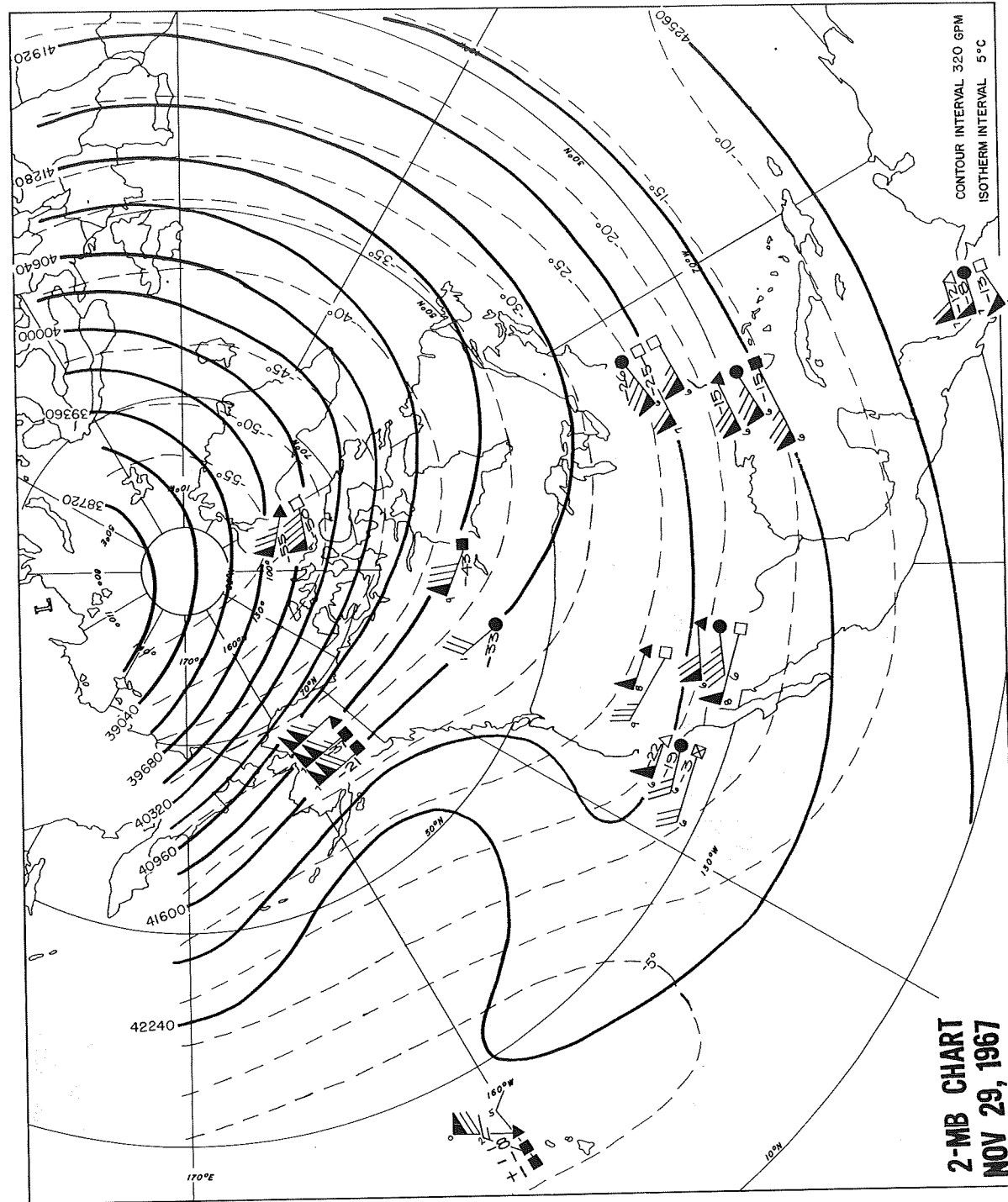




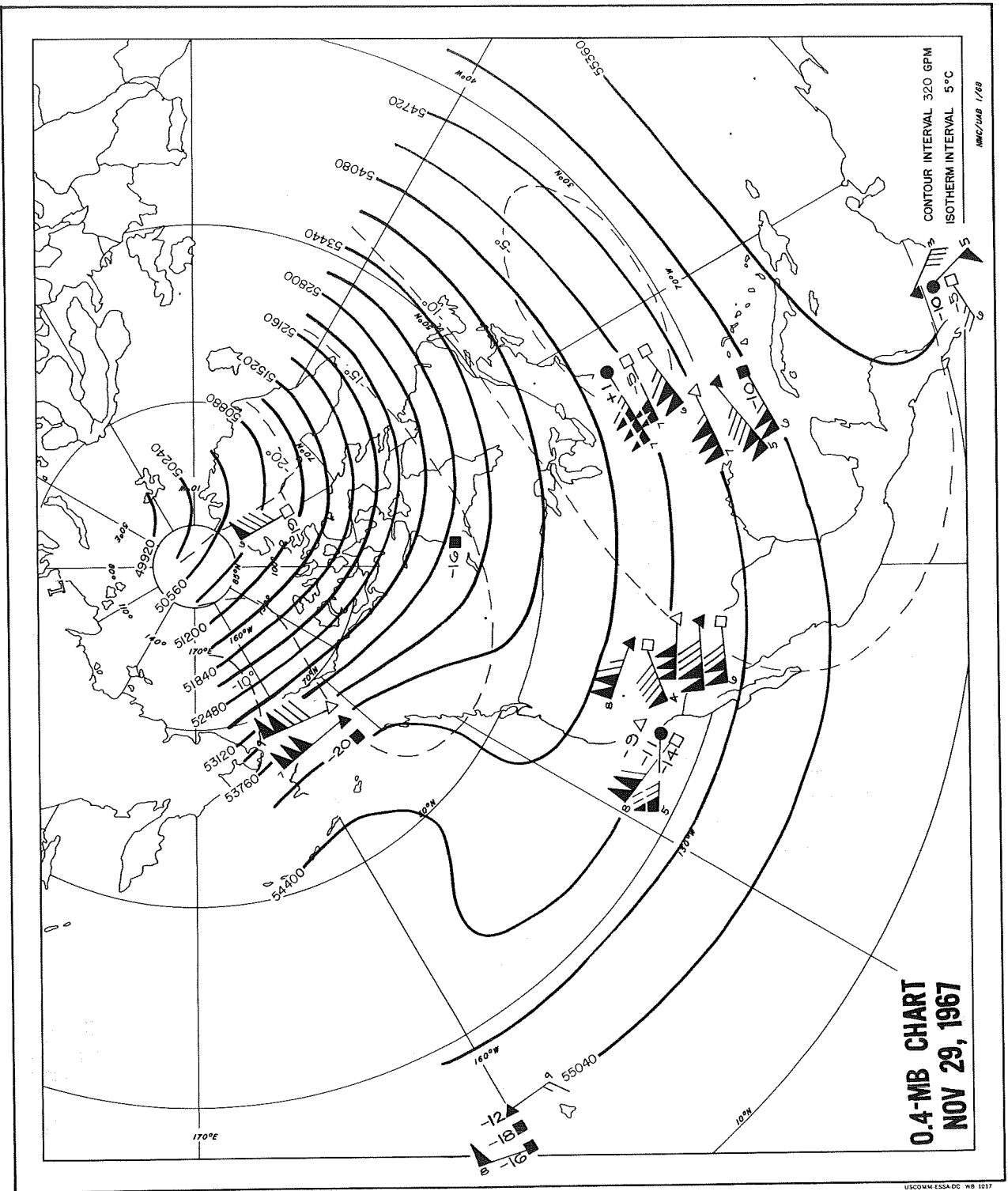
NMC/UMB 1/68

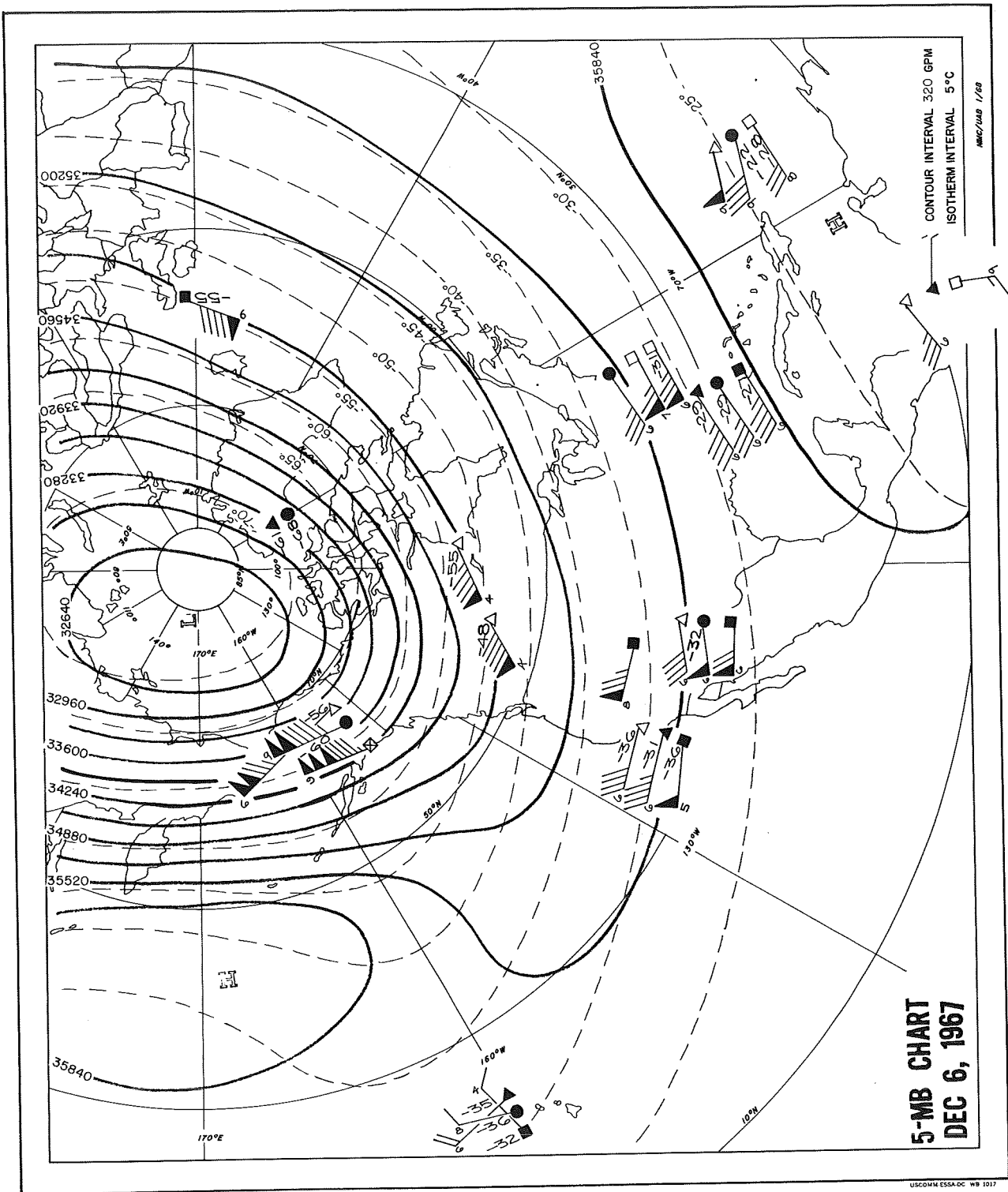
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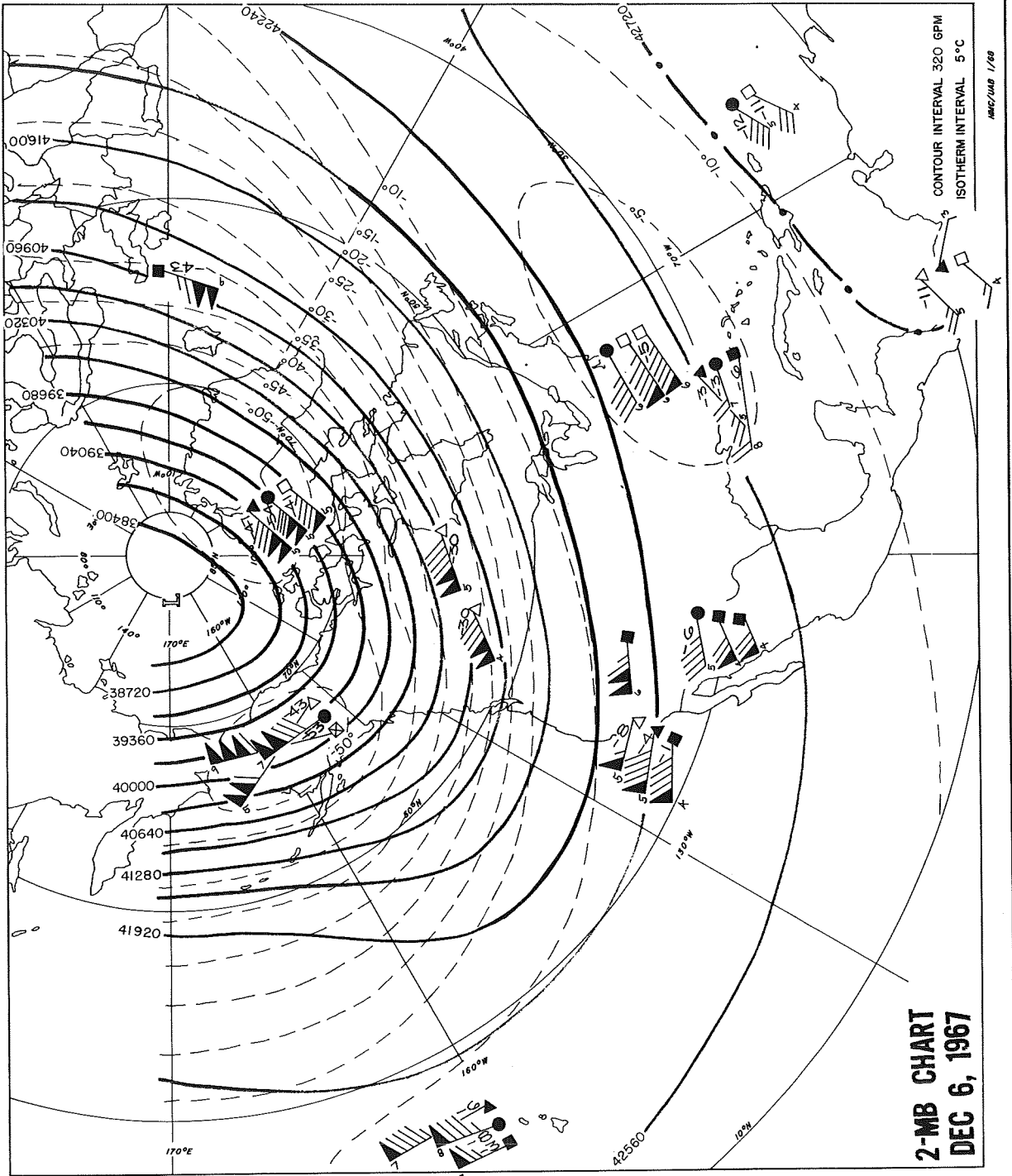




USCOMM ESSA DC WB 1917





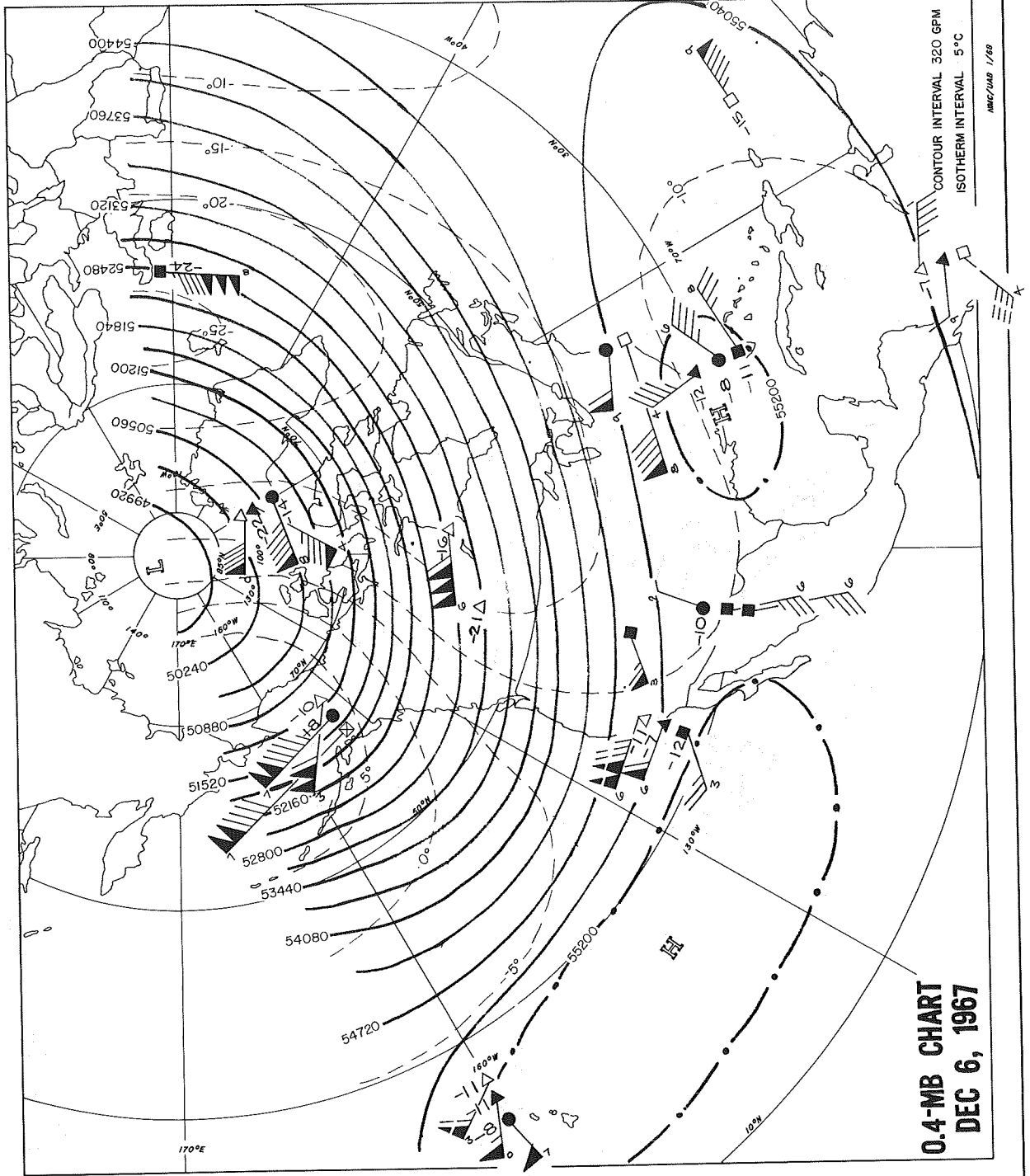


2-MB CHART
DEC 6, 1967

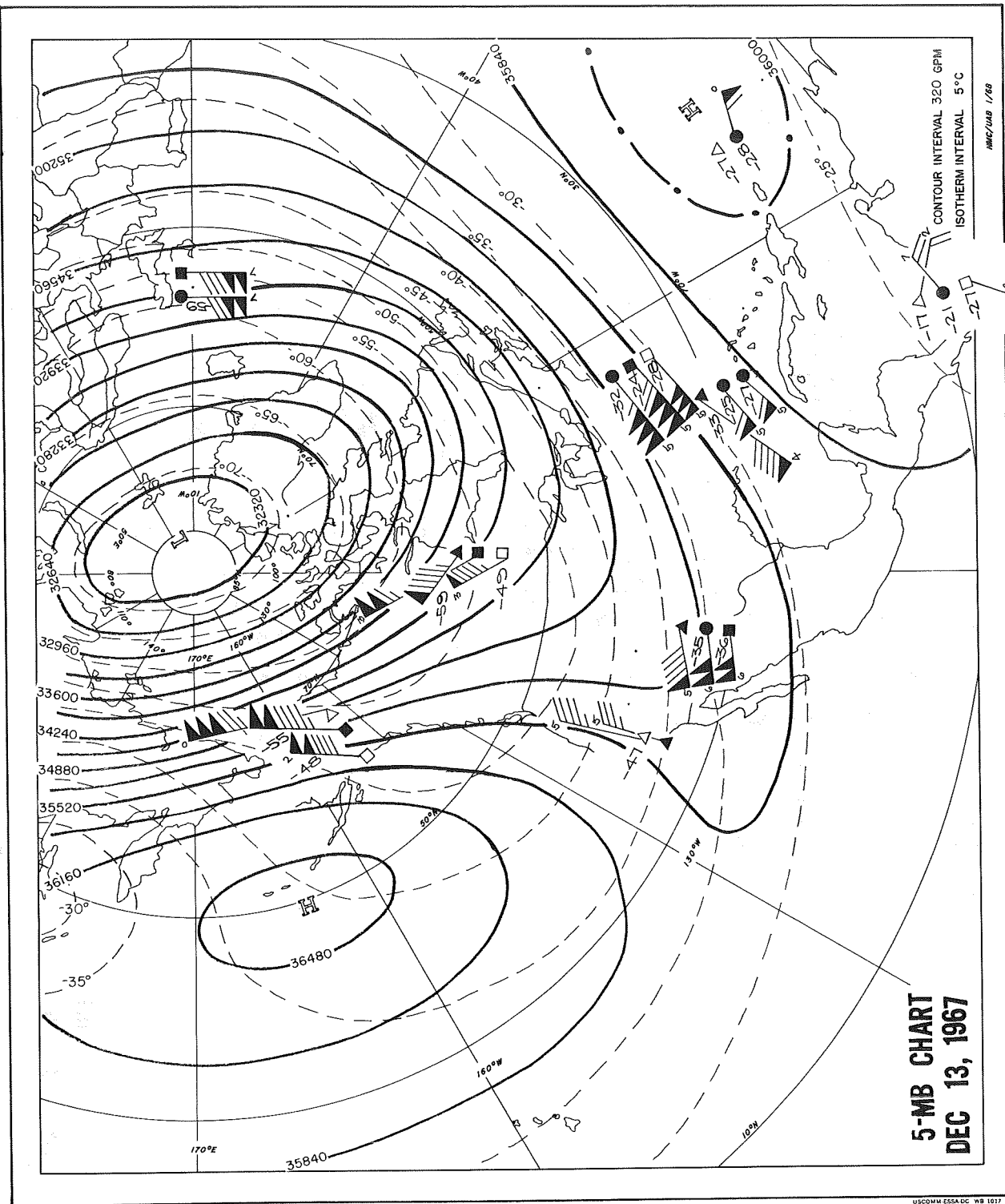
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ISOTHERM INTERVAL 5°C

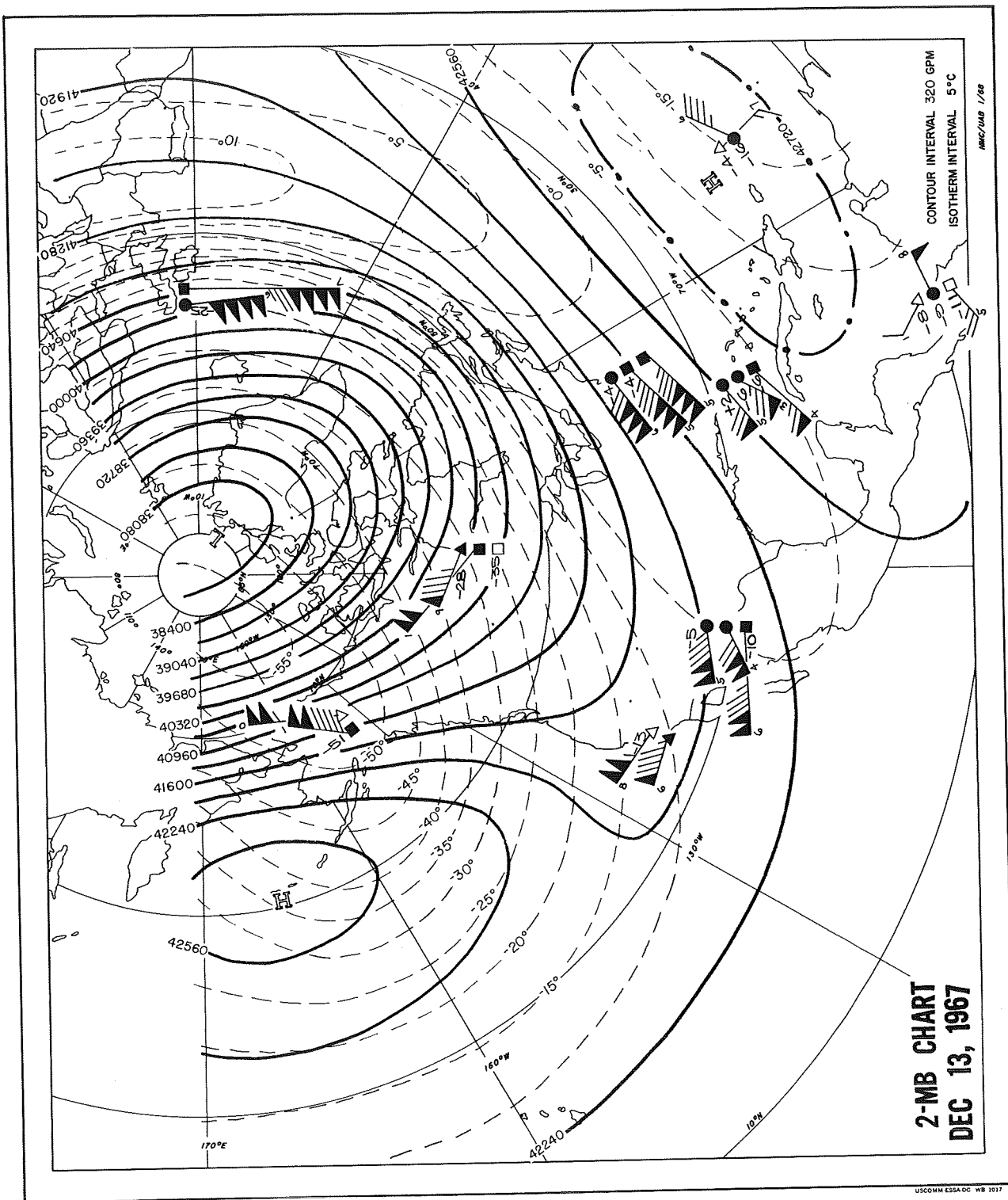
AMC/UMB 1/68

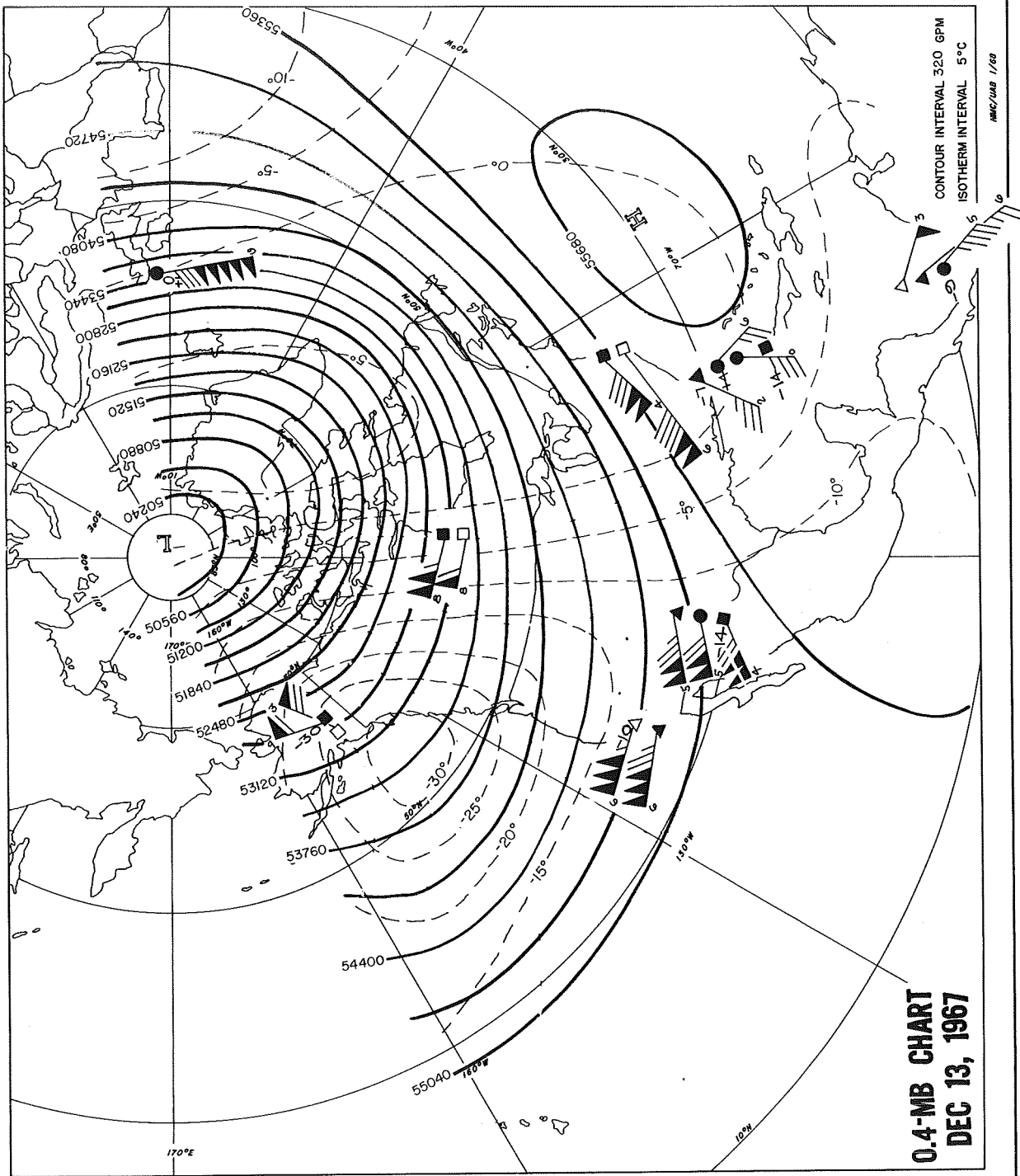
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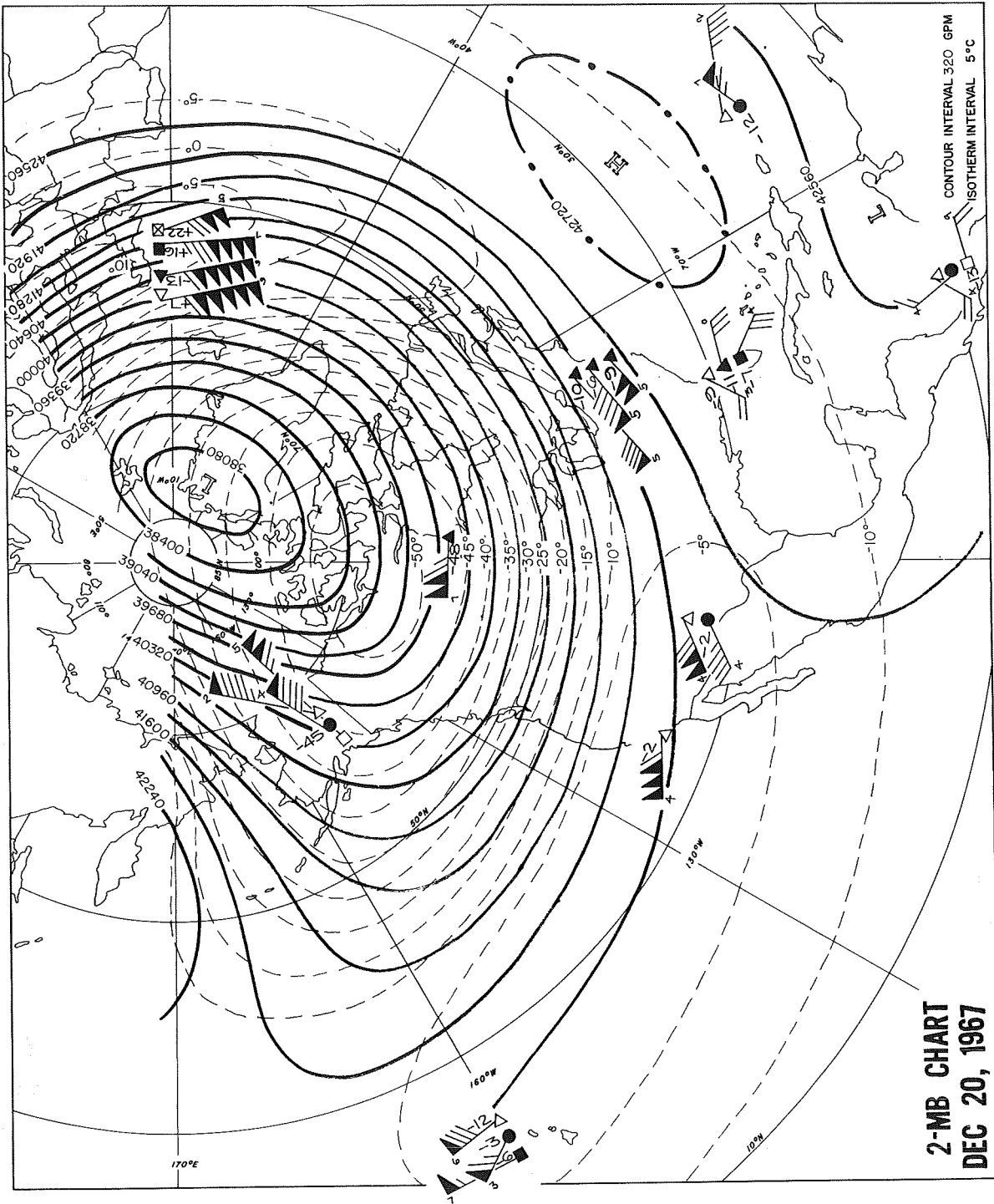
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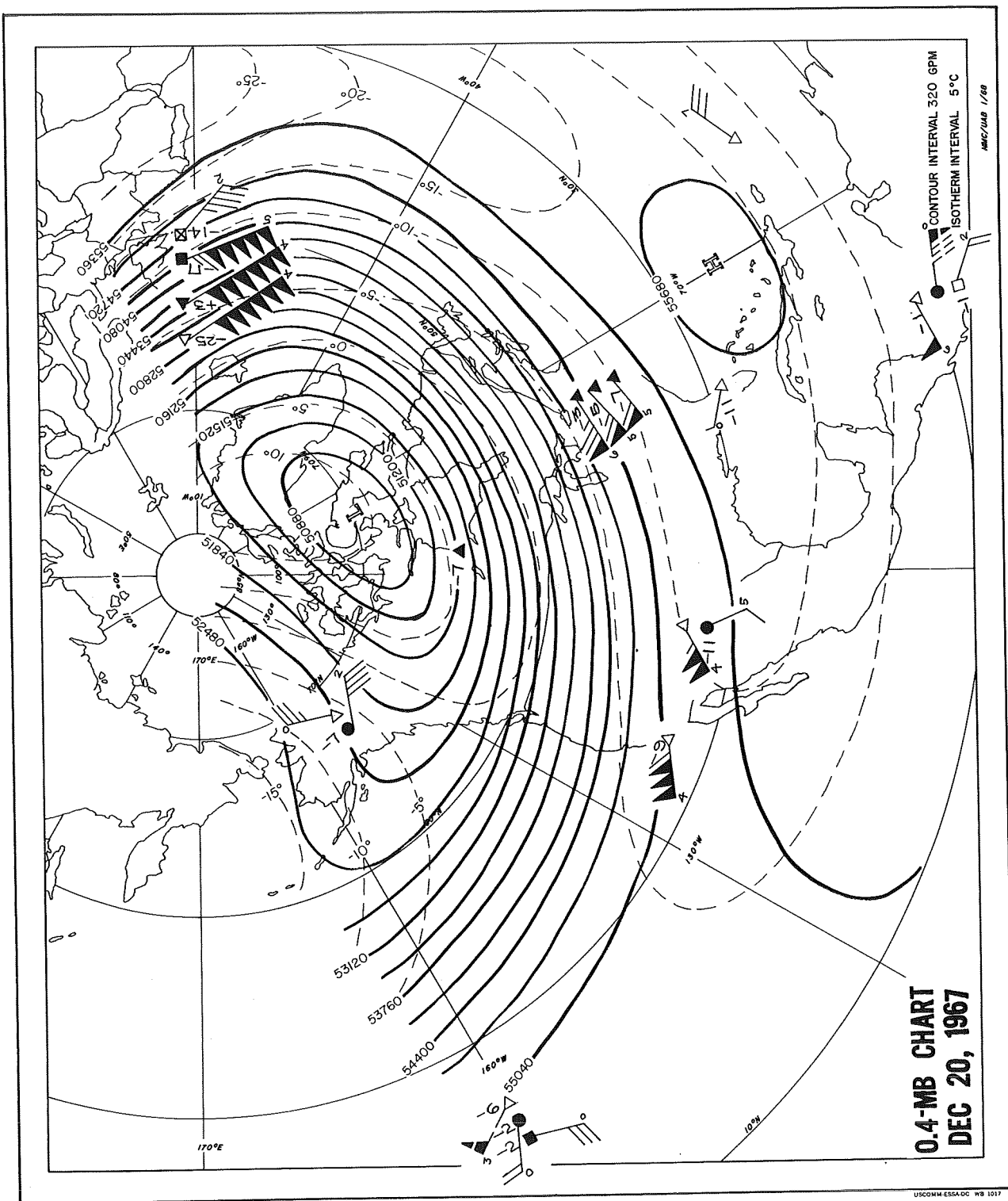


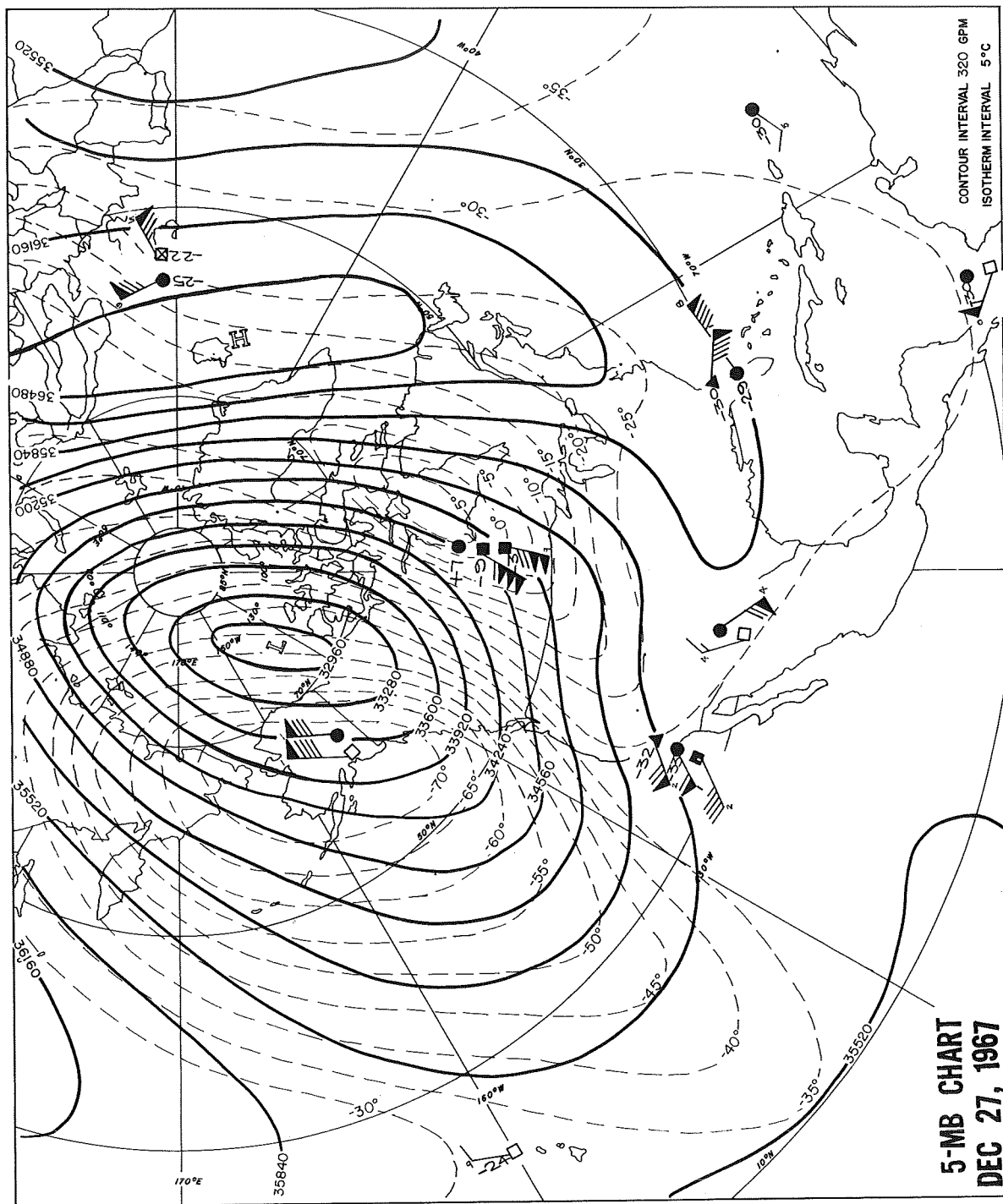




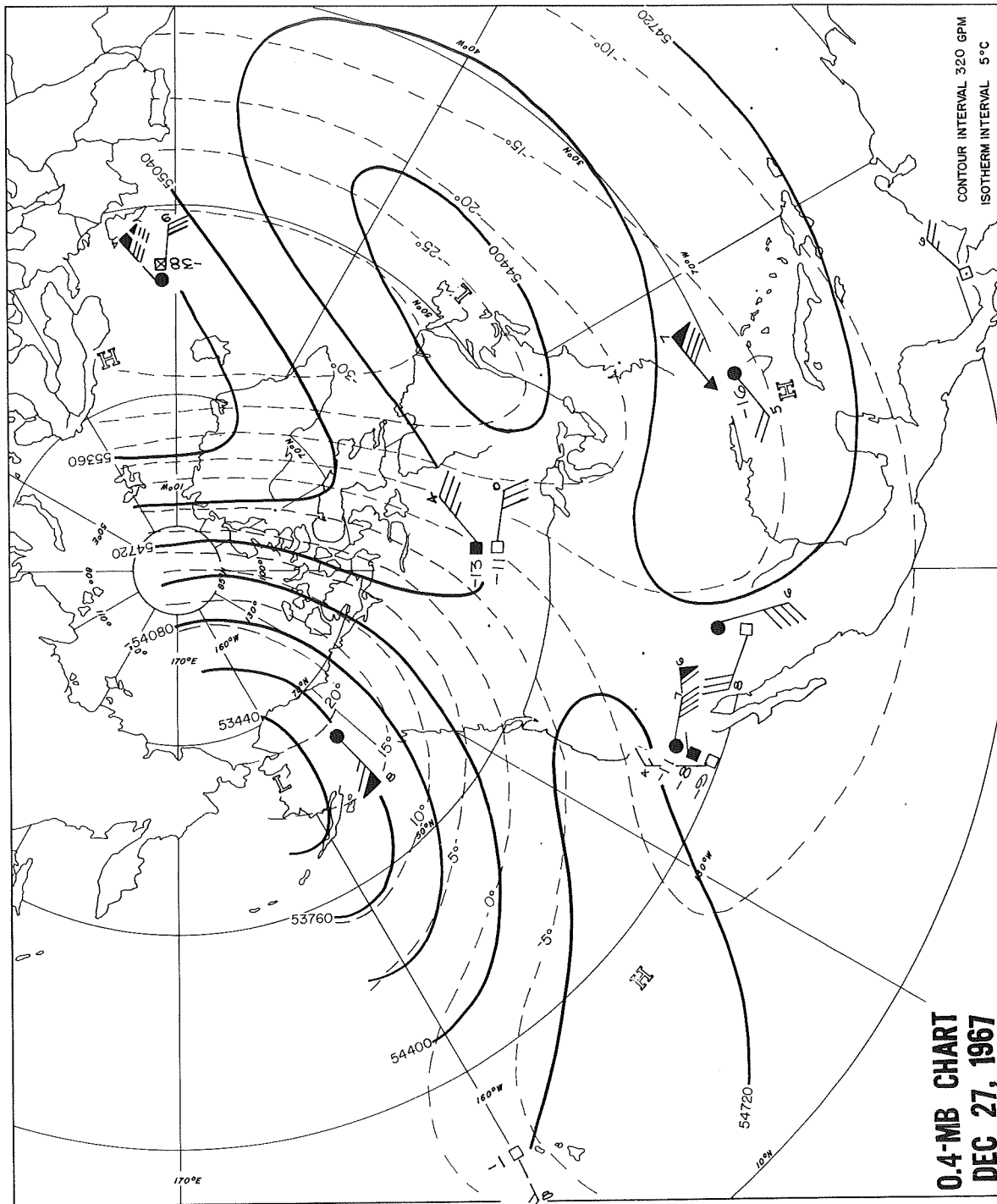
USCOMM ESSADC WB 1017







USCOMB ESSA DC WB 1027



0.4-MB CHART
DEC 27, 1967

USCOMM ESSA DC WB 1017

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